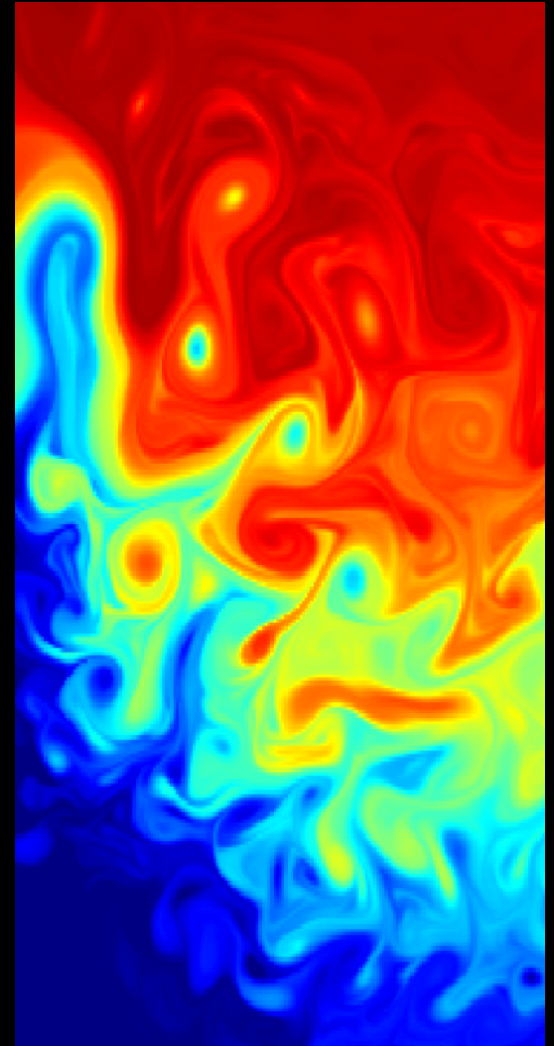


# The role of closed gyres in the zonal transport of the ACC

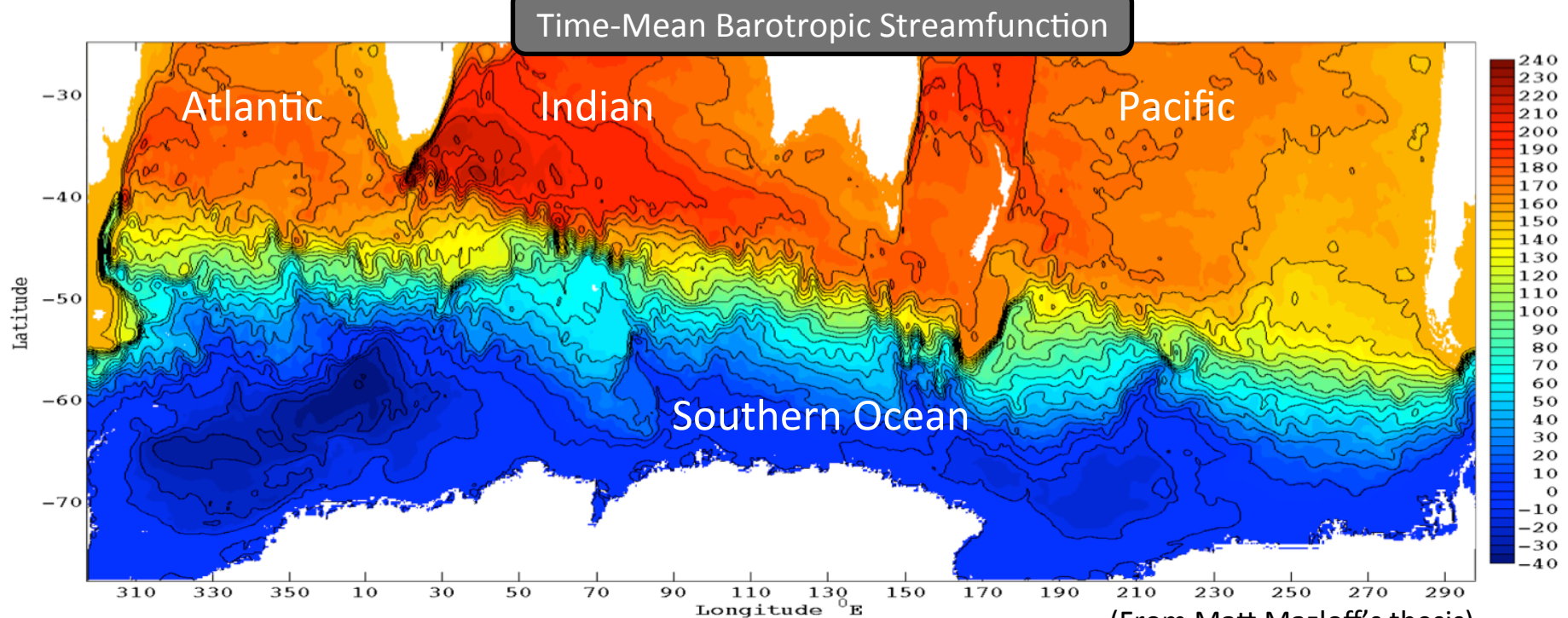
**L.-P. Nadeau and R. Ferrari**



# INTRODUCTION

## Antarctic Circumpolar Current

- Most voluminous ocean current (transport= 130 Sv)
- Global importance: mixes together Atlantic, Indian and Pacific Oceans waters
- **Still lack a robust quantitative theory for what sets the magnitude of the circumpolar transport**

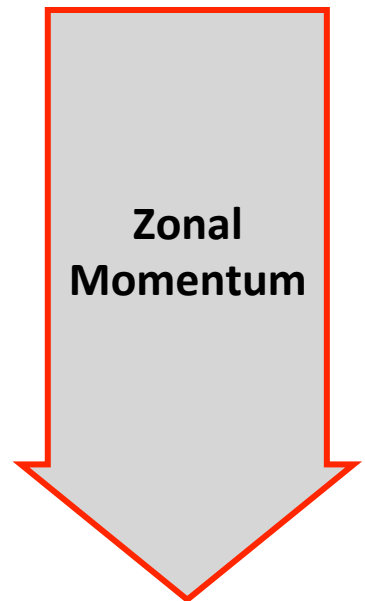


# INTRODUCTION

## Consensus:

Transport = Balance of 3 terms in Zonal Momentum Budget

- Surface → *Input* by the wind
- Interior → *Downward transport* by geostrophic eddies
- Bottom → *Sink* by topographic form drag



# INTRODUCTION

Focus of most recent studies

- Surface → *Input* by the wind
- Interior → *Downward transport* by geostrophic eddies
- Bottom → *Sink* by topographic form drag

↑  
Explains Transport  
↓

Zonal  
Momentum

# INTRODUCTION

Focus of most recent studies

- Surface →

*Input* by the wind



Explains Transport



- Interior →

*Downward transport*  
by geostrophic eddies

Zonal  
Momentum

**Implicit assumption:** *topographic drag is  
slaved to balance the  
downward eddy transport of momentum*

## Questions:

- Does the topographic form drag exert an active control on the (baroclinic) transport?  
➡ *yes*
- Is the effect of the topography on the ACC better thought of as a local or a global constraint?  
➡ *global*

## Questions:

- Does the topographic form drag exert an active control on the (baroclinic) transport?  
→ *yes*
- Is the effect of the topography on the ACC better thought of as a local or a global constraint?

