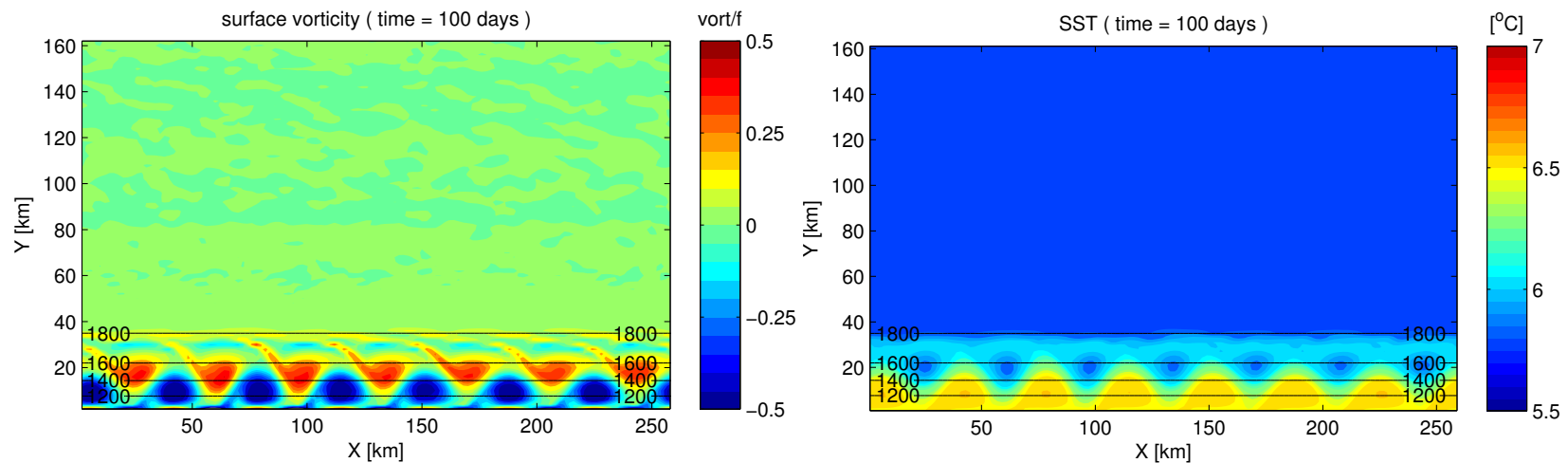


Meso and sub meso scale dynamics of coastal currents along a steep shelf bathymetry



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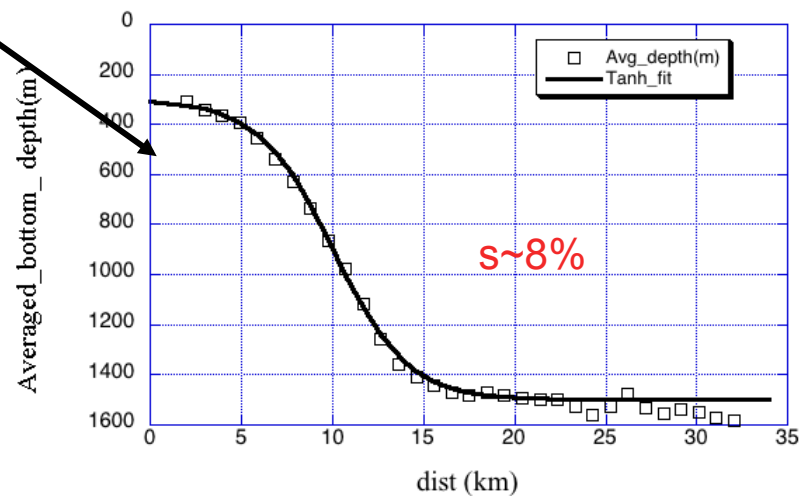
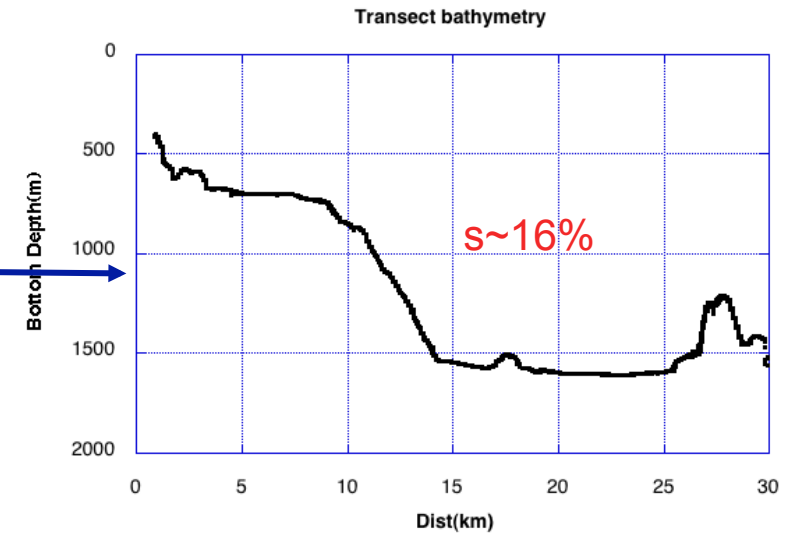
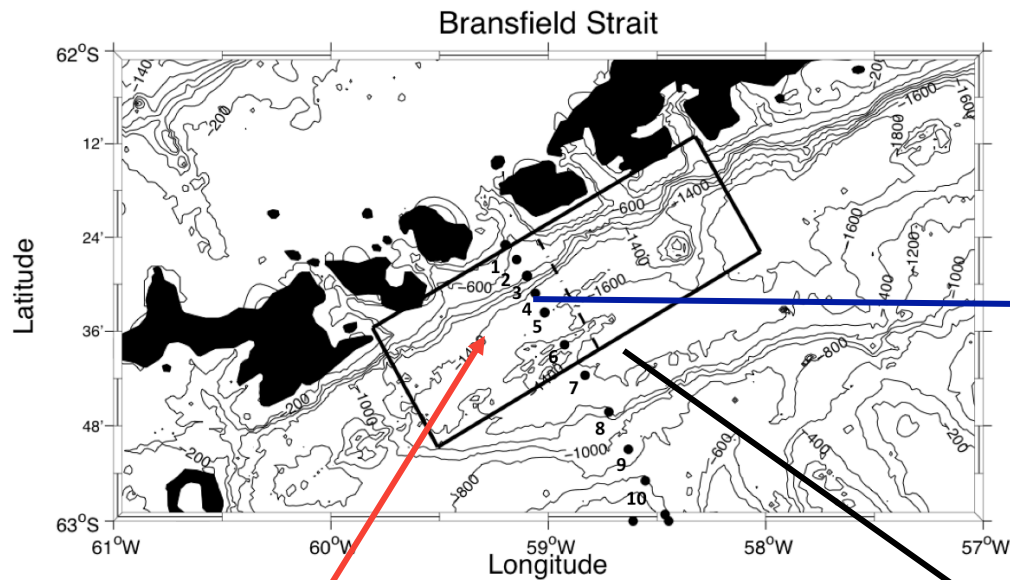
Meso and sub meso scale dynamics of coastal currents...



