

An observational estimate of the direct atmospheric response to the Arctic sea ice loss in the cold season

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Work in Progress



Observational estimates of sea ice loss impacts

- **Connection** between the observed sea ice loss and slow atmospheric changes can easily be established, but **attribution** is difficult since SST, GHG and aerosol, snow cover, ...also slowly decrease or increase
- **AGCMs** can be used to single out the **direct** impact of sea ice loss, but may be affected by model biases
- Hence, **attribution** should still be attempted using observations

Our basic assumptions

- The *direct* atmospheric response to the slow Arctic sea ice loss is the same as that to *interannual pan-Arctic* sea ice fluctuations *with identical spatial patterns*

This disentangles the SIC impact from slow anthropogenic and forced climate variations

The response to pan-Arctic sea ice patterns differs from cumulative regional effects

- The response is sufficiently linear to be estimated by lag regression analysis

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- The response is sufficiently linear to be estimated by lag regression analysis
- Attribution is derived from *lag multiple regression* if there are significant concomitant interannual SST or snow cover changes

Often disregarded in previous studies

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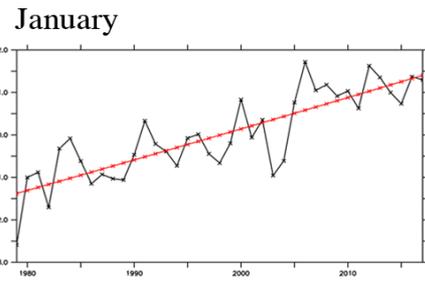
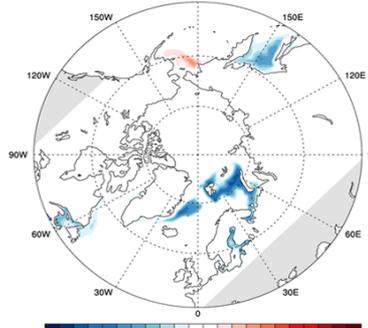
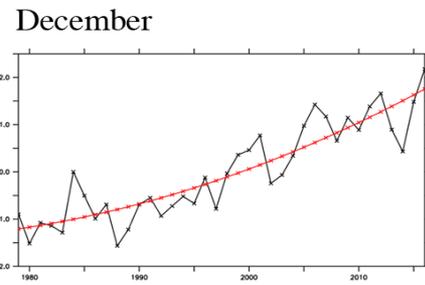
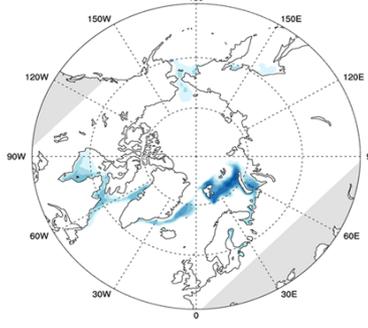
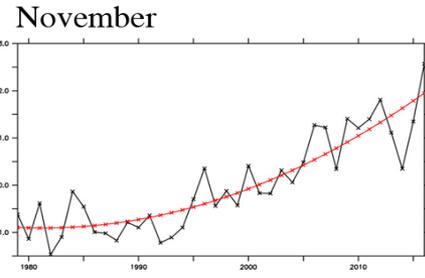
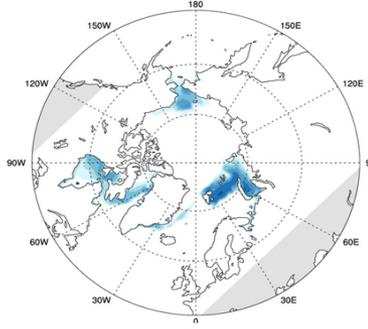
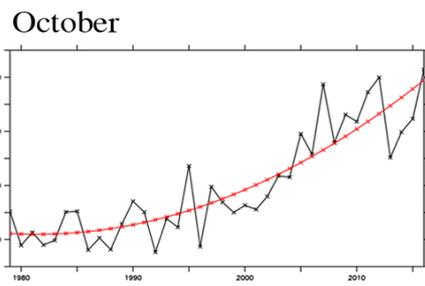
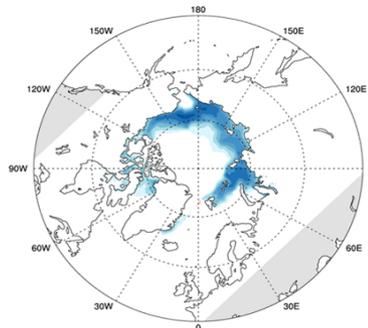
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Often disregarded in previous studies

Observational data

- Monthly sea ice concentration from passive microwave measurements **1979-February 2017**
- ERA-Interim, HadISST, snow cover from Rutgers University

Main patterns of sea ice loss



Sea ice loss is well represented by first EOF and a quadratic fit to PC1

Removing the quadratic fit leads to interannual to fluctuations of *identical* spatial pattern