→ *global*



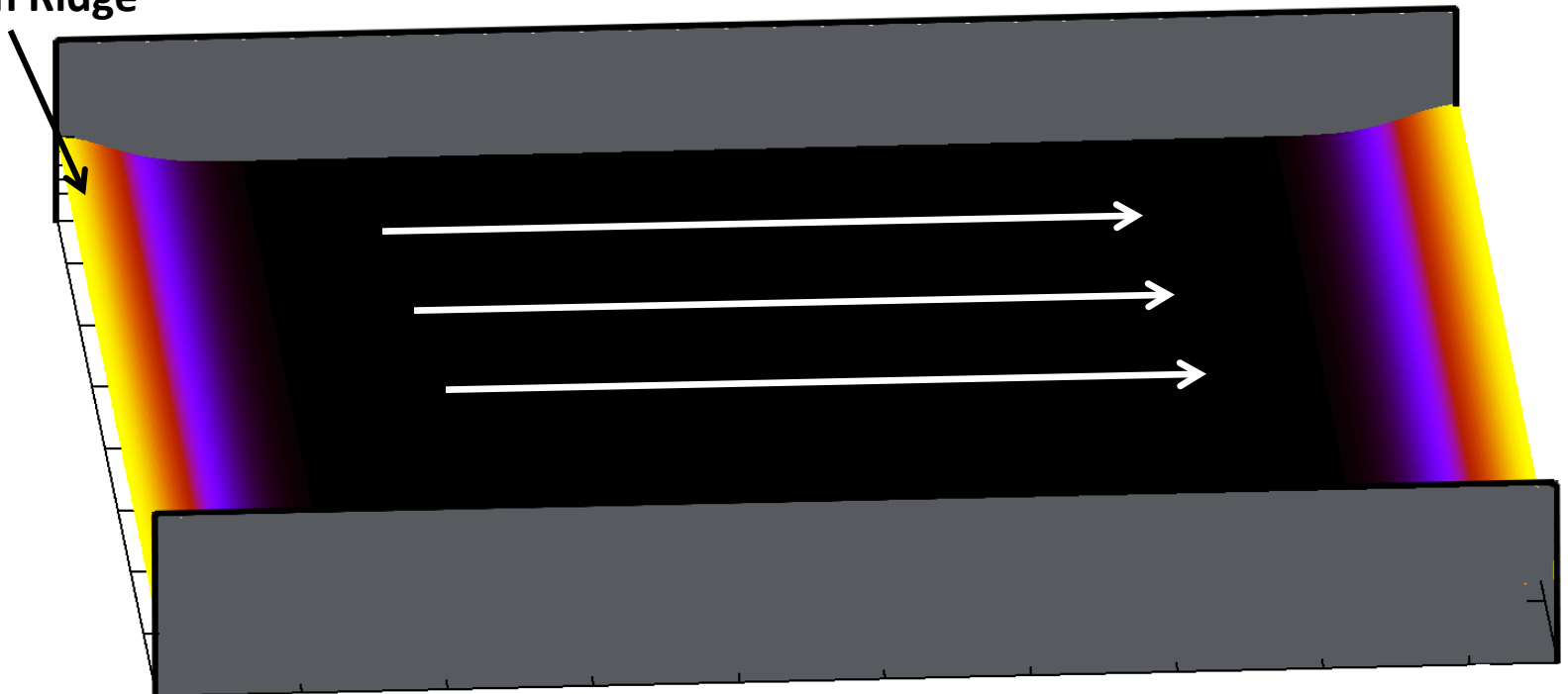
*closed gyres*

# EXPERIMENTAL SETUP

1. Closed box with topography
2. Channel with topography
3. Flat bottom Channel

- **QG** and **Primitive Eq.** numerical simulations
- Wind forcing but **no buoyancy forcing**

Gaussian Ridge





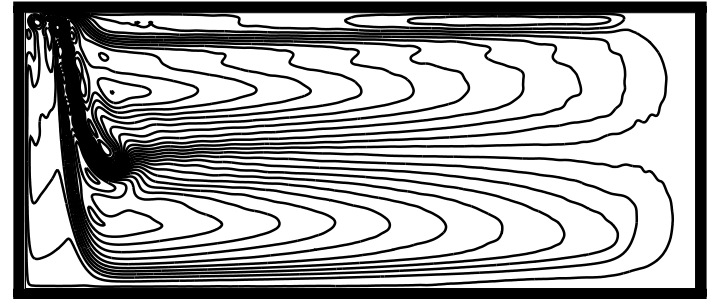
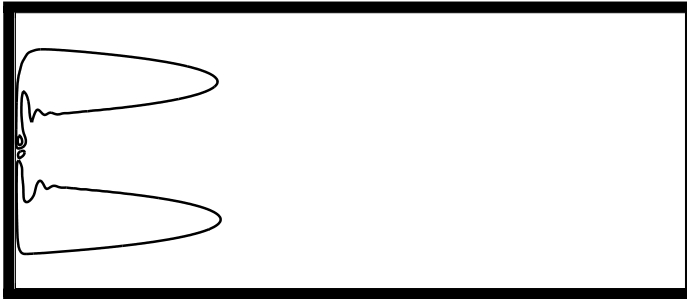
# CLOSED GYRES