Motivations: $1/36^\circ$ Mediterranean Sea (*MED36 runs*)

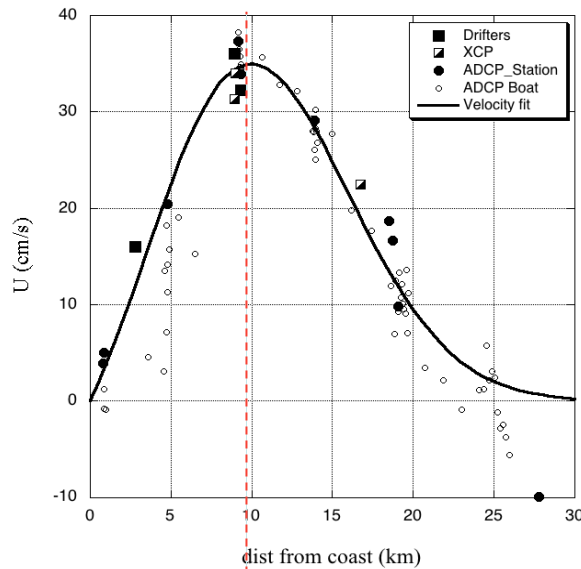


Buoyant coastal current : Bransfield strait bathymetry (Antarctica)



Buoyant coastal current : stable Bransfield current (Antarctica)

VELOCITY PROFILE

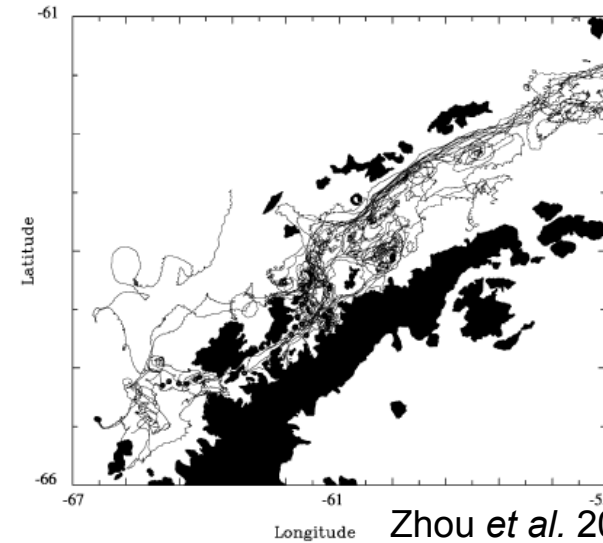


$$Ro = \frac{U}{fL} \approx 0.2 - 0.3$$

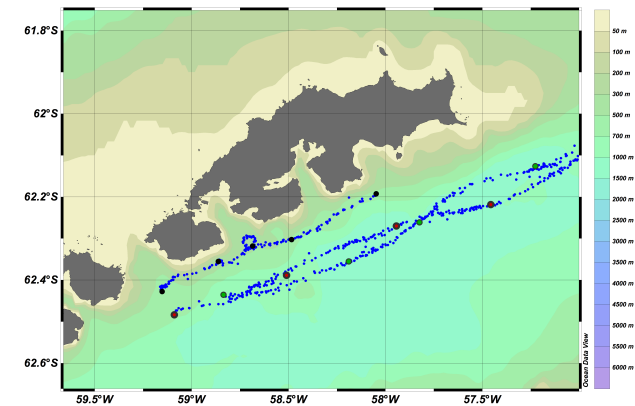
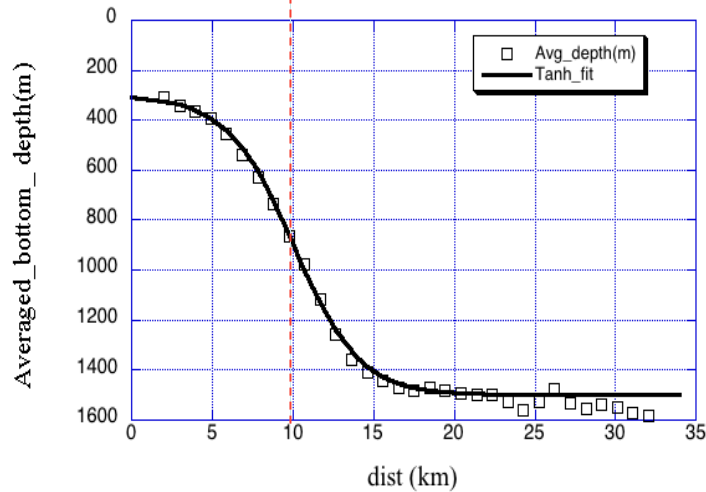
$$Bu = \left(\frac{R_d}{L} \right)^2 \approx 1$$

GEOSTROPHIC CURRENT

DRIFTERS TRAJECTORIES



Zhou et al. 2002

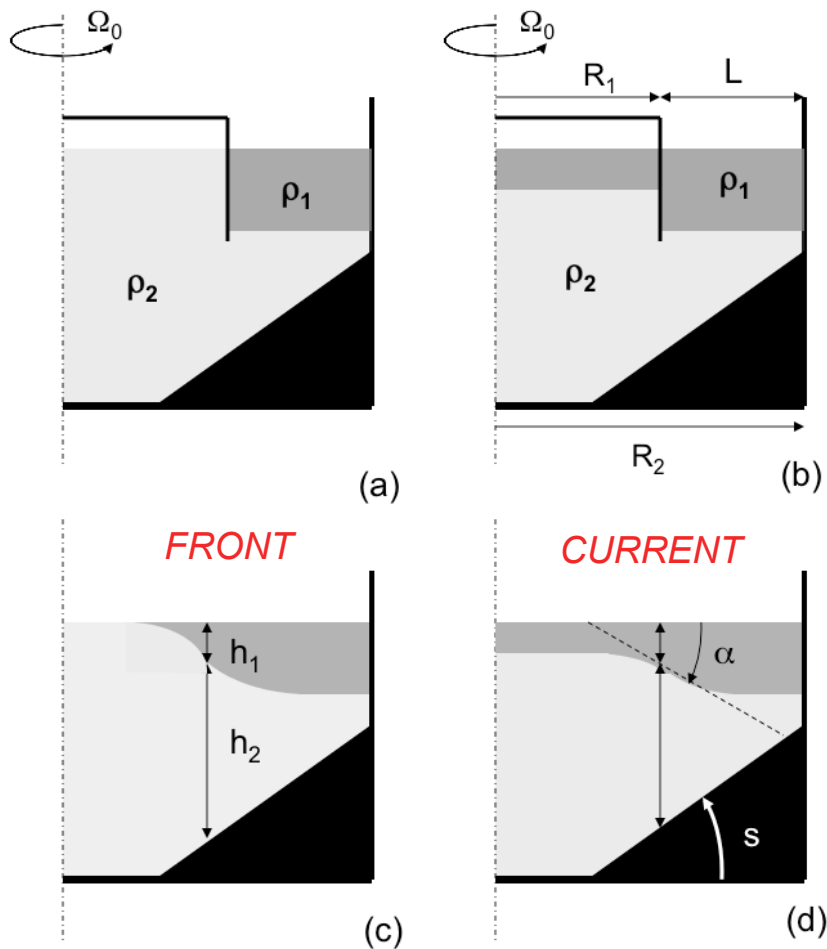


Poulin et al. 2014

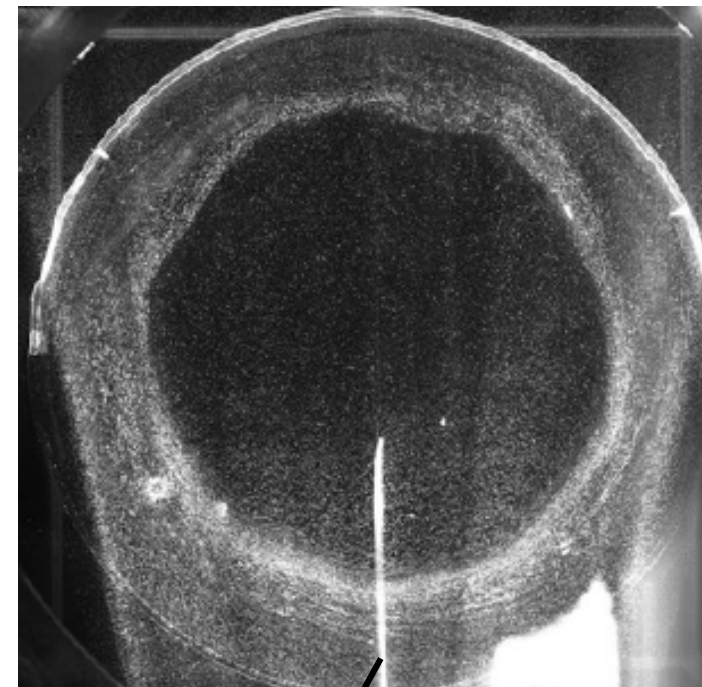
What do we learn from laboratory experiments ?



