

Eddy kinetic-energy redistribution in quasi-geostrophic flows: implication for midlatitude winter storms.

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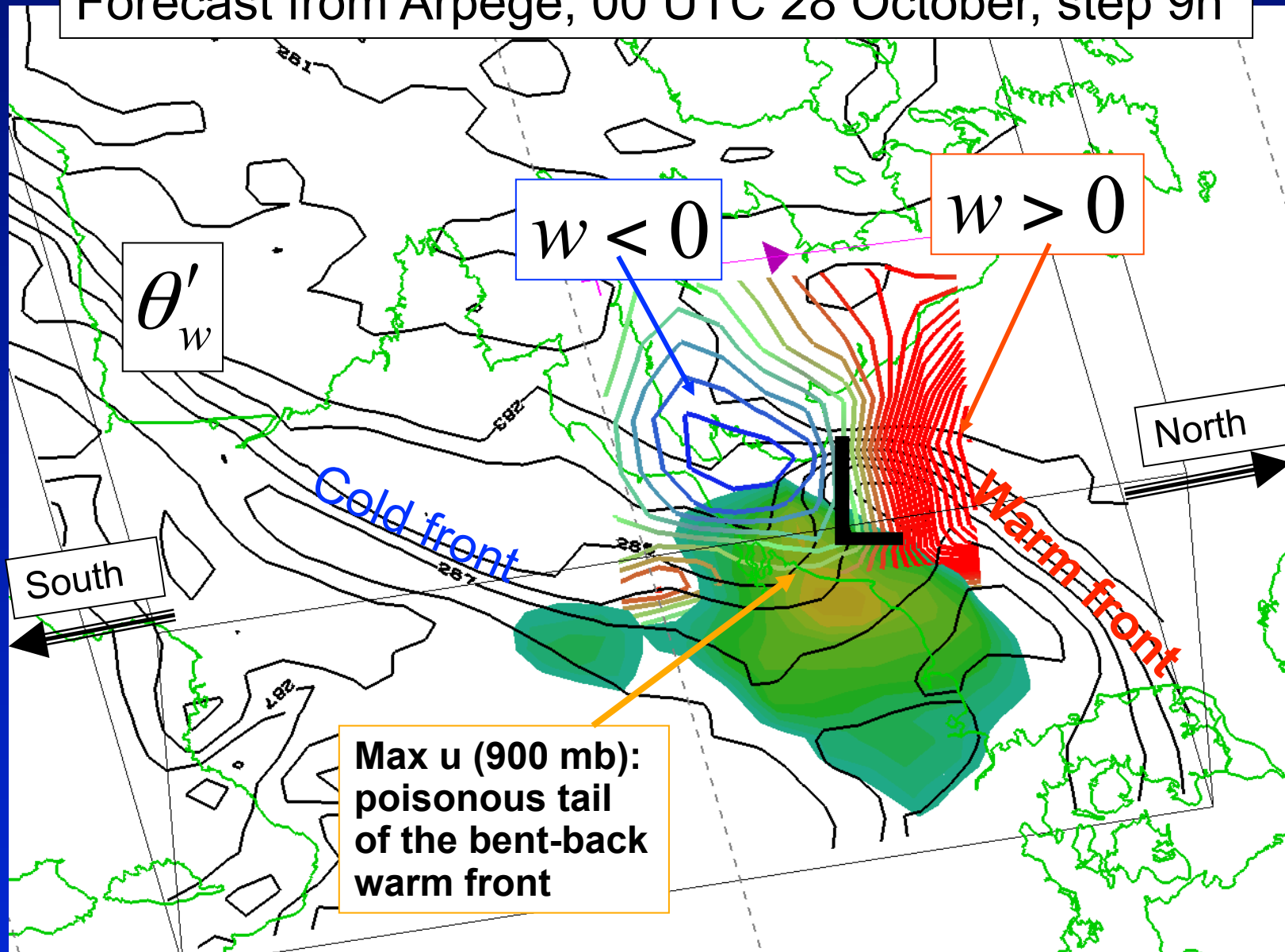
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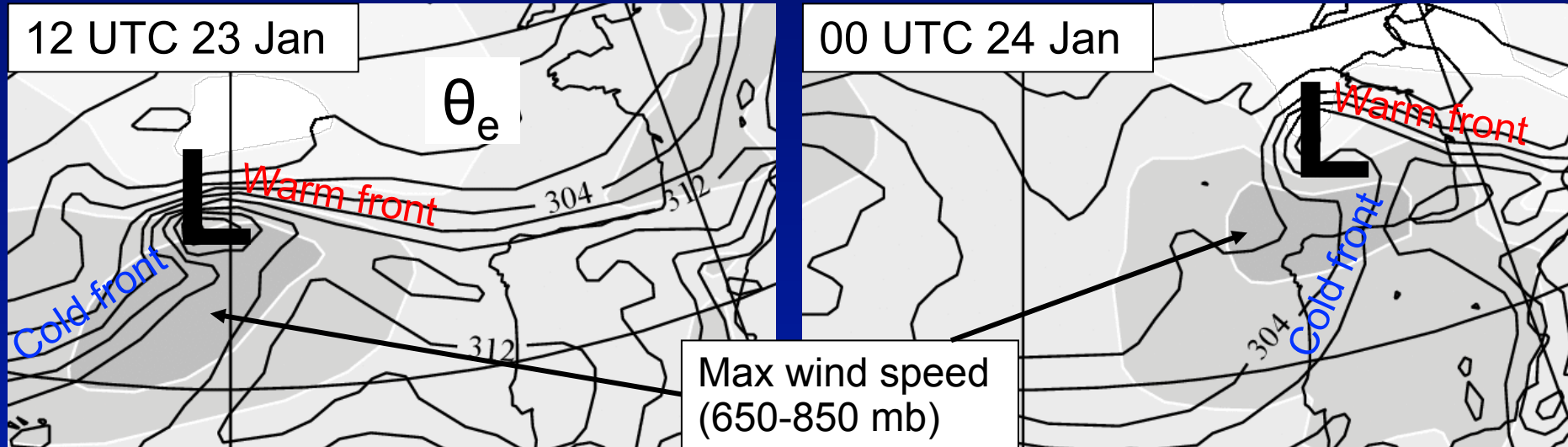


A 3D easterly view of the storm Christian (Oct 2013)