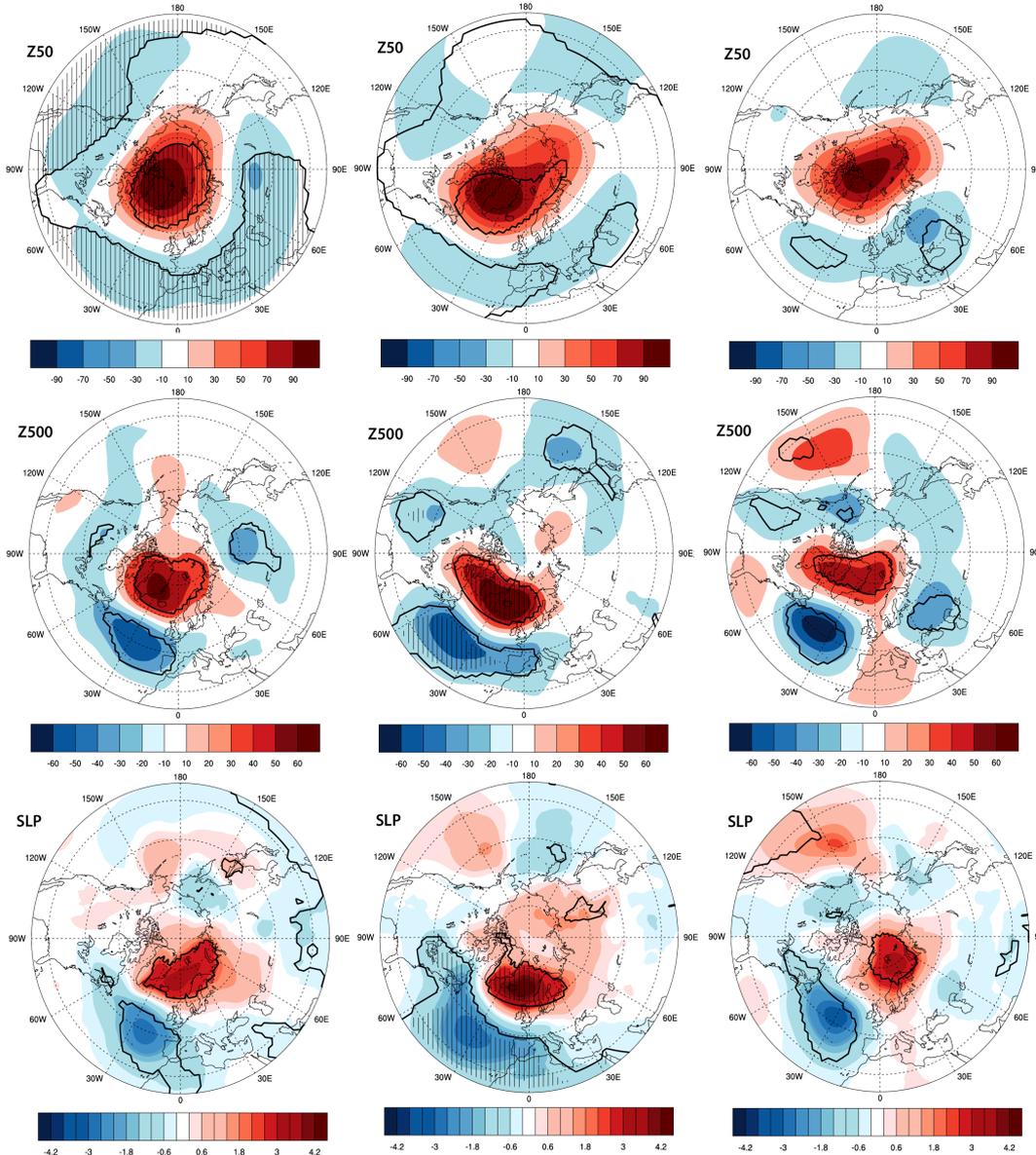
Lag regression is performed on Interannual fluctuations

Lag regression on standardized interannual SIC fluctuations in November

December (lag 1)

January (lag 2)

February (lag 3)



Negative AO in the stratosphere
Field significant in December

Negative NAO in the troposphere
Largest and field significant
in January (2-month lag)

Weaker amplitude in February

but

Amplitudes must be divided by 2
for typical SIC changes

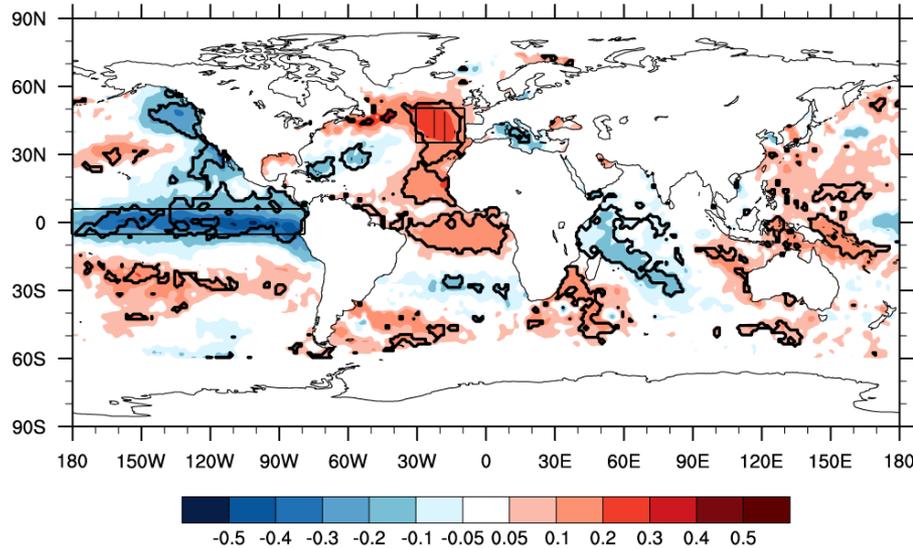
February signal is similar and field
significant when regressed on
December SIC; weaker in March

Black line
10% significance
Hatching
FDR of 10%

Are there concomitant SST and snow cover anomalies in November?