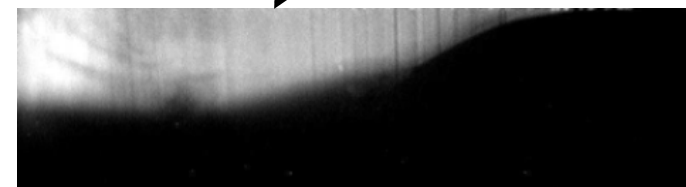
Idealized configuration: experimental setup



*Initial and adjusted configurations
side view of the two layer salt stratifications*



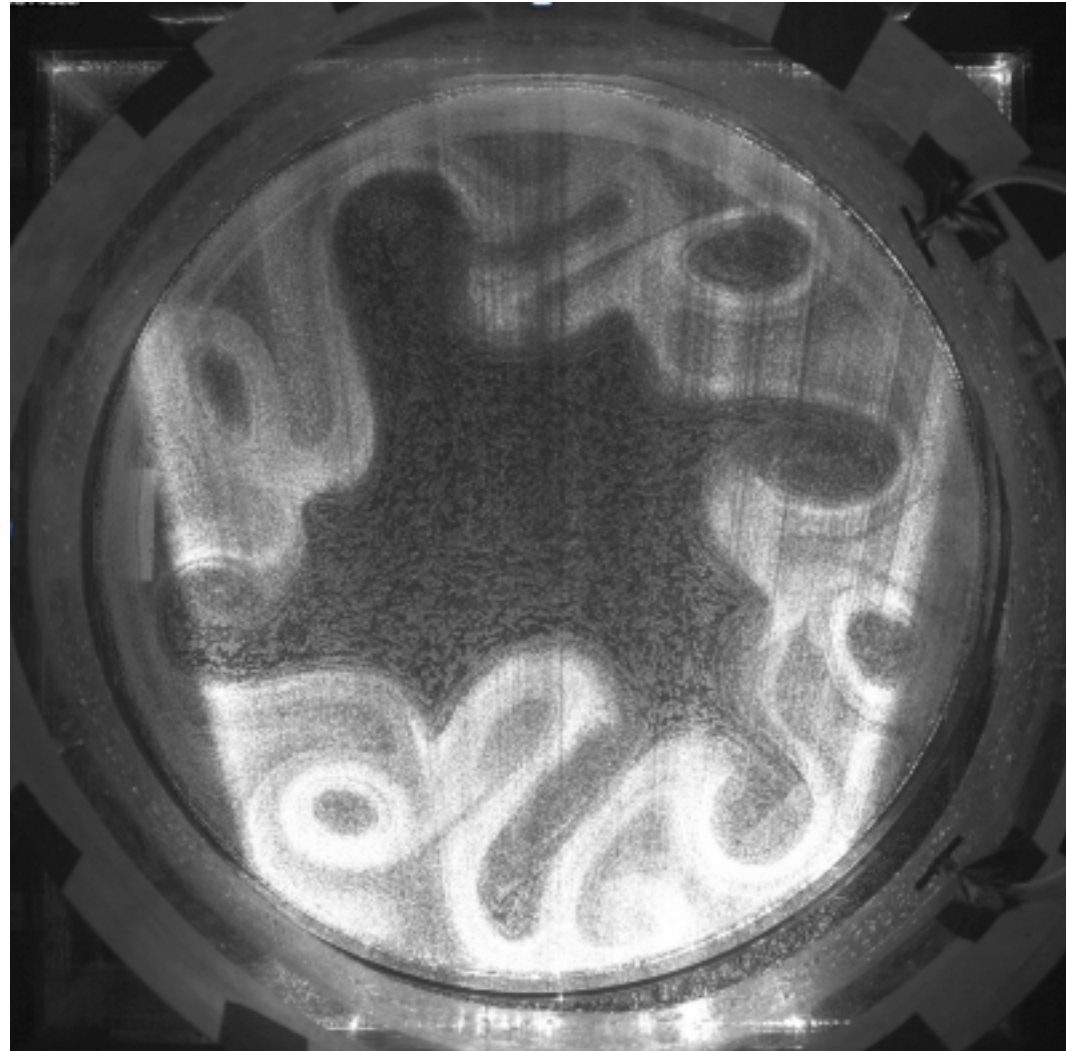
top view with PIV particles



side view LIF visualization



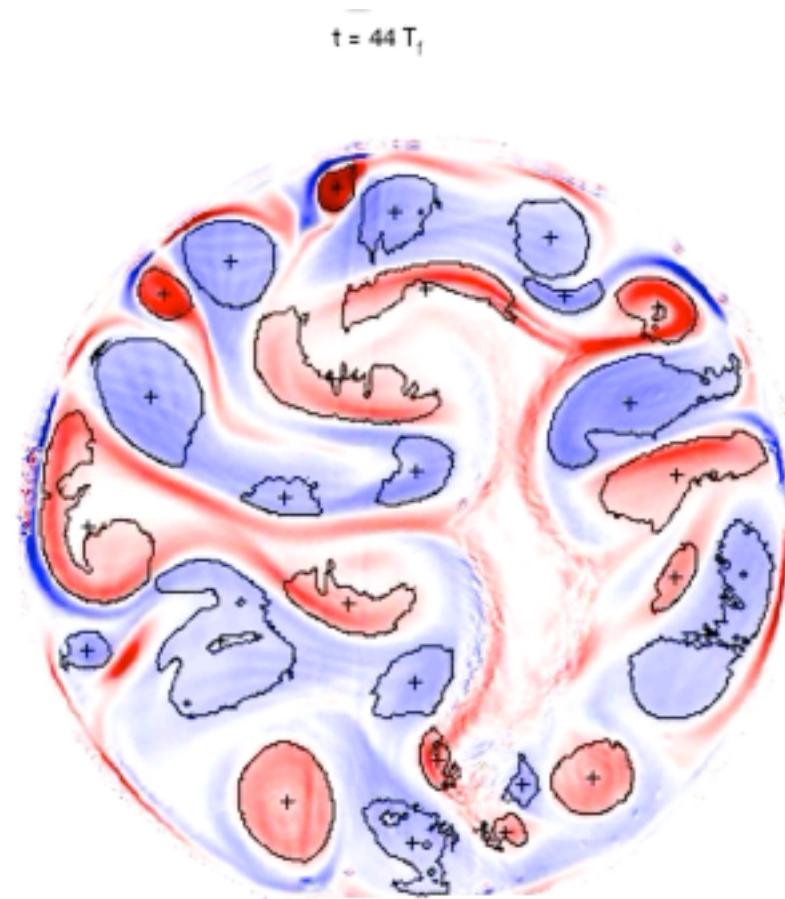
High resolution
PIV measurements:
4800x3200 pixels camera



COASTAL FRONT
FLAT BOTTOM CASE

High resolution
PIV measurements:
surface vorticity

blue: anticyclonic
red: cyclonic



COASTAL FRONT
FLAT BOTTOM CASE

Black contours Okubo-Weiss criterion
(Isern-Fontanet et al. 2003)