Barotropic Streamfunction

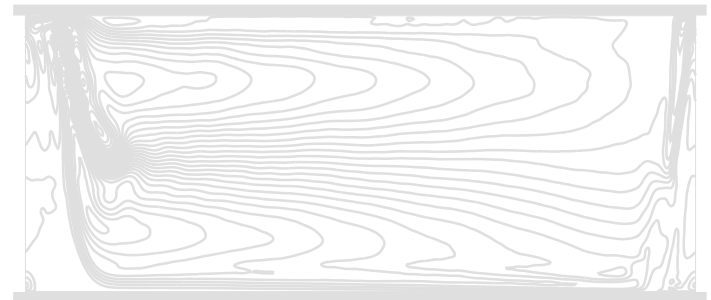
Weak Forcing

Strong Forcing

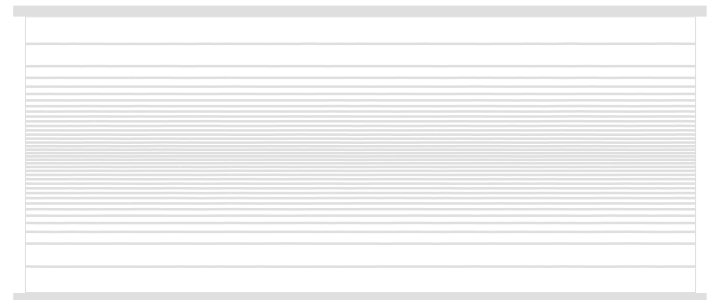
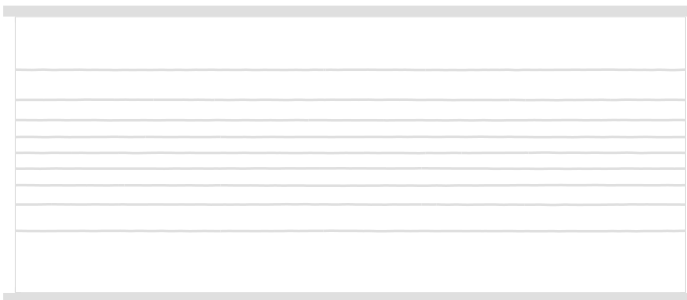
Box  
+ Ridge



Channel  
+ Ridge



Channel  
Flat

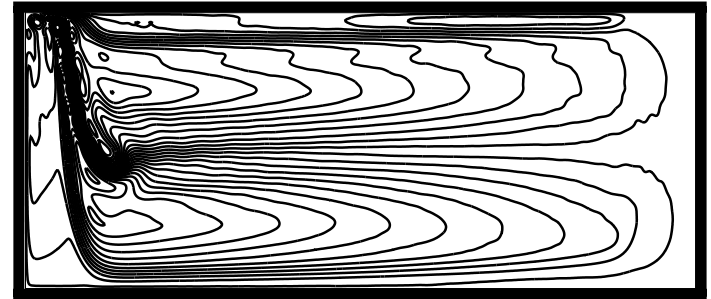
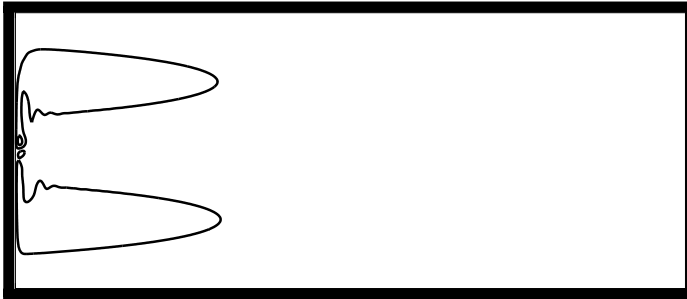


# CLOSED GYRES

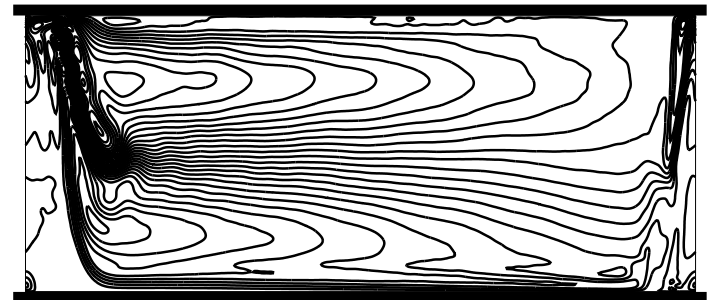
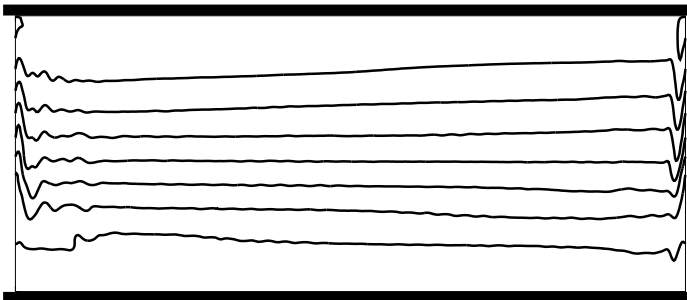
Weak Forcing