Regression of
quadratically
detrended

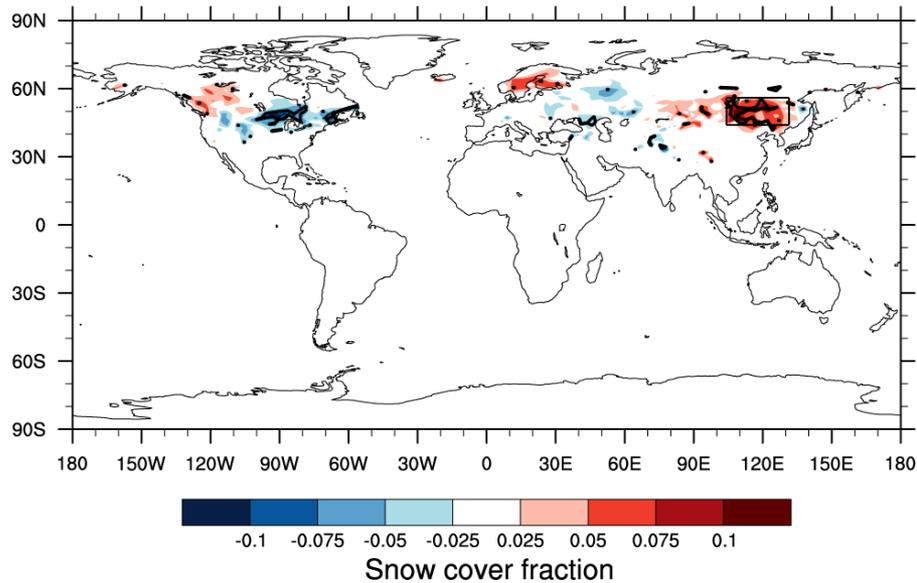
SST



Weak La Nina conditions
with very weak Indian cooling

Warm northeast Atlantic
(no significant impact)

Snow cover

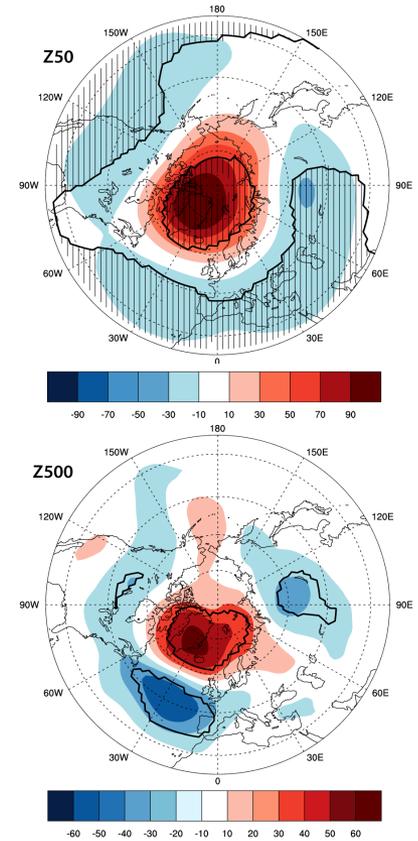


More snow in Siberia
with less in North America

Multiple regression on 3 indices
(SIC, equ. Pac. SST, Siberian snow)

Multiple regression in December on 3 standardized indices in November (lag 1)