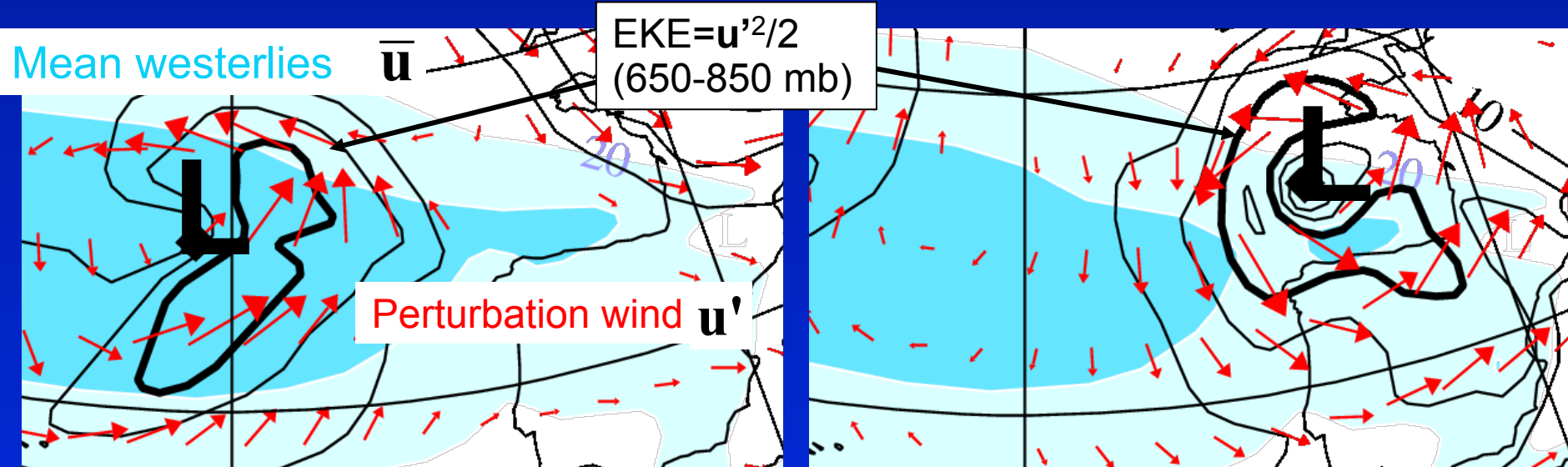
Forecast from Arpege, 00 UTC 28 October, step 9h



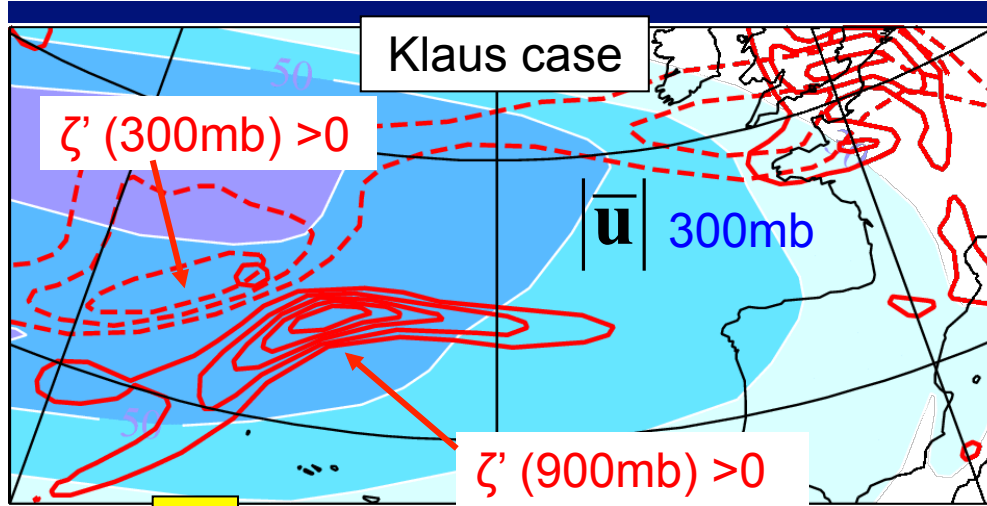
Example of the storm Klaus (24/01/09)