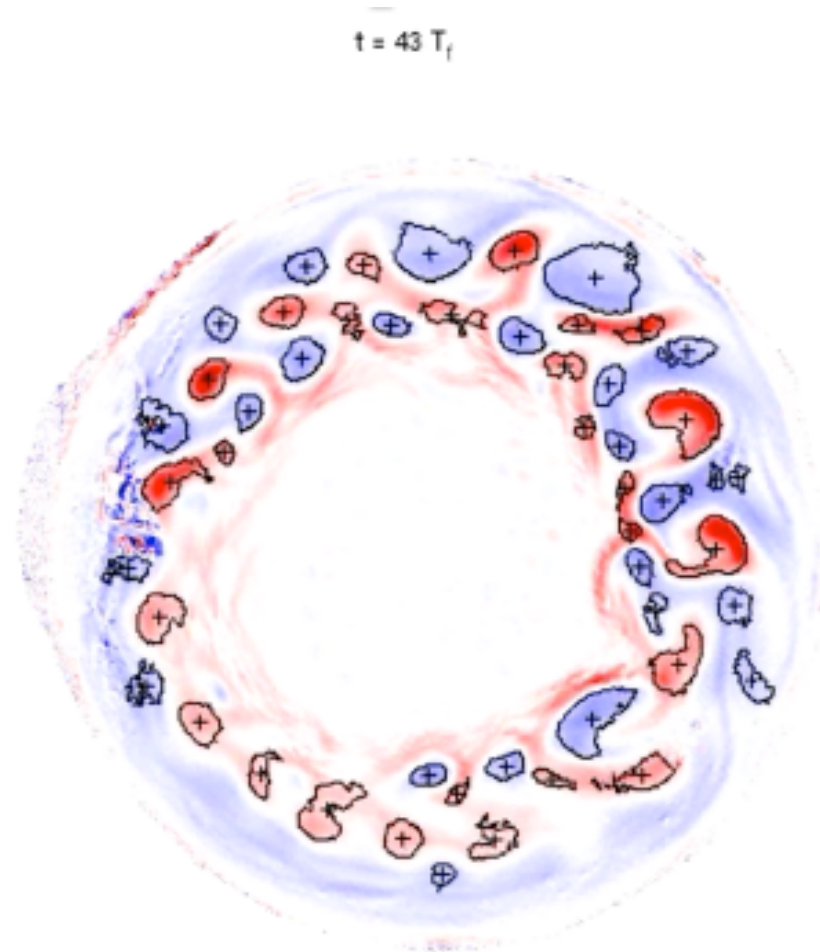
High resolution
PIV measurements:
surface vorticity

blue: anticyclonic
red: cyclonic

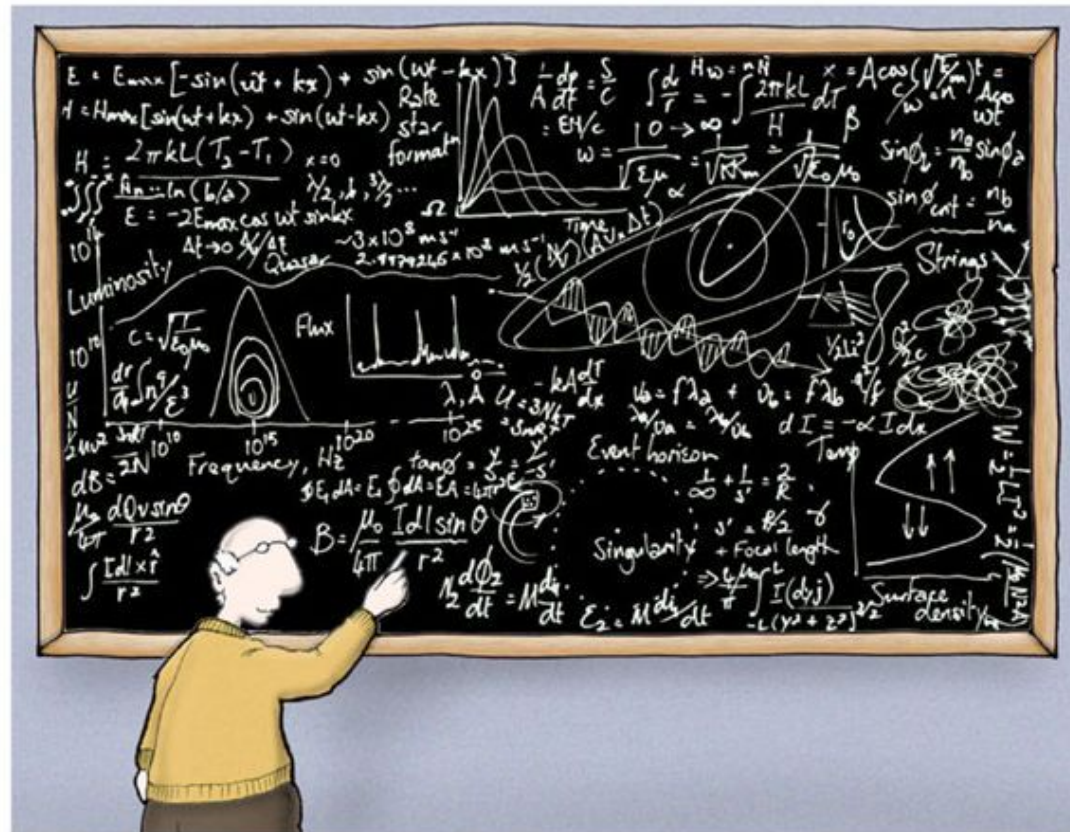
LINEAR SHELF

Topographic parameter

$$T_0 = \frac{s}{\alpha} = -1.3$$

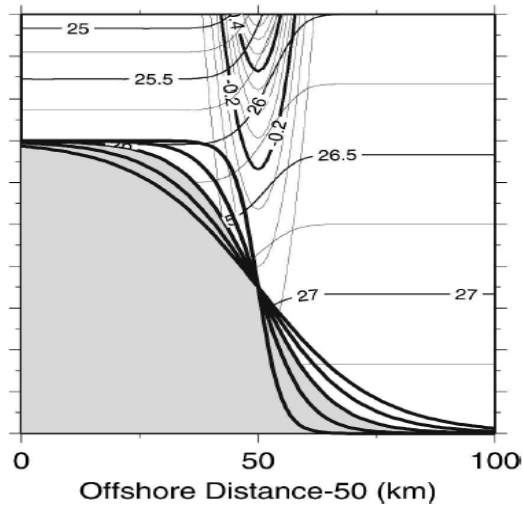


What do we learn from linear stability analysis ?

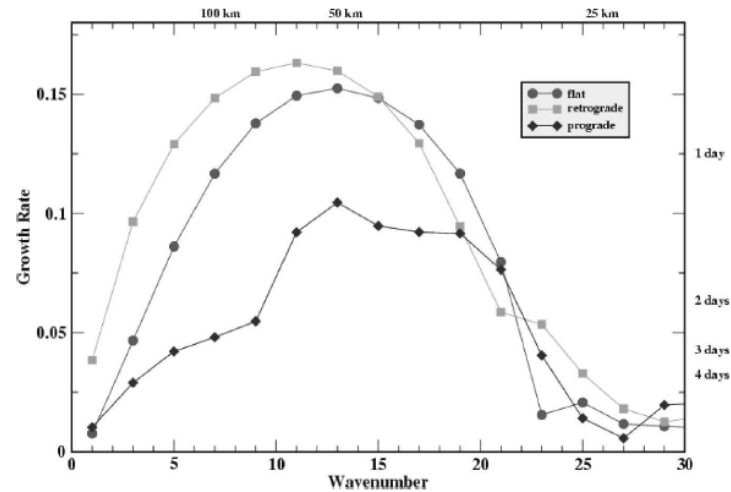


Some previous studies ... not exhaustive !

Coastal current over shelf slope



Shelf slope **destabilize** the current



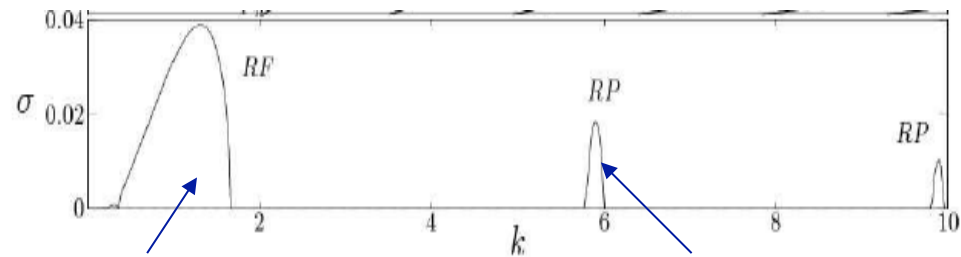
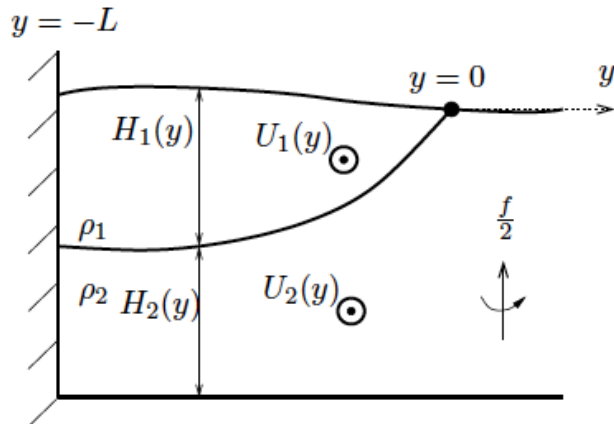
Lozier & Reed, *JPO*(2005)