Simple regression



Black line

10% significance

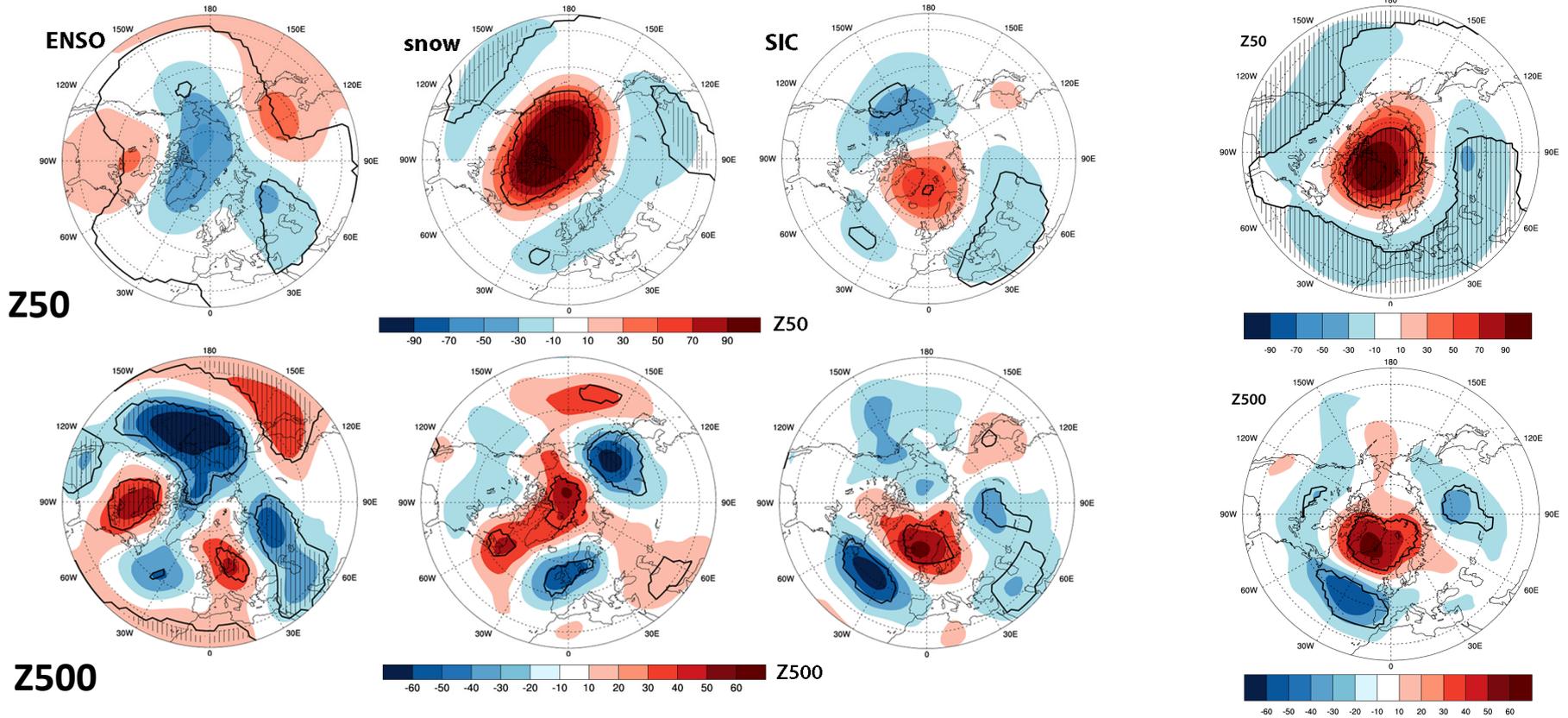
Hatching

FDR of 10%

Multiple regression in December on 3 standardized indices in November (lag 1)

Statistical significance estimated by block bootstrap method 500 permutations

Simple regression



Black line

10% significance

Hatching

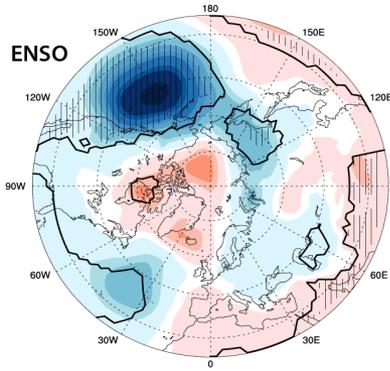
FDR of 10%

Stratospheric AO- signal is due to synchronous Siberian snow cover increase

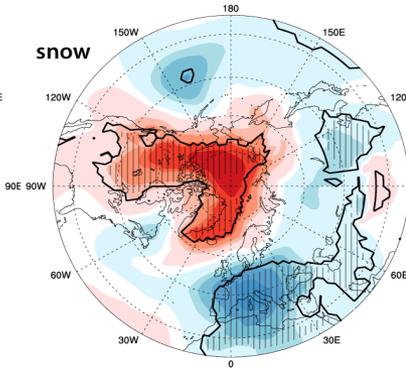
Tropospheric NAO- signal is mostly due to SIC

Multiple regression in January on 3 standardized indices in November (lag 2)