Approach: separation into a background flow and a perturbation by time filtering



Idealized study: use of the two-layer quasi-geostrophic model

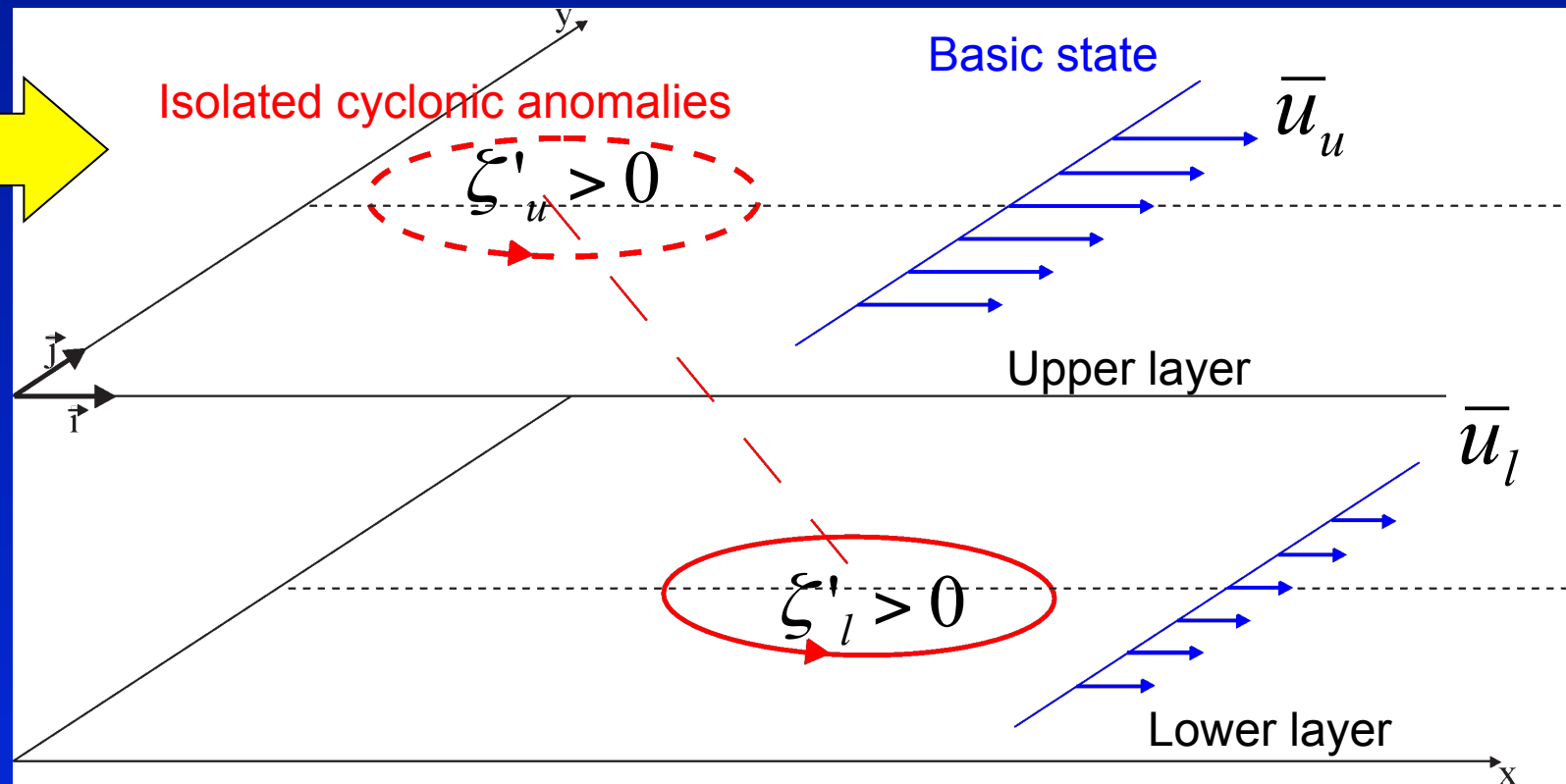


$$q_u = \Delta\psi_u + f - \lambda^{-2}(\psi_u - \psi_l)$$

$$q_l = \Delta\psi_l + f + \lambda^{-2}(\psi_u - \psi_l)$$

$$\partial_t q_k + \mathbf{u}_k \cdot \nabla q_k = 0$$

λ Rossby radius of deformation



The eddy kinetic energy (EKE) equation

$$\frac{\partial}{\partial t} \frac{\mathbf{u}'^2}{2} = -\mathbf{u}' \cdot (\mathbf{u} \cdot \nabla \mathbf{u})' - \mathbf{u}'_a \cdot \nabla \Phi'$$

Φ geopotential

\mathbf{u} geostrophic wind

\mathbf{u}_a ageostrophic wind

Pressure work

$$-\mathbf{u}'_a \cdot \nabla \Phi' = \omega' \frac{\partial \Phi'}{\partial p} - \nabla \cdot (\mathbf{u}'_a \Phi') - \frac{\partial}{\partial p} (\omega' \Phi')$$

Baroclinic conversion Horizontal ageostrophic geopotential fluxes Vertical ageostrophic geopotential fluxes

$$\mathbf{u}_a = \nabla \chi - \frac{1}{f_0} \mathbf{k} \times \nabla \phi_a$$

$$\nabla^2 \phi_a = 2J(u, v) + \nabla \cdot \left(\frac{\beta y}{f_0} \nabla \Phi \right)$$

Diagnostic equation for ageostrophic scalar ϕ_a

$$s^2 \nabla^2 \omega + f_0^2 \frac{\partial^2 \omega}{\partial p^2} = -2 \nabla \cdot \mathbf{Q} + f_0 \beta \frac{\partial v}{\partial p}$$

Omega equation

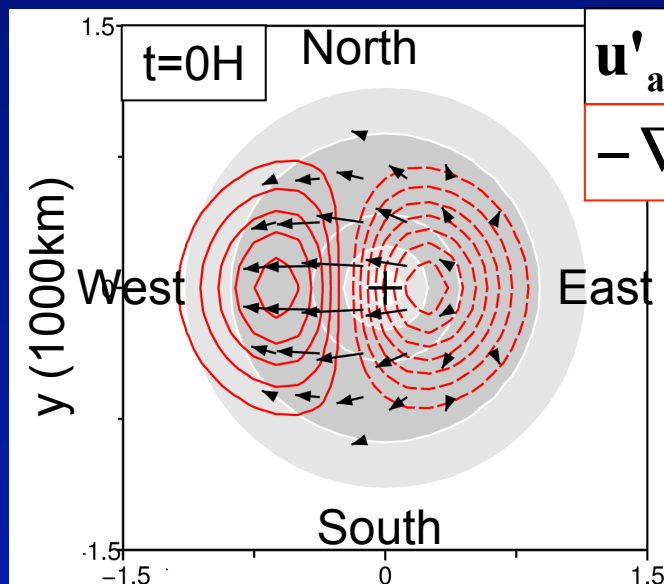
$$\nabla^2 \chi = \frac{\partial \omega}{\partial p}$$

Q-vector

EKE redistribution at the lower layer in presence of the vertical shear

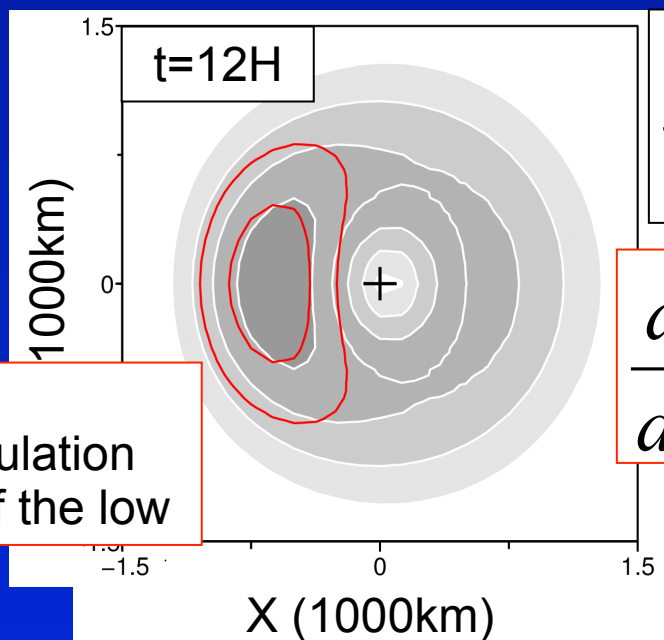
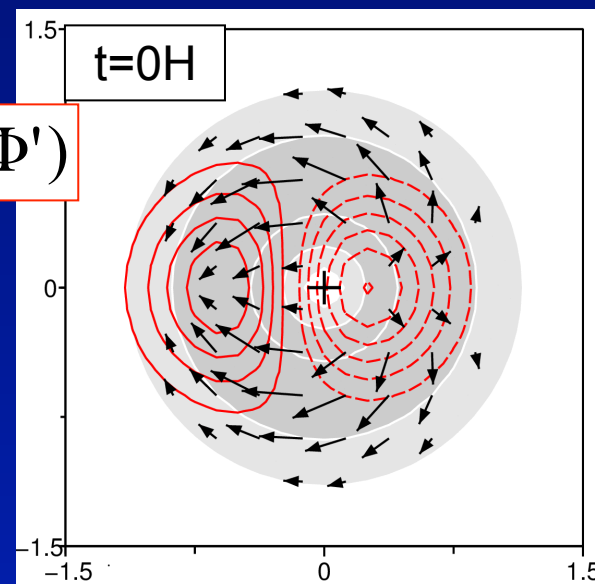
f-plane, linear