Strong Forcing

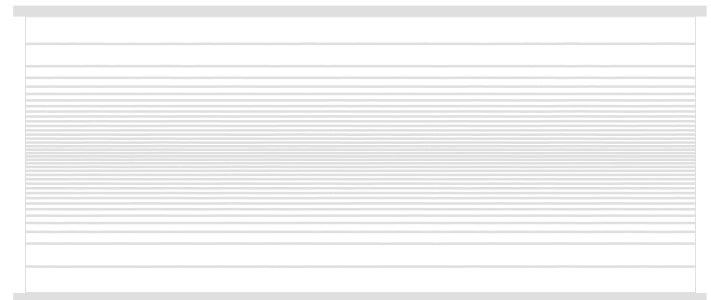
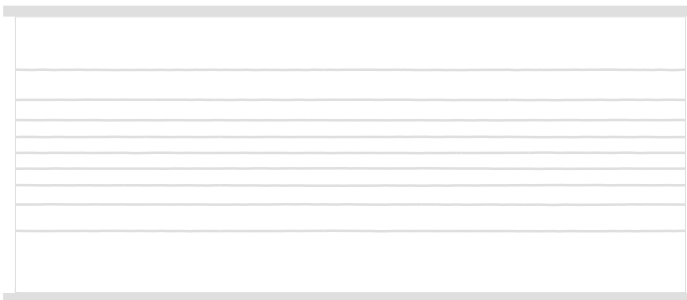
Box  
+ Ridge



Channel  
+ Ridge



Channel  
Flat

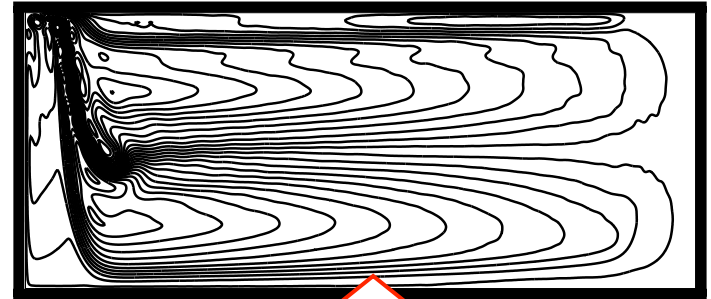
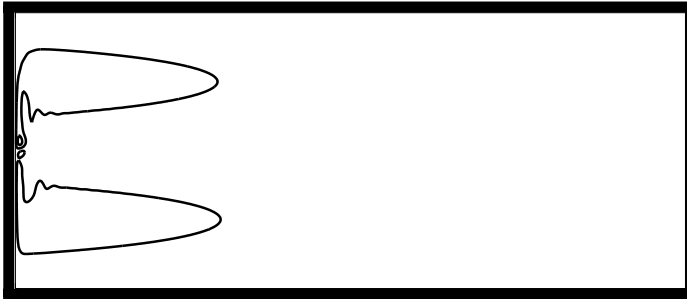


# CLOSED GYRES

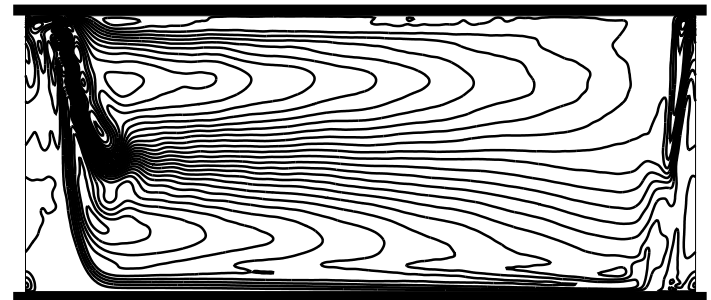
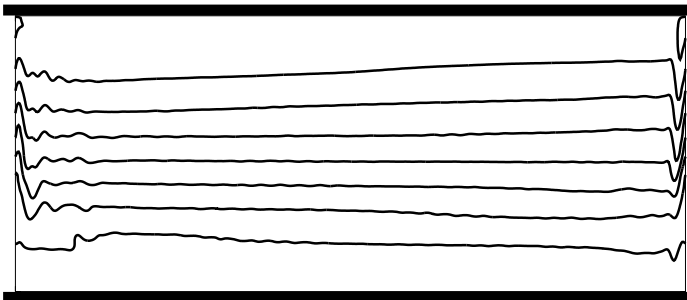
Weak Forcing