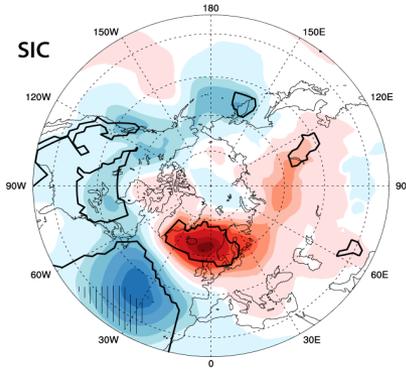
SLP



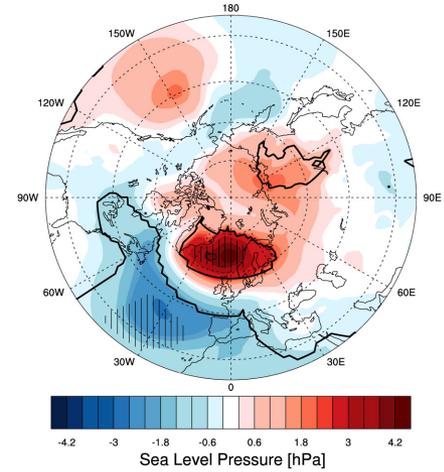
snow



SIC



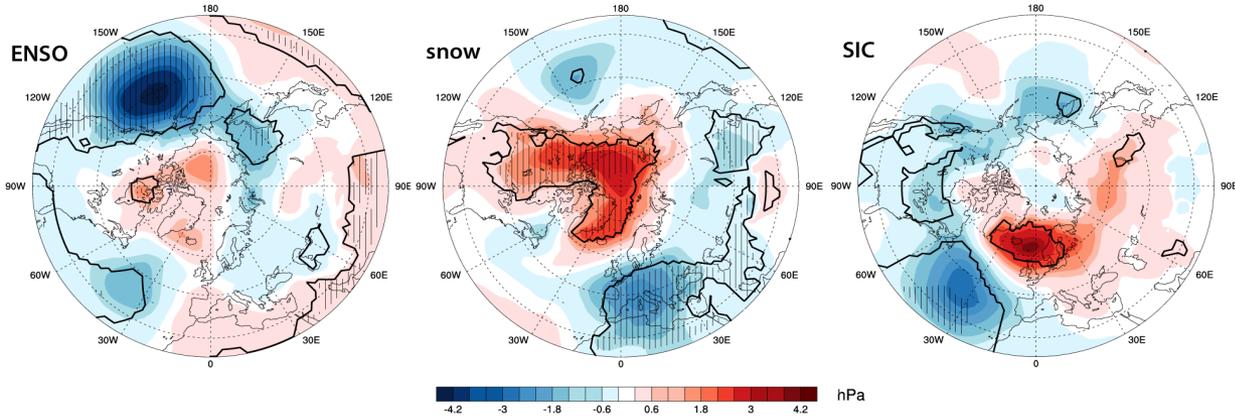
Simple regression



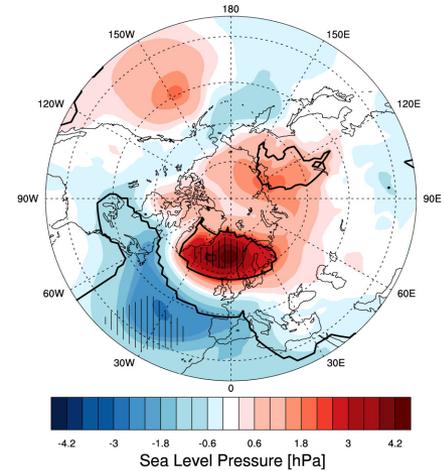
Negative NAO signal in January is largely due to SIC variability

Multiple regression in January on 3 standardized indices in November (lag 2)

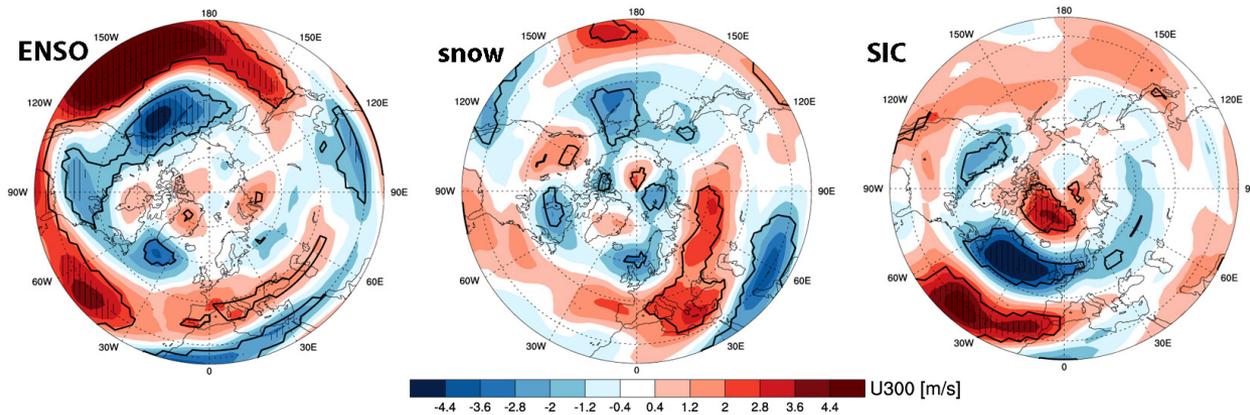
SLP



Simple regression



U300



Negative NAO signal in January is largely due to SIC variability

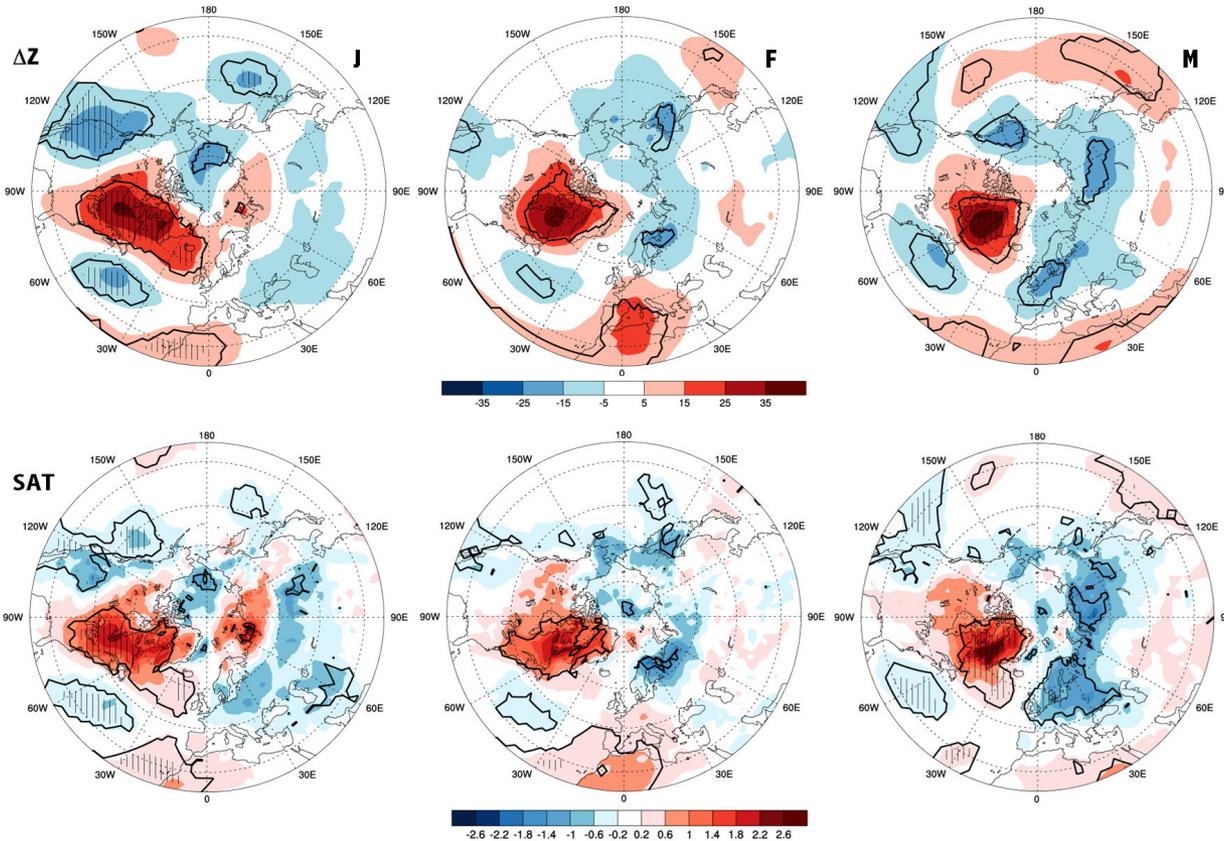
Southward shift on the tropospheric jet in the North Atlantic sector