β -plane, linear



$$\mathbf{u}'_a \Phi'$$

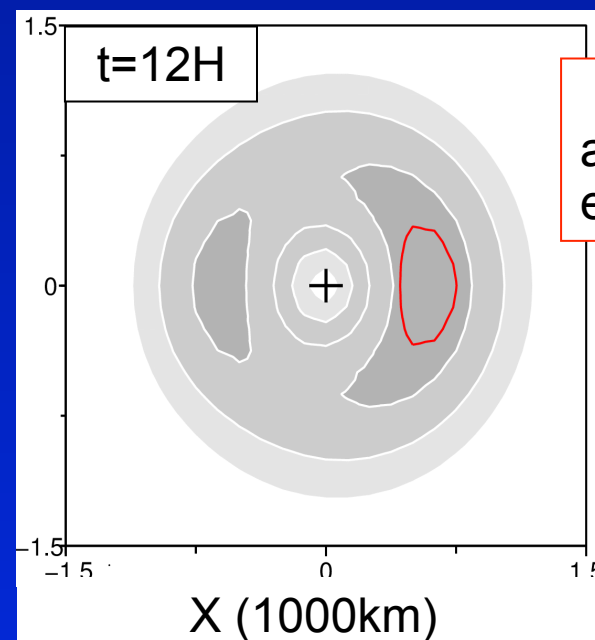
$$-\nabla \cdot (\mathbf{u}'_a \Phi')$$



$$\frac{\mathbf{u}'^2}{2}$$

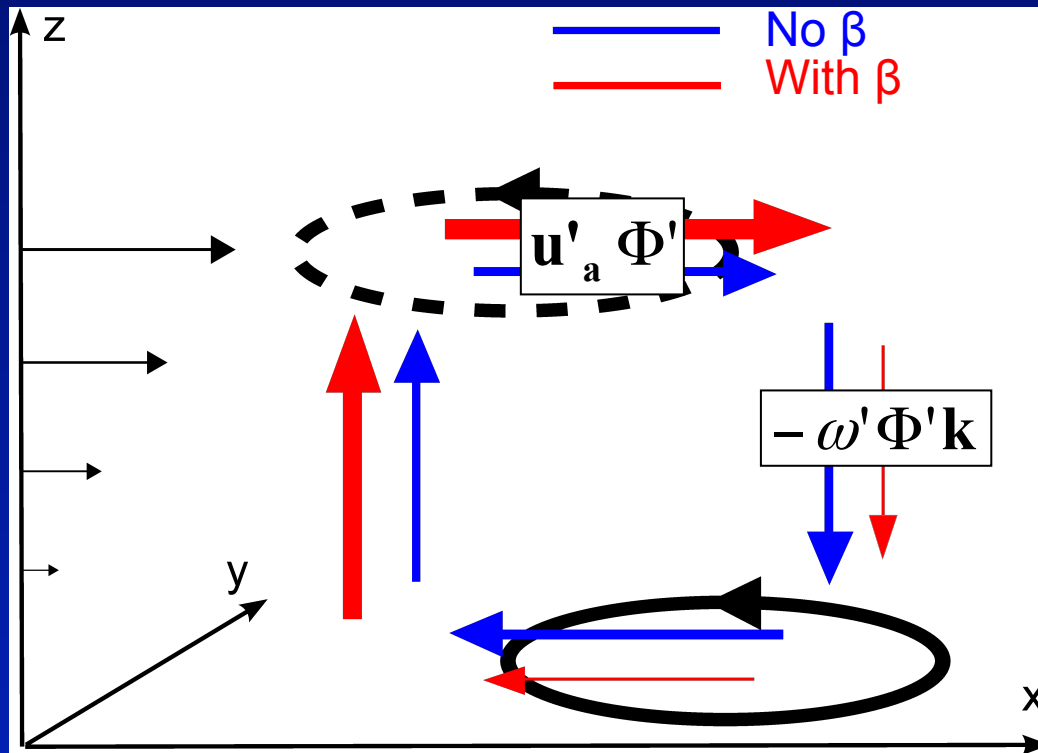
$$\frac{d}{dt} \frac{\mathbf{u}'^2}{2}$$

EKE accumulation west of the low



EKE accumulation east of the low

Role of β and the vertical shear in EKE redistribution



β increases the upward transfer and decreases the downward transfer, so a sink of energy for the lower layer !!!

Wavelike disturbance assumption (wavenumbers m, K)

$$u'_{al} \approx \frac{m^2}{f_0(K^2 + 2\lambda^{-2})} \left\{ -(\bar{u}_{u0} - \bar{u}_{l0})\lambda^{-2}(\psi'_u + \psi'_l) + \frac{\beta}{K^2} \left((K^2 + \lambda^{-2})\psi'_l + \lambda^{-2}\psi'_u \right) \right\}$$