S

λ_{\max}

σ_{\max}

Coastal front over flat bottom

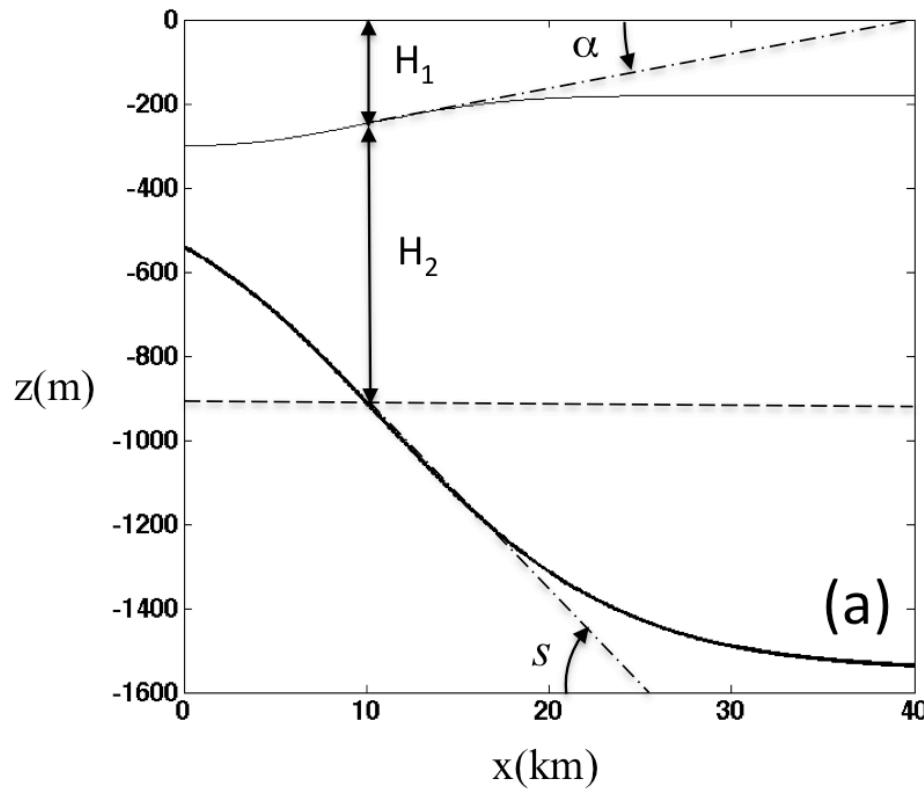


Frontal-Rossby modes

Poincaré-Rossby modes

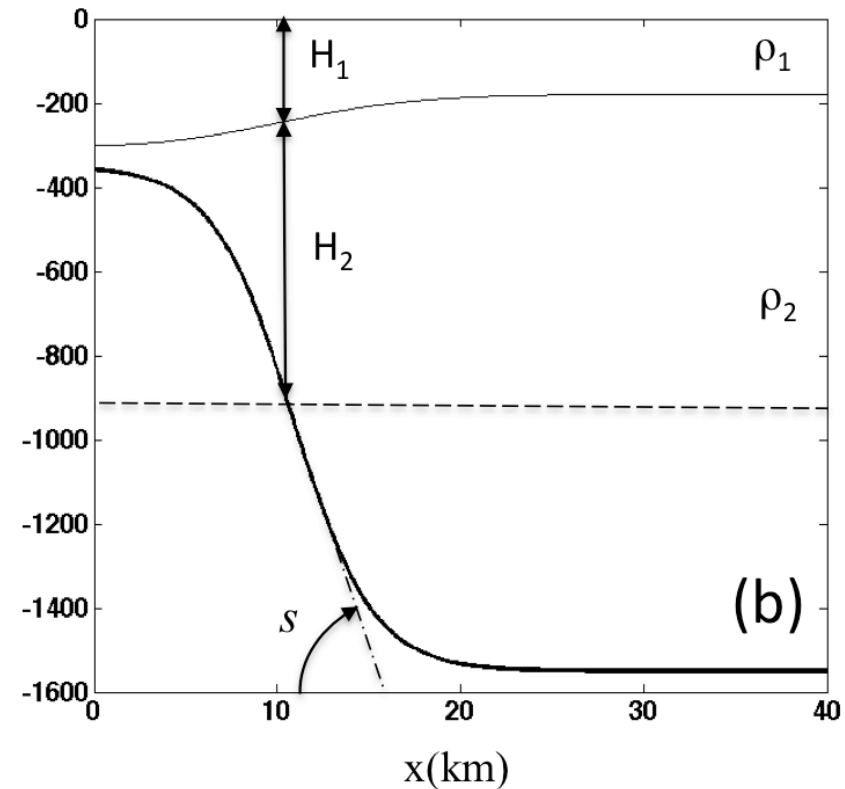
Boss et al. (1996); Gula, Zeitlin & Bouchut, *JFM* (2010)