SAT and 1000-500 hPa thickness response to standardized SIC fluctuations

January (lag 2)

February (lag 3)

March (lag 3)



In addition, warm SAT above SIC retreat at lag 0 and, less, at lag 1

No evidence of warm Arctic cold Eurasia (WACE), except perhaps in March

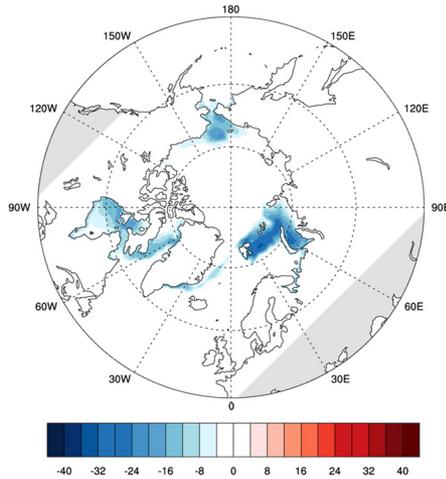
(stronger influence of Siberian snow)

Response estimated by multiple regression (similar to regular regression)

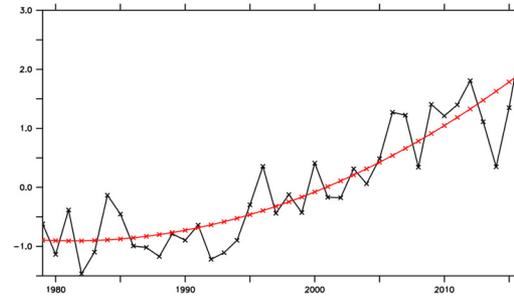
Significant warming in northeast America, weak cooling in March in northern Europe

Estimation of the direct atmospheric impact of the sea ice loss

Assuming that the direct response to the sea ice loss is given by that to the interannual fluctuations, scaling leads to a strong estimated impact between 1979 and 2016

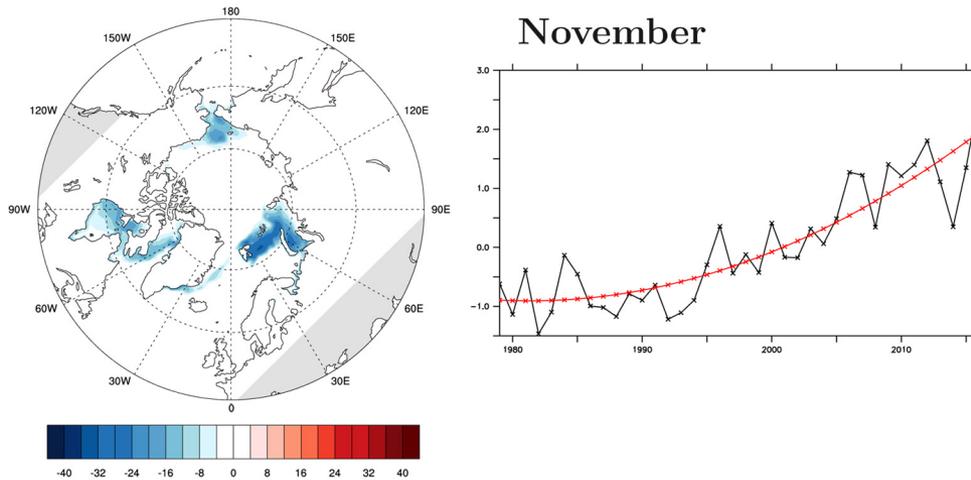


November



Estimation of the direct atmospheric impact of the sea ice loss

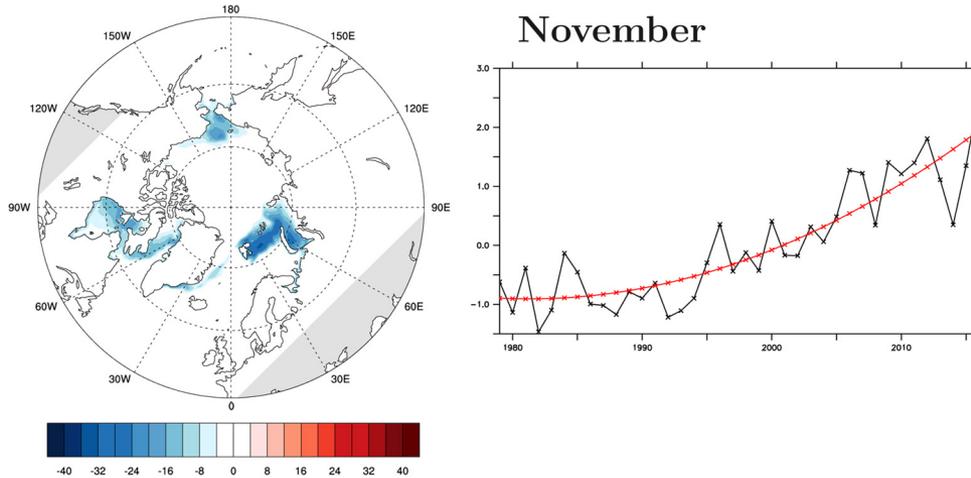
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- **The NAO has become more negative**
- Z500 has increased in January and February by up to 65 m over the subpolar gyre and decreased by 40 m in the subtropical North Atlantic
- **SAT has increased by up to 2.5 K in northeastern North America**
- SAT has decreased in March by up to 2 K over northern Europe
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Such large changes are masked by the response to increasing GHG concentration, SST, decreasing snow cover and indirect sea ice influence