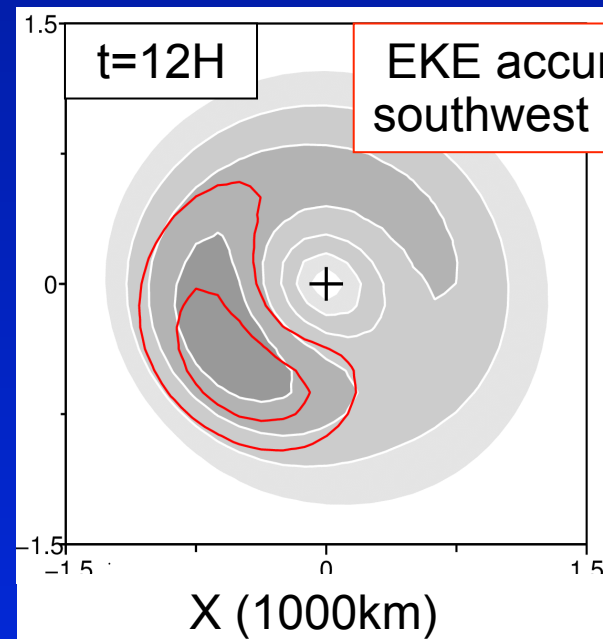
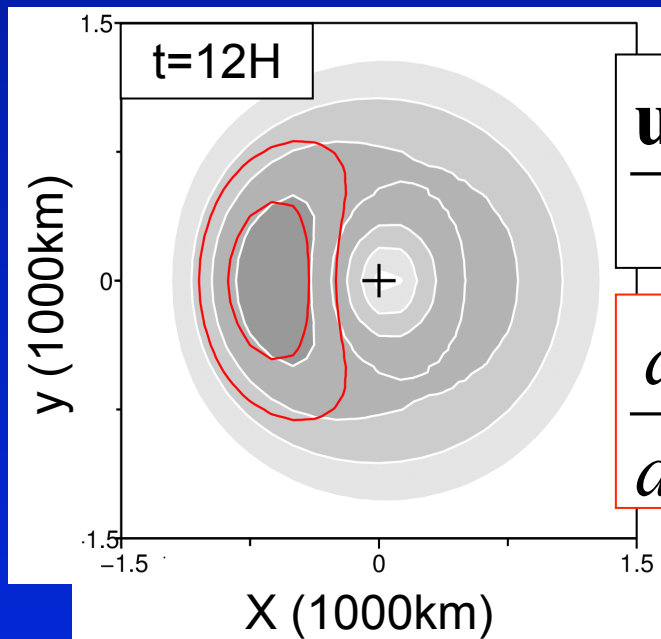
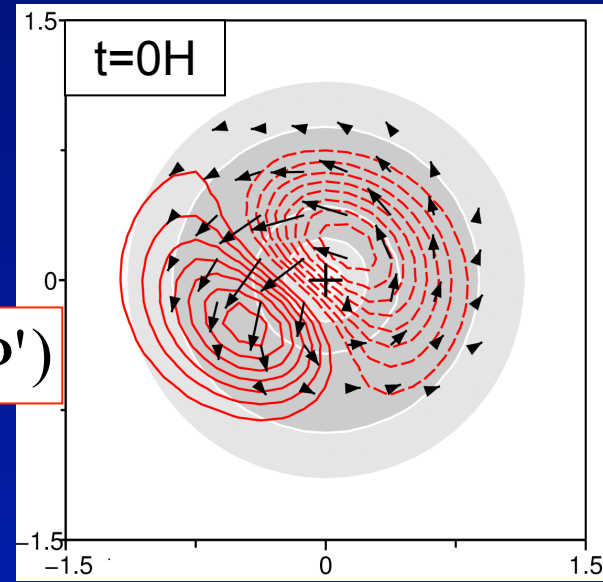
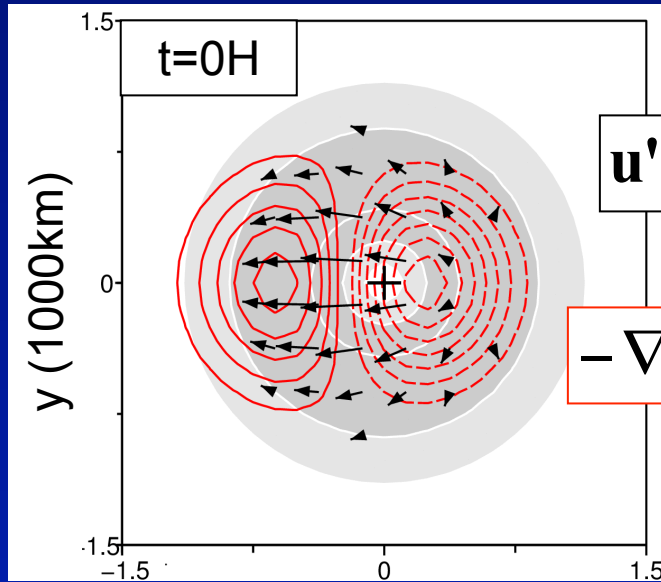
$$\omega' \approx \frac{K^2\lambda^{-2}}{s\lambda^{-1}(K^2 + 2\lambda^{-2})} \left\{ -(\bar{u}_{u0} - \bar{u}_{l0})(v'_u + v'_l) + \frac{\beta}{K^2} (v'_u - v'_l) \right\}$$