Shallow-water model: idealized two layers configuration



Vertical stratification parameter

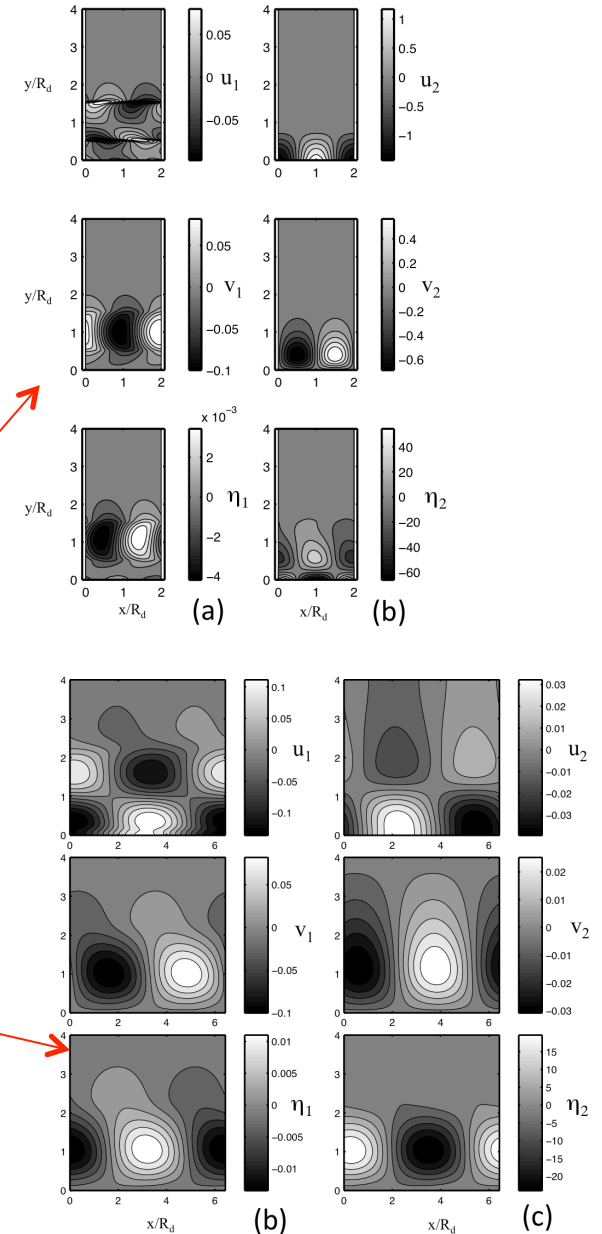
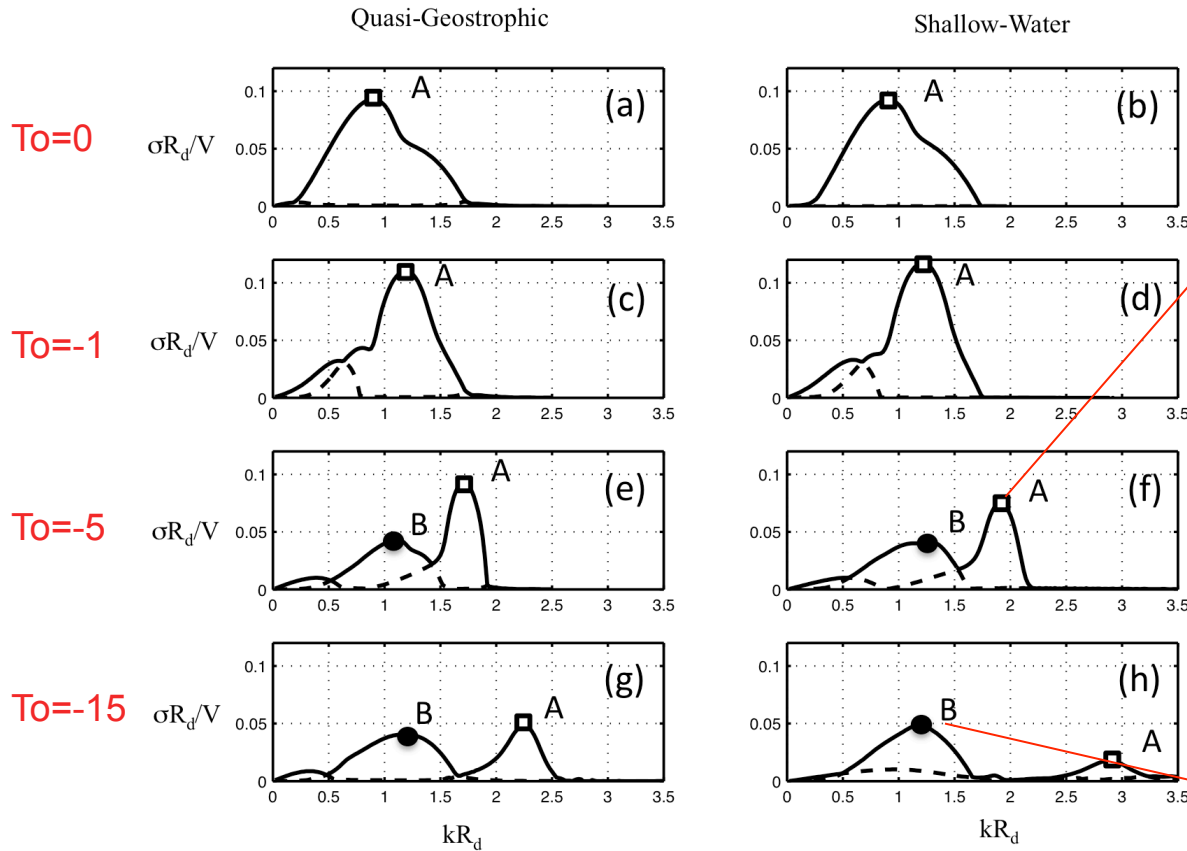
$$\gamma = H_1 / H_2 = 0.4$$