Strong Forcing

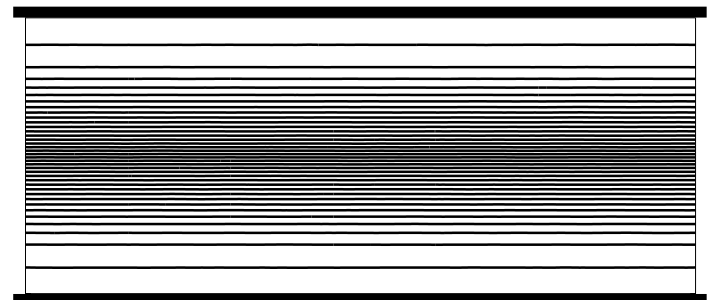
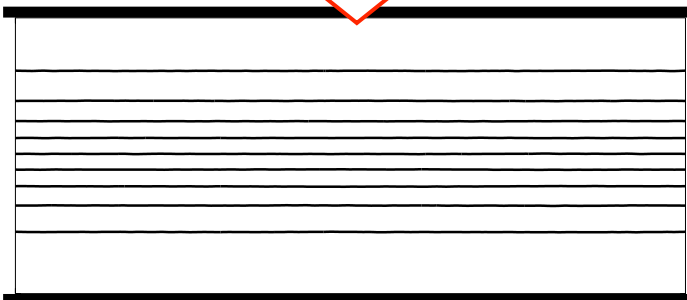
Box  
+ Ridge



Channel  
+ Ridge

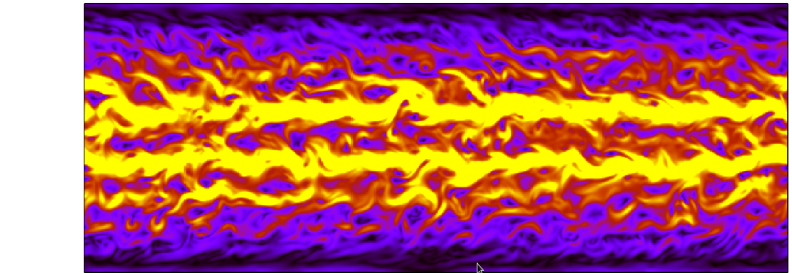
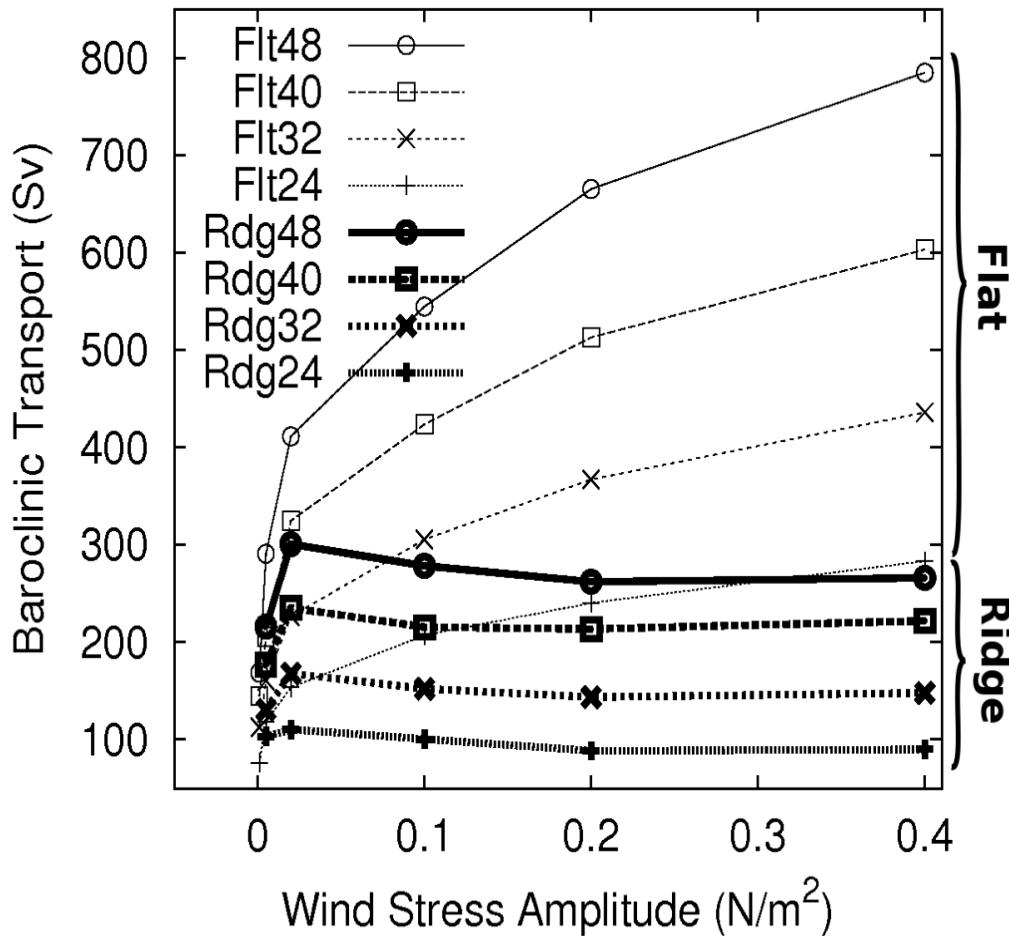