Role of nonlinearities in EKE redistribution

f-plane

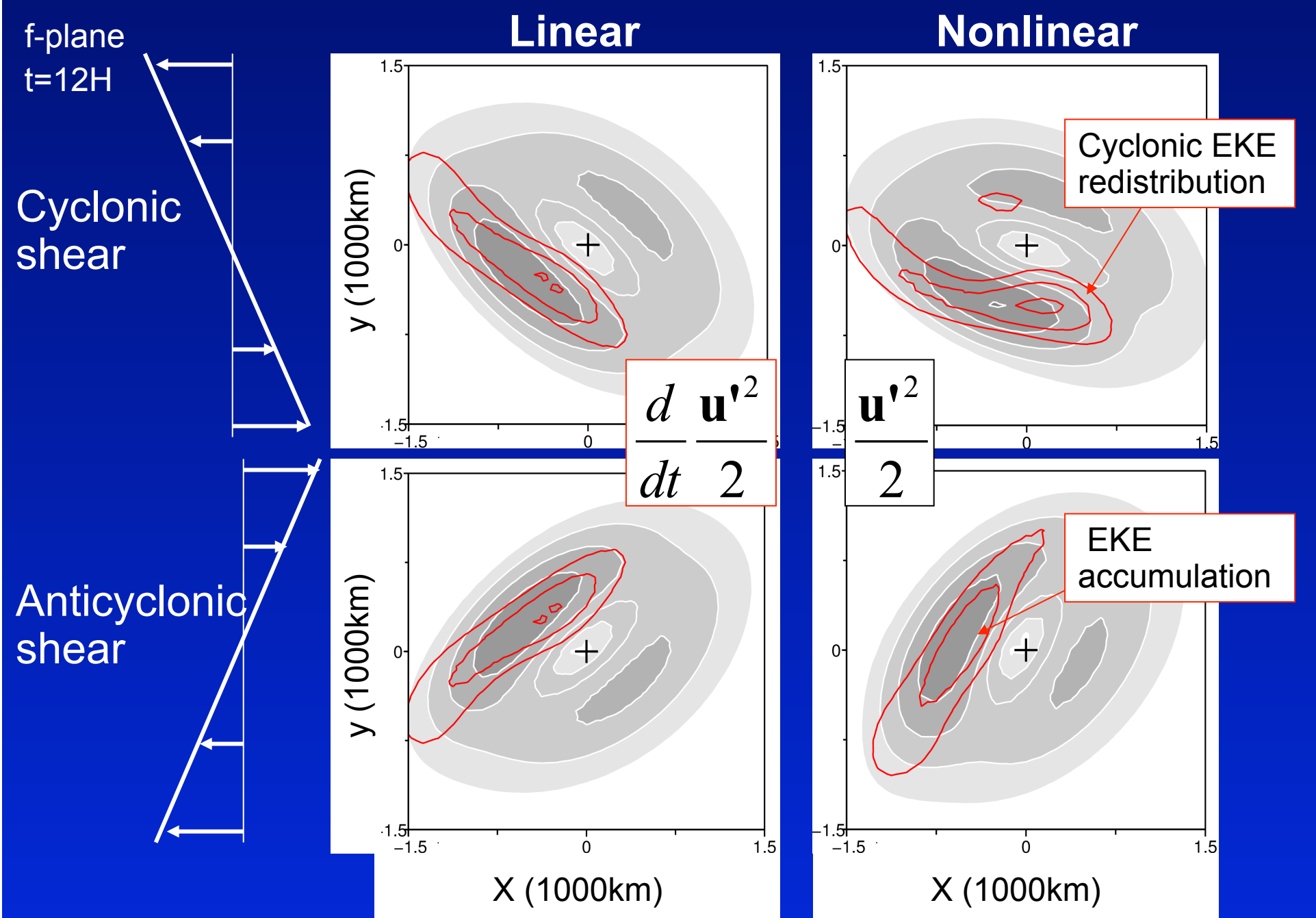
Linear

Nonlinear

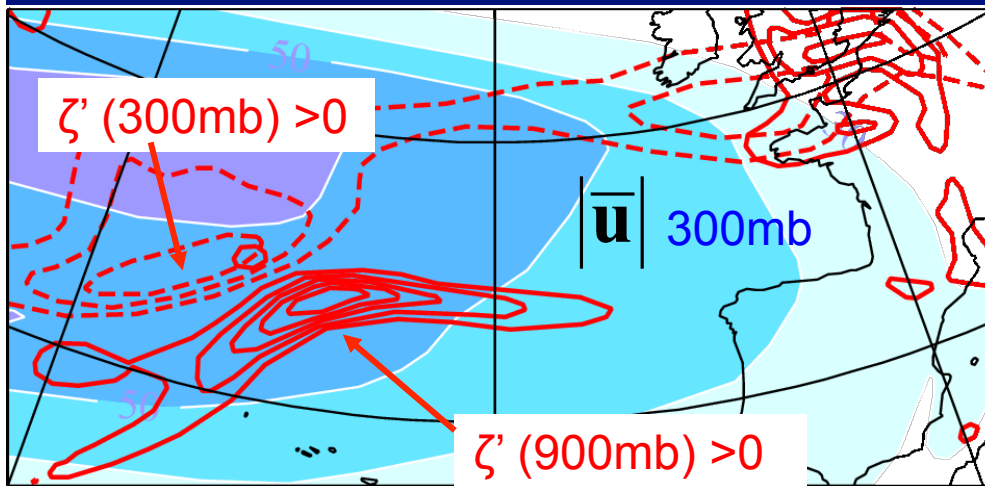


$$\frac{d}{dt} \frac{\mathbf{u}'^2}{2}$$

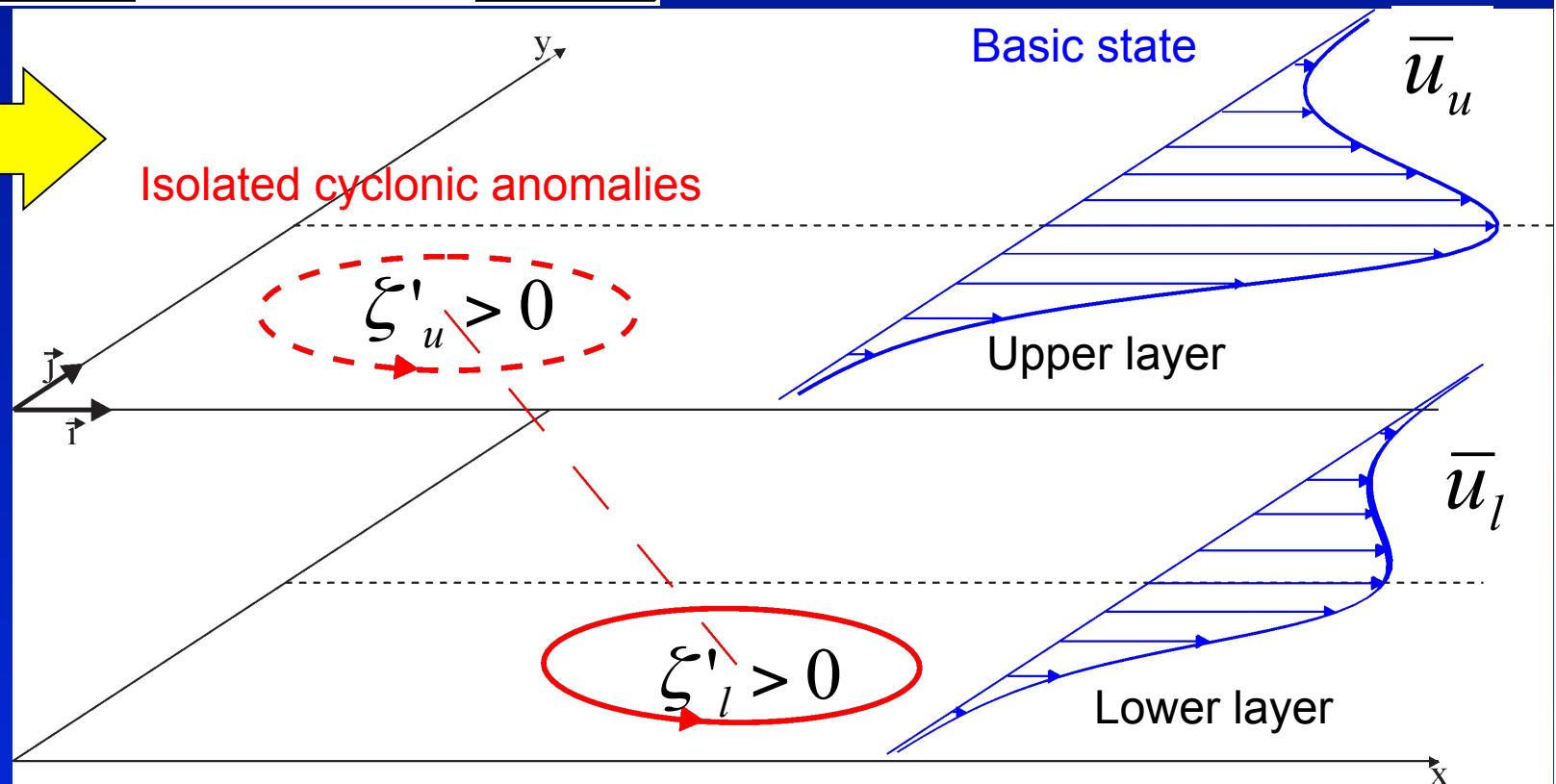
Role of background lateral shear in EKE redistribution