Topographic parameter

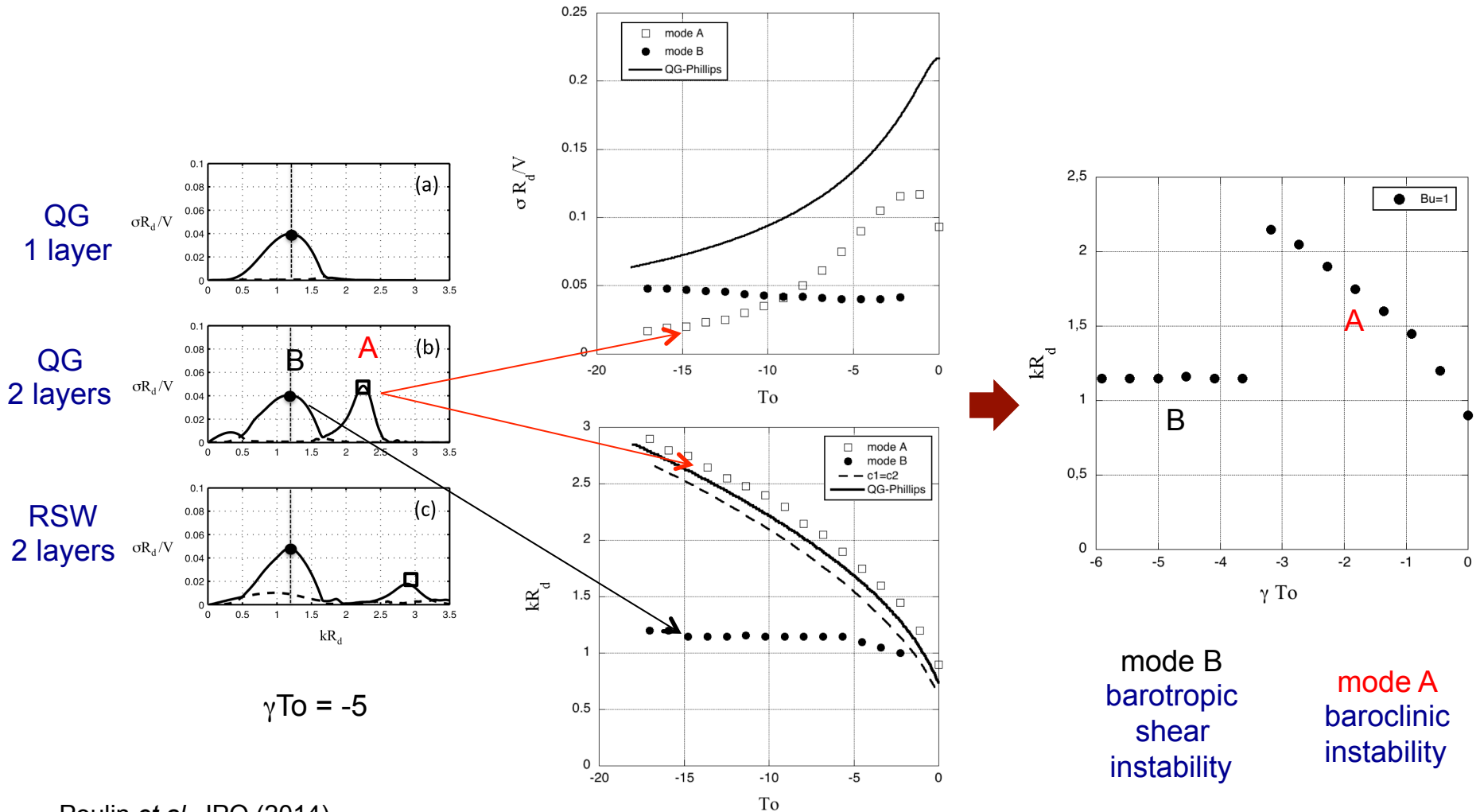
$$T_0 = s / \alpha < 0$$

Shallow-water model: idealized two layers configuration



Two unstable modes: A,B

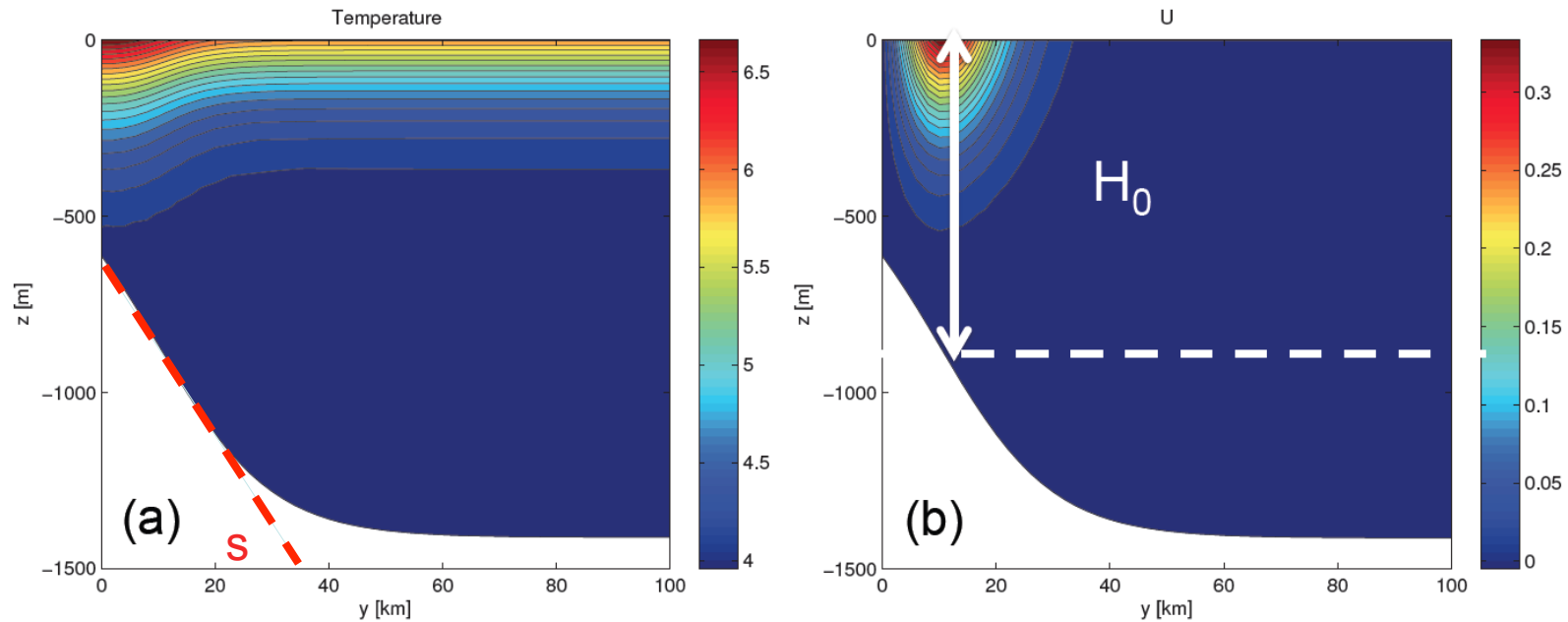
Shallow-water model: idealized two layers configuration



What do we learn from numerical simulations ?



ROMS coastal current: continuous stratification



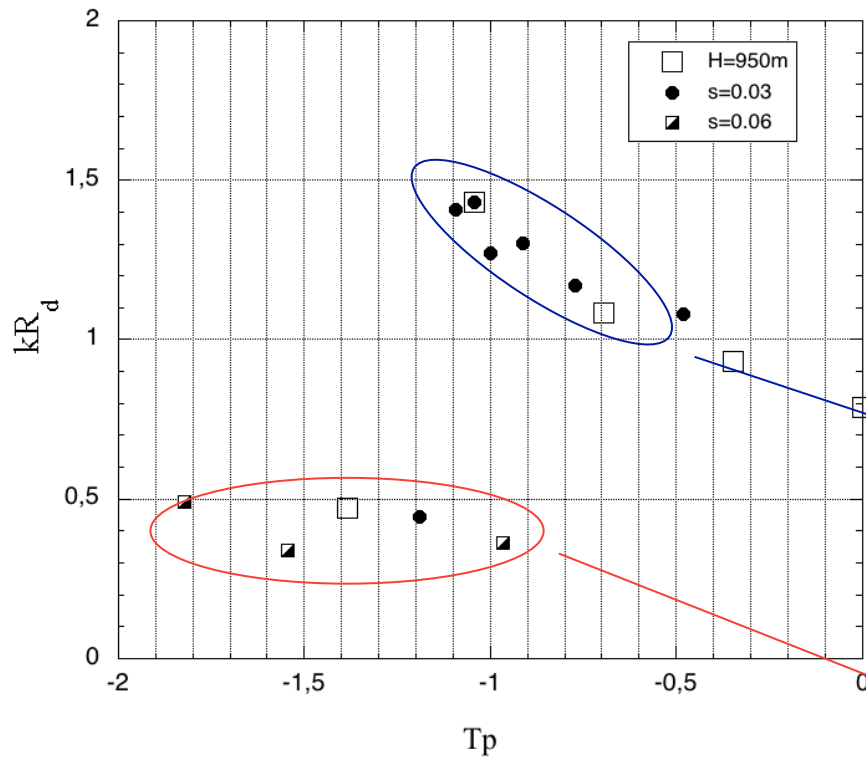
s : shelf slope below the jet

H_0 : water depth at the jet position

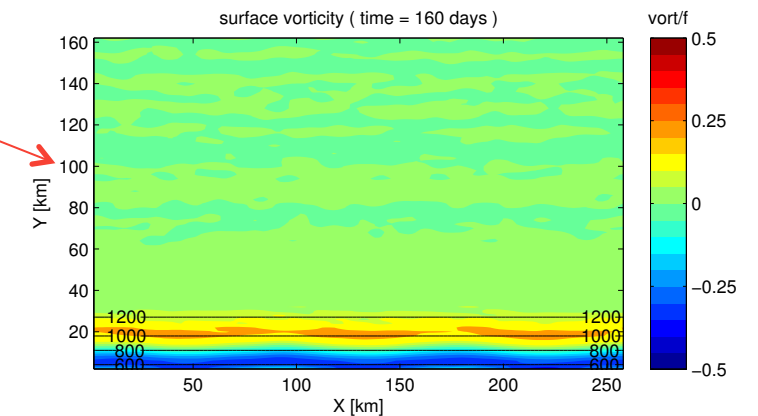
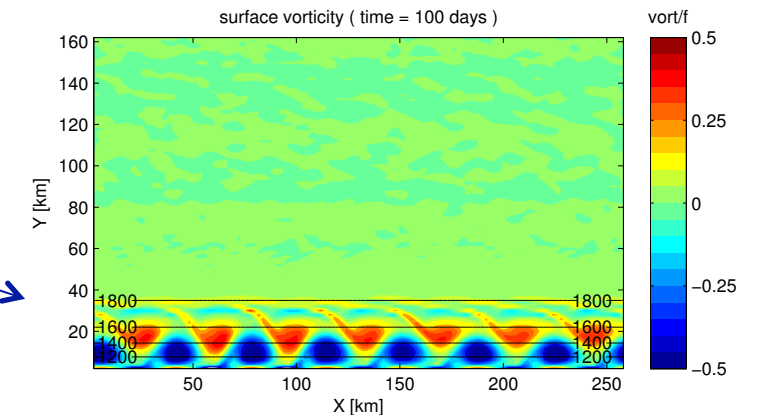
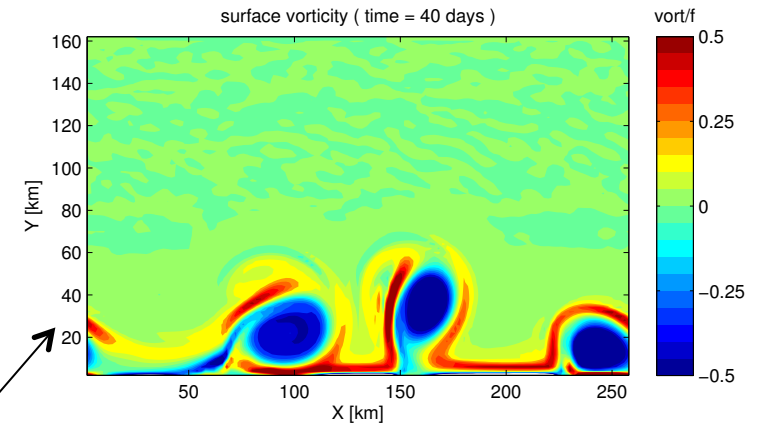
R_d : first baroclinic deformation radius

U_0 : surface jet velocity (constant)