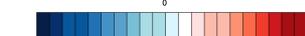
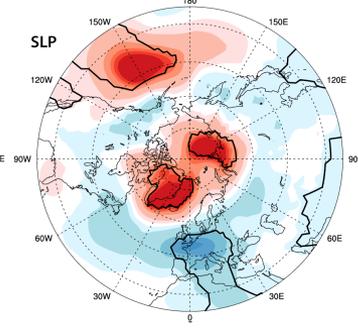
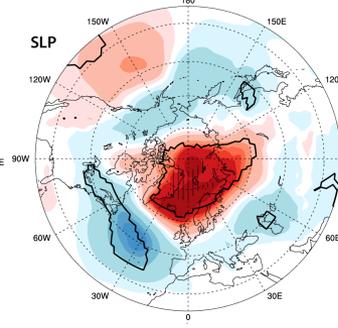
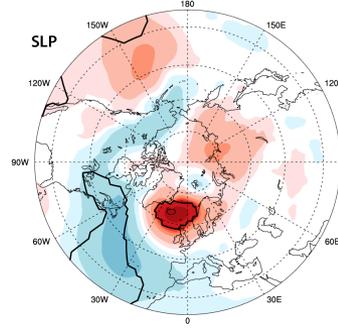
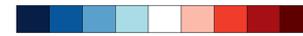
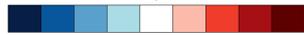
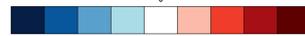
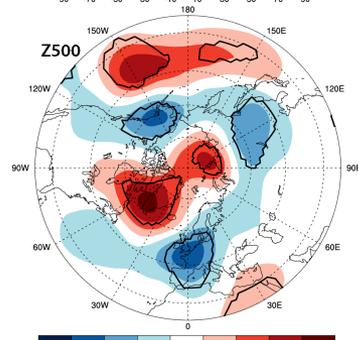
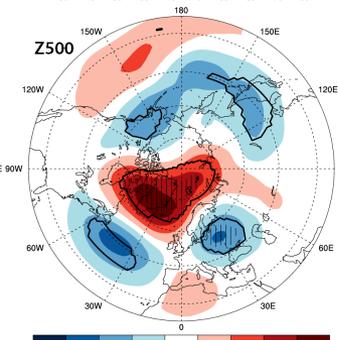
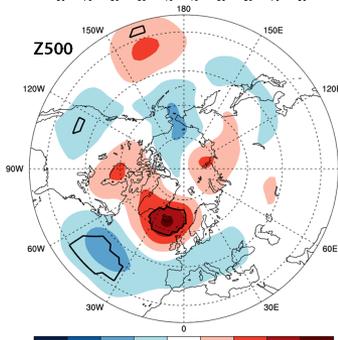
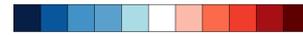
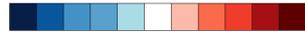
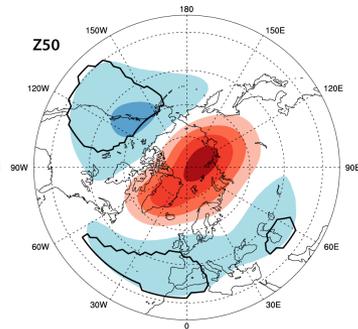
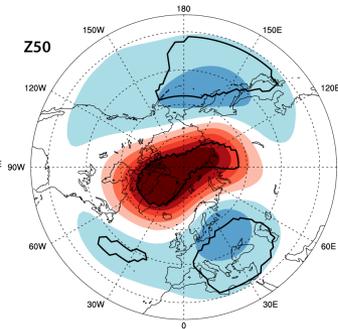
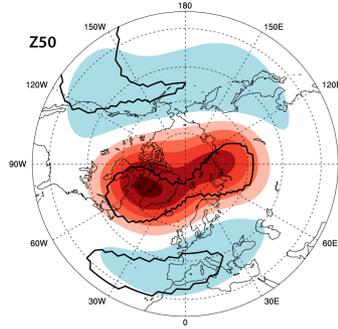
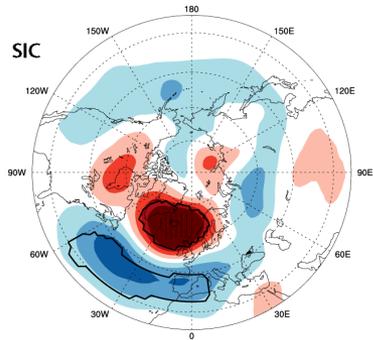
Lag regression on standardized interannual SIC fluctuations in December

January (lag 1)

February (lag 2)