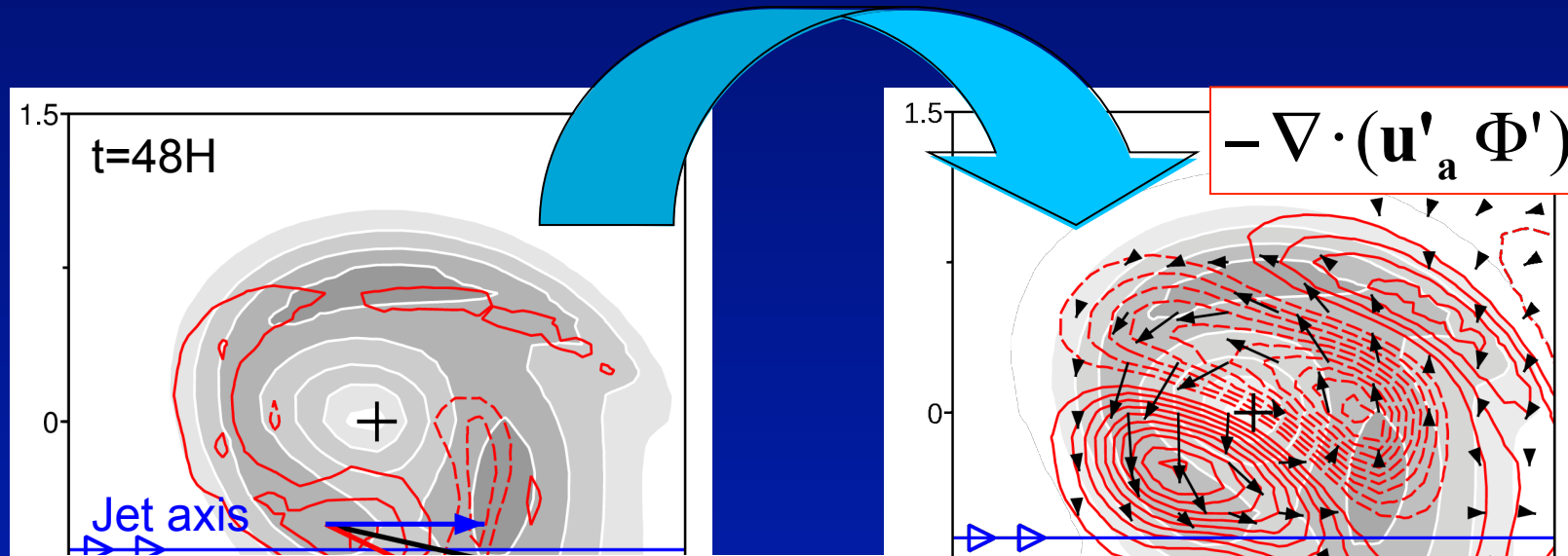
Isolated cyclonic anomalies initiated south of a westerly jet



Isolated cyclonic anomalies initialized south of a meridionally confined zonal jet



Evolution of the idealized jet-crossing cyclone



$$u'_{al} \approx \frac{m^2}{f_0(K^2 + 2\lambda^{-2})} \left\{ \begin{aligned} & -(\bar{u}_u - \bar{u}_l)\lambda^{-2} \\ & + \frac{\beta - \partial_y^2(\bar{u}_u + \bar{u}_l)}{K^2} \left((K^2 + \lambda^{-2})\psi'_l + \lambda^{-2}\psi'_u \right) \\ & + \frac{K^2 + 2\lambda^{-2}}{K^2} \left(\partial_y^2 \bar{u}_u - \partial_y^2 \bar{u}_l \right) \psi'_l \end{aligned} \right\}$$

$$\omega' \approx \frac{K^2 \lambda^{-2}}{s\lambda^{-1}(K^2 + 2\lambda^{-2})} \left\{ \begin{aligned} & -(\bar{u}_u - \bar{u}_l)(v'_u + v'_l) + \frac{\beta - \partial_y^2(\bar{u}_u + \bar{u}_l)}{K^2} (v'_u - v'_l) \end{aligned} \right\}$$

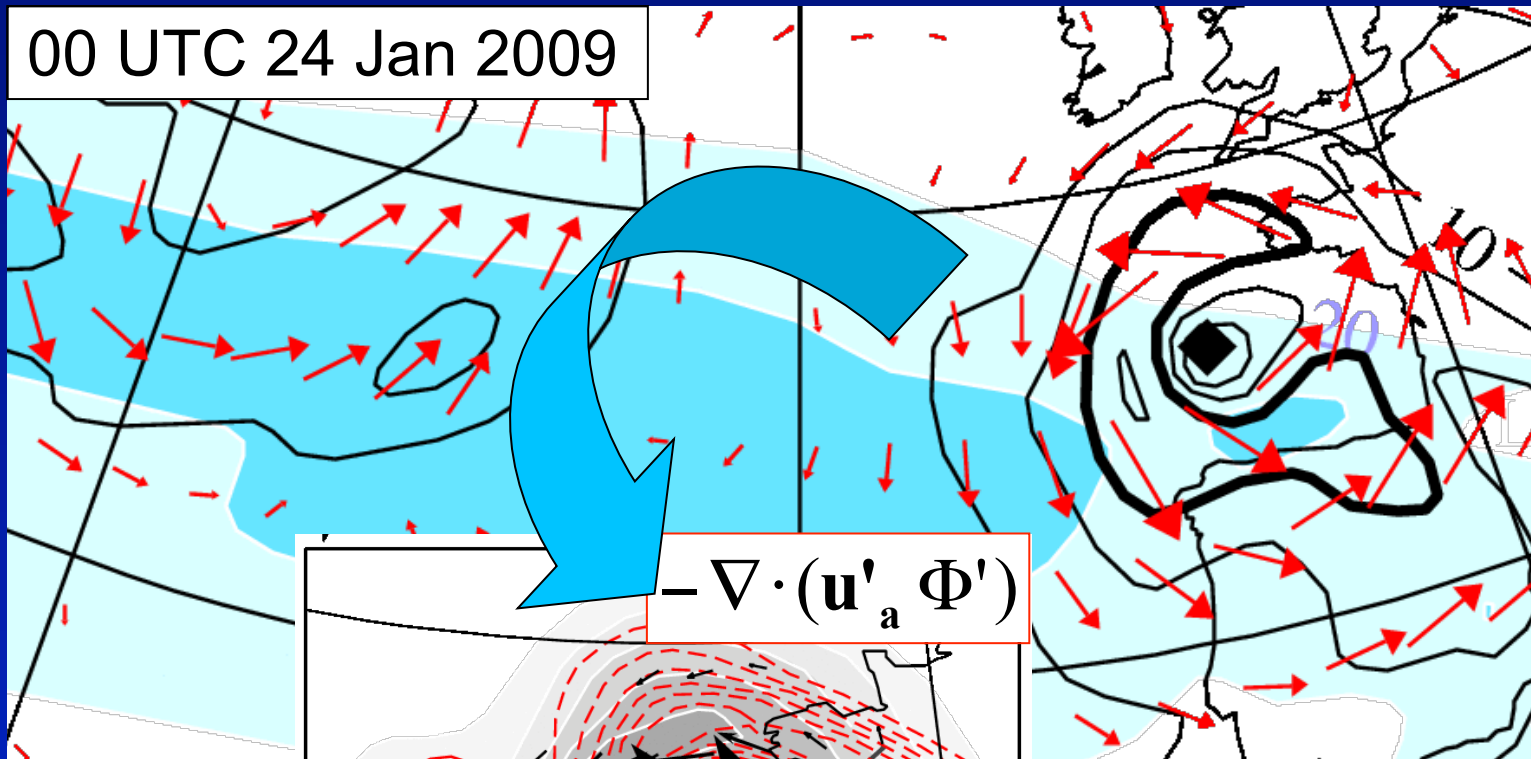
Interpreting EKE distribution during for the storm Klaus