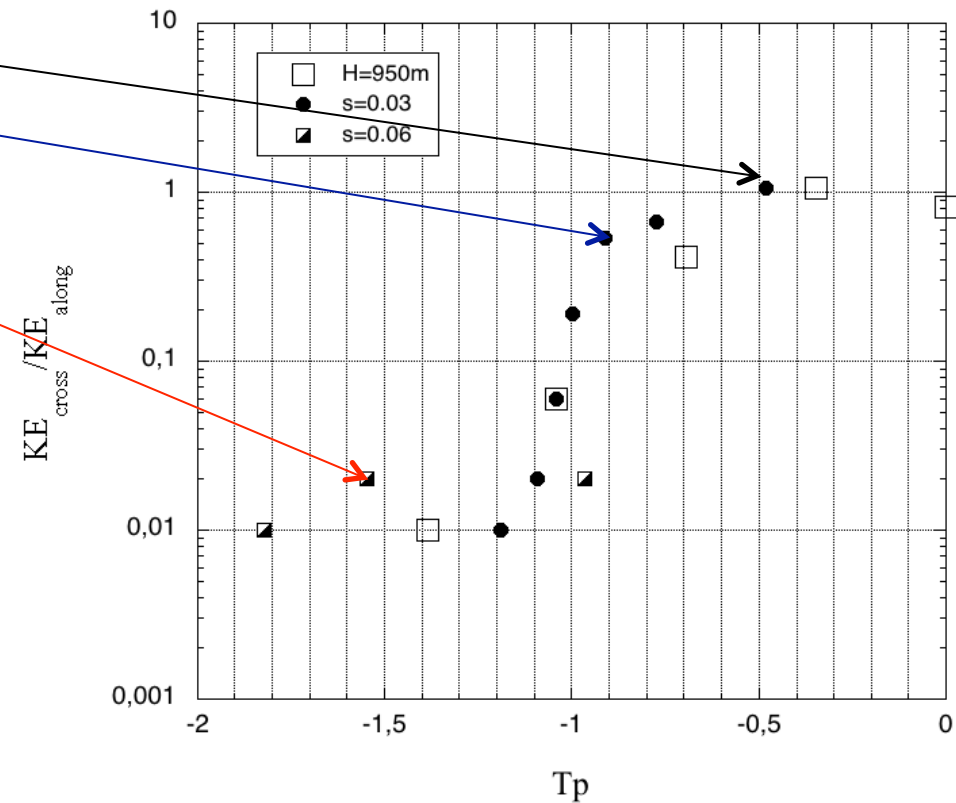
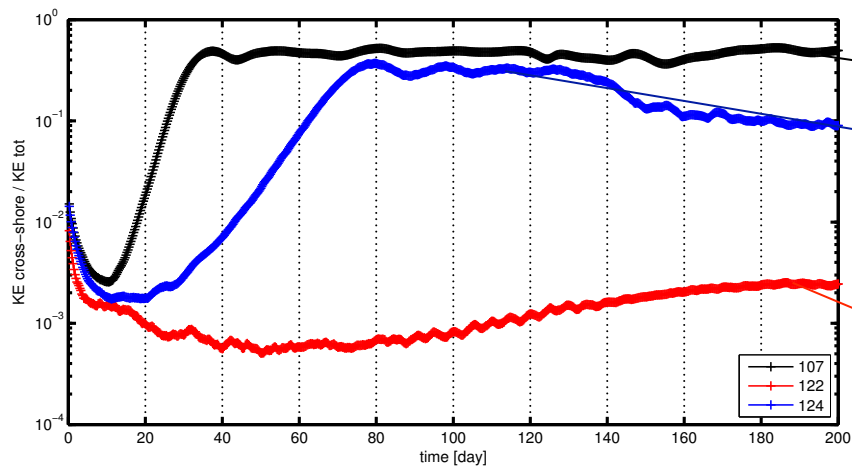
ROMS coastal current: wavelength selection and eddy formation



$$T_p = \frac{s(1 + h_1/h_2)^2 f R_d^2}{H_0 U_0}$$



ROMS coastal current: non linear saturation



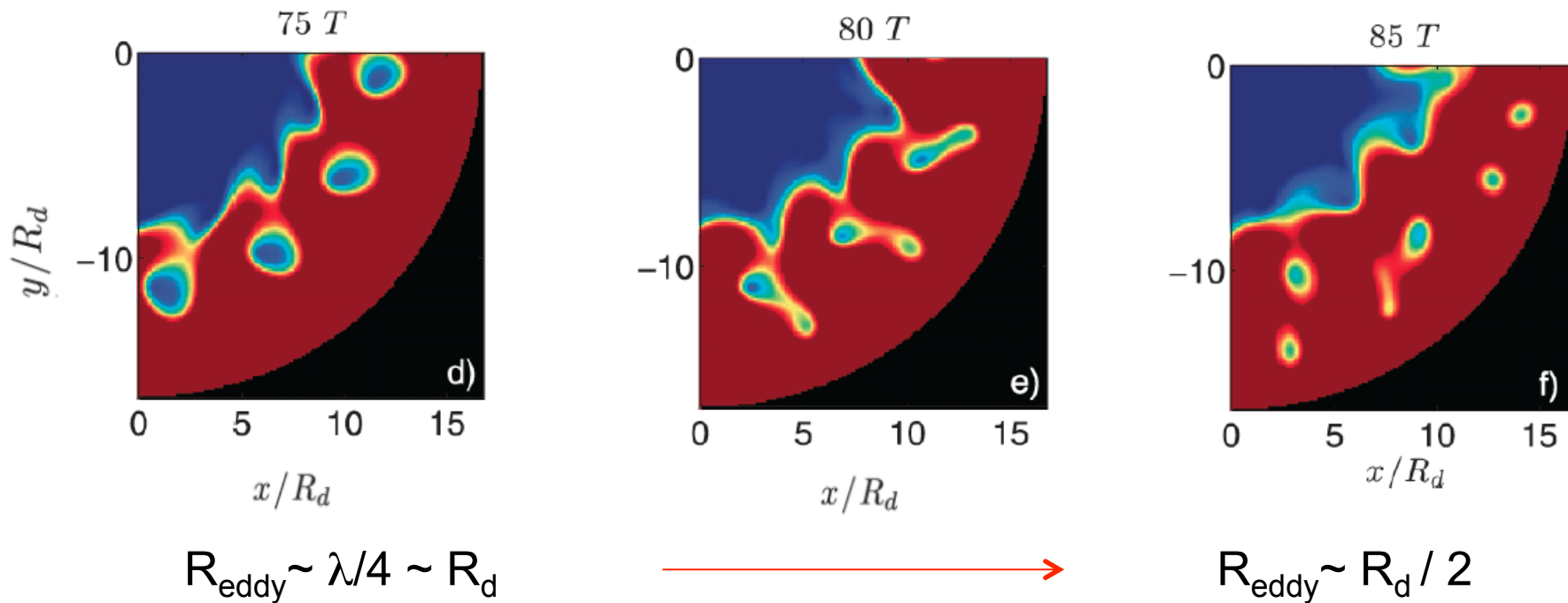
Standard baroclinic instability

Small-scale shelf trapped eddies

Strong non-linear stabilization

Direct cascade towards **sub meso-scale eddies**

NEMO model (high grid resolution $1/8 R_d$ equivalent to $1/72^\circ$ MED) :
coastal front in idealized circular geometry



splitting of cyclones over the shelf

Pennel *et al.* JPO (2012)

We doesn't need ageostrophic instabilities (Rossby-Kelvin or Rossby-Poincaré) to trigger sub meso scale structures !



CONCLUSIONS

The coastal shelf steepness induces on geostrophic coastal currents

Three dynamical regimes

- 1- Standard baroclinic instability, meso-scale anticyclones
strong cross-shelf exchanges
- 2- Rossby-TopoRossby instability, **trapped** anticyclones & cyclones
reduced cross-shelf exchanges, local cyclonic upwelling
- 3- Strong non-linear stabilization, **weak** barotropic shear disturbances
no cross-shelf exchanges, strong along shore transport

Direct **cascade towards sub meso-scale eddies** may occurs !

Our goal: relevant **generalized topographic (slope) parameters** for two-layer and 3D models and **impacts on coastal biology**.