March (lag 3)

Multiple regression
in January on SIC



No field significant
stratospheric signal

The negative NAO
is mostly due to SIC

Black line
10% significance

Hatching
FDR of 10%

Summary

- Interannual SIC fluctuations with the sea ice loss pattern are followed by a NAO- from December to March (stronger in January and February)
- The stratospheric AO- signal is primarily due to Siberian snow cover

(the sea ice loss influences the wintertime atmospheric circulation primarily via tropospheric processes; the maximum response is found at 2-month lag)
- There is a fast warming above sea ice retreat in a warming of northeastern North America, but the WACE pattern is not due to the sea ice loss, except perhaps in March

Summary

- Interannual SIC fluctuations with the same pattern as the sea ice loss in November and December are followed by a negative NAO- in the troposphere from December to March (stronger in January and February)

- The stratospheric AO- signal is primarily due to Siberian snow cover

(the sea ice loss influences the wintertime atmospheric circulation primarily via tropospheric processes; the maximum response is found at 2-month lag)

- There is a fast warming above sea ice retreat, but no clear evidence that the WACE pattern is due to the variability of the sea ice loss pattern

(WACE seems to be in part driven by a sea ice seesaw between the Barents-Kara Seas and the Greenland Sea, see Mori et al. 2019, not by the *overall* SIC decrease)

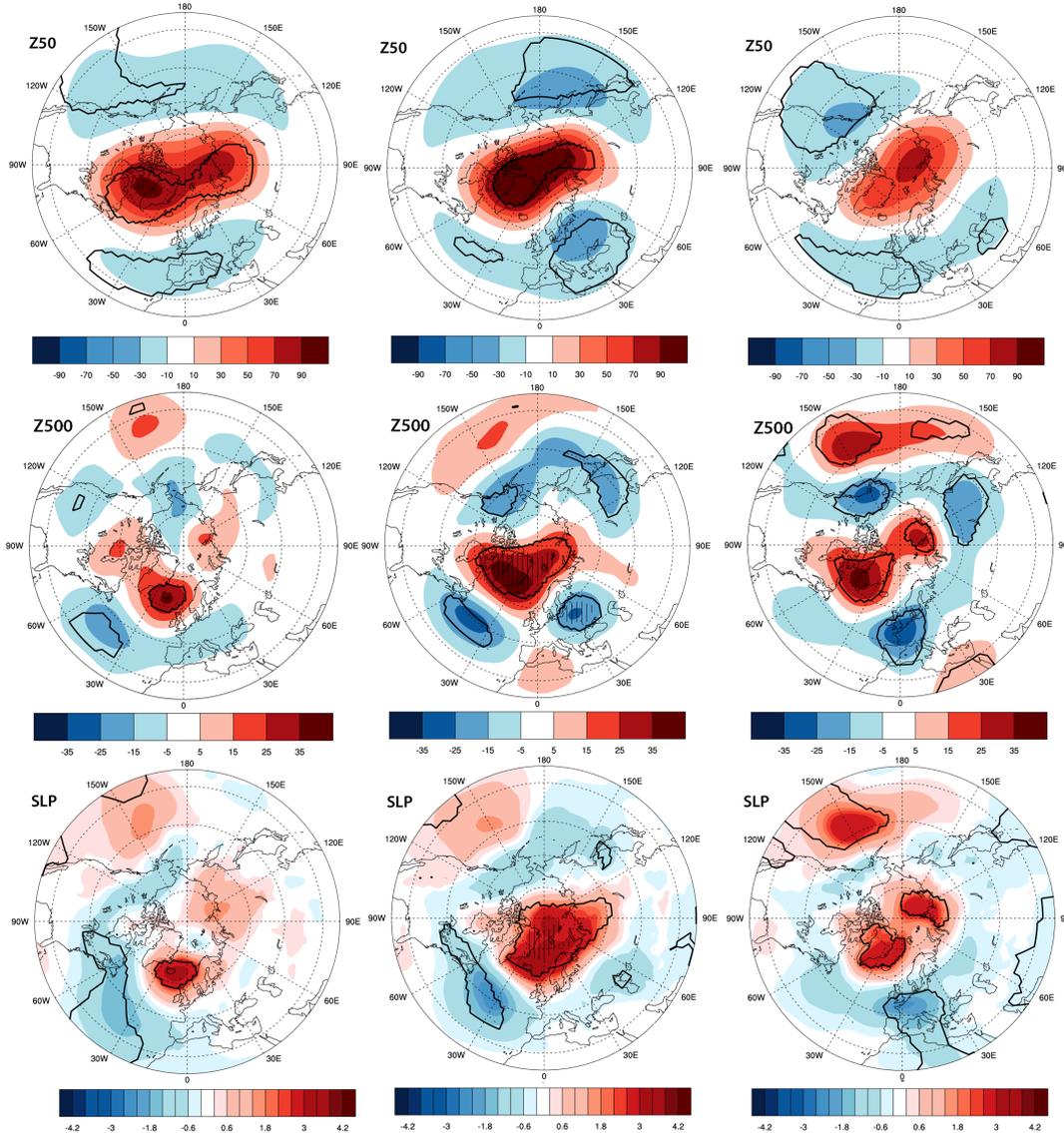
- No significant large-scale impact was found in October and November

Lag regression on standardized interannual SIC fluctuations in December

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February (lag 2)

March (lag 3)



No field significant stratospheric signal

Clear NAO-tropospheric signal (strongest at lag 2)

Black line
10% significance

Hatching
FDR of 10%