environment $\bar{\mathbf{u}}$

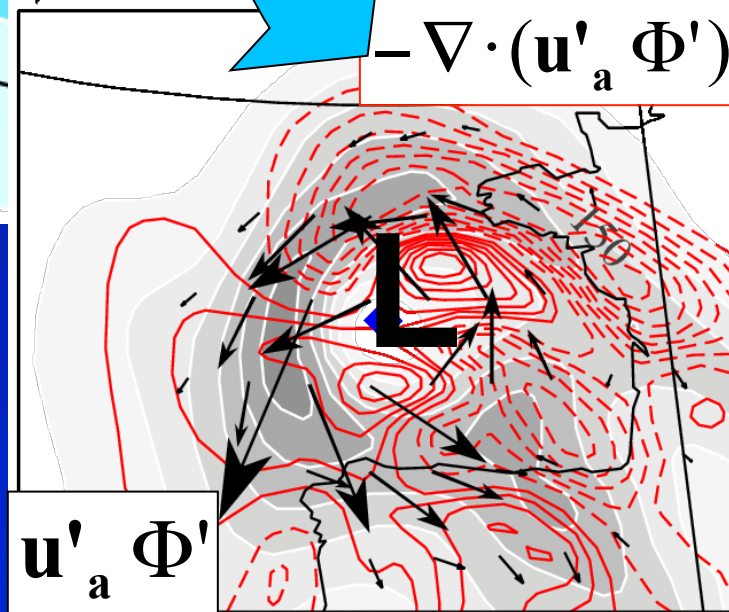
$$\text{EKE} = \mathbf{u}'^2 / 2$$

\mathbf{u}'

00 UTC 24 Jan 2009

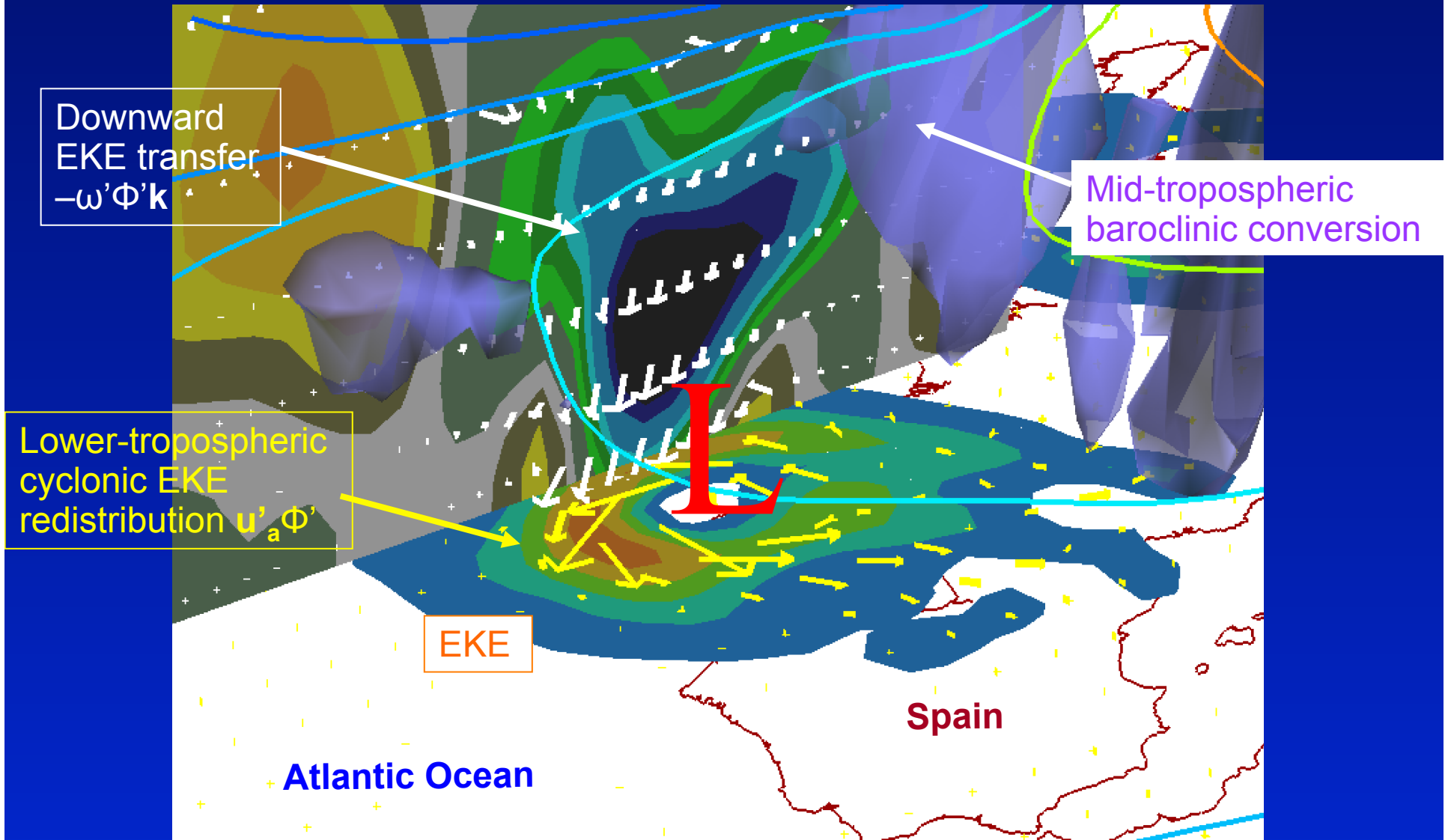


$$-\nabla \cdot (\mathbf{u}'_a \Phi')$$



$\mathbf{u}'_a \Phi'$

A 3D view as a summary



Conclusion: the background flow controls EKE redistribution processes !
Outlook: how is it related to the formation of mesoscale jets (sting jets, Browning, 2004; Gray et al. 2011) ?