Channel  
Flat



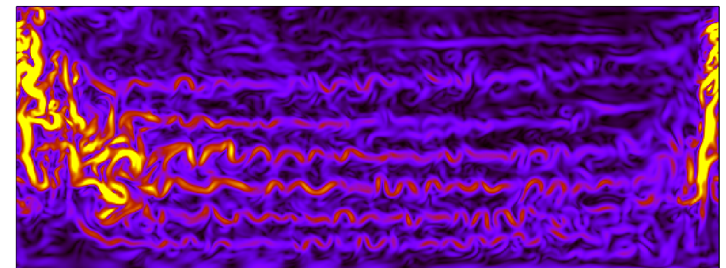
# COMPARISON FLAT AND RIDGE

## Baroclinic Transport



Flat Bottom:  $T \sim \tau^{0.2}$

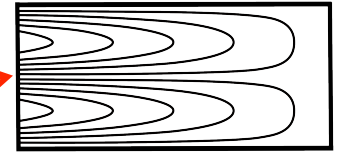
Gaussian Ridge: Saturated Transport



# CLOSED GYRES

## Gaussian Ridge: Saturated Transport

$$\psi_{SV}(x, y) = -\frac{1}{\rho_0 H \beta} \int_x^{L_x} \text{curl}_z \boldsymbol{\tau} \, dx$$



Wind Forcing

$\tau_0=0.02$

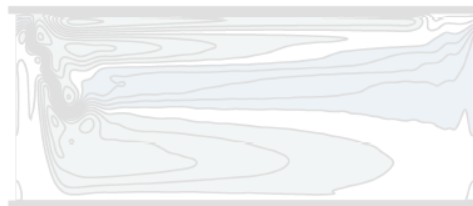
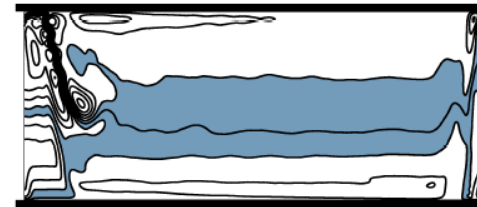
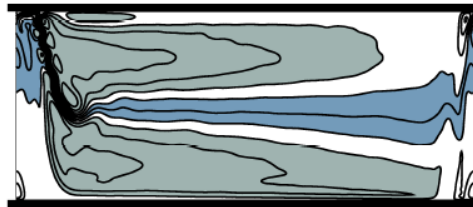
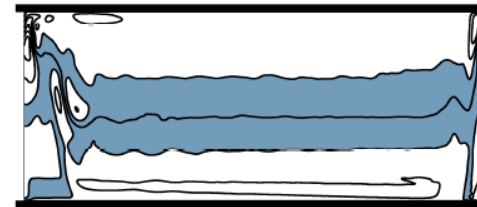
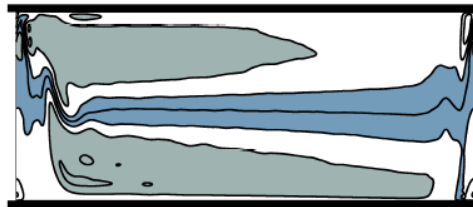
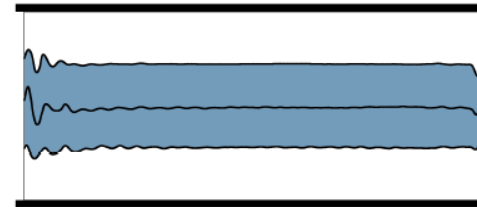
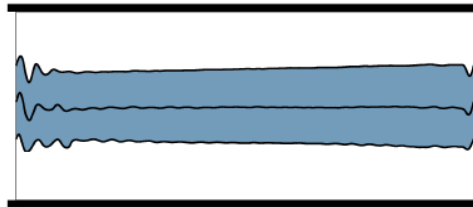
$\tau_0=0.2$

$\tau_0=0.4$

Primitive  
Equations  
 $\tau_0=0.4$

$\Psi_{BT}$

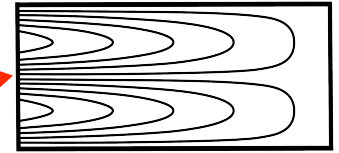
$\Psi_{diff} = \Psi_{BT} - \Psi_{SV}$



# CLOSED GYRES

## Gaussian Ridge: Saturated Transport

$$\psi_{SV}(x, y) = -\frac{1}{\rho_0 H \beta} \int_x^{L_x} \text{curl}_z \boldsymbol{\tau} \, dx$$



Wind Forcing

$\tau_0=0.02$

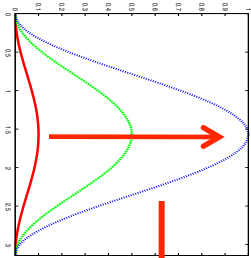
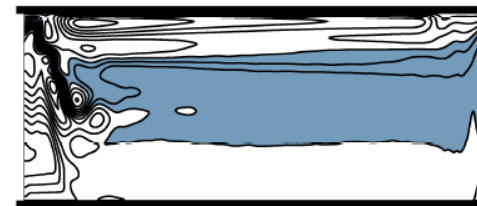
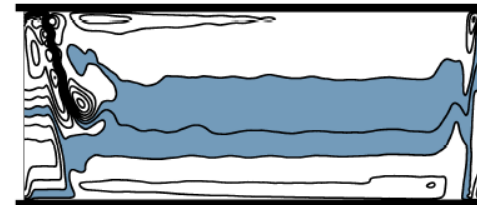
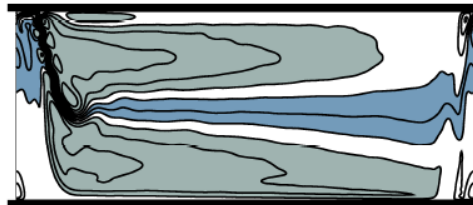
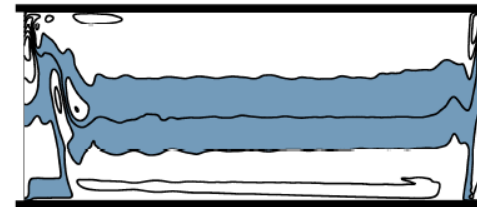
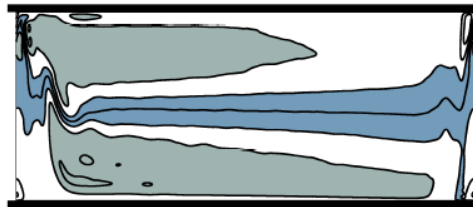
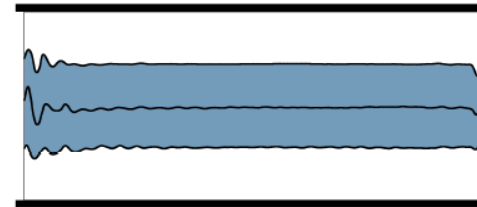
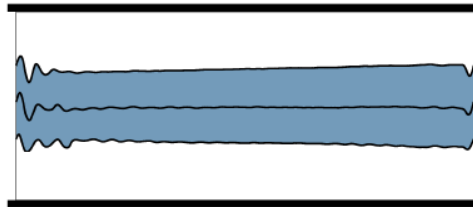
$\tau_0=0.2$

$\tau_0=0.4$

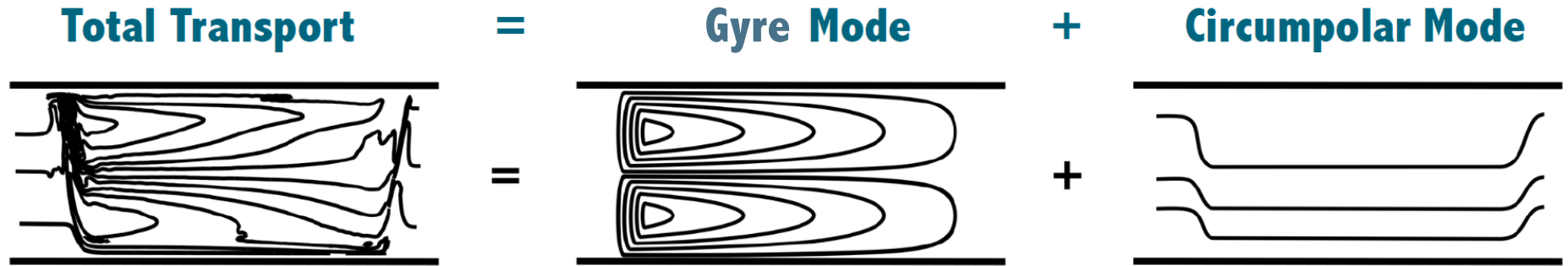
Primitive  
Equations  
 $\tau_0=0.4$

$\Psi_{BT}$

$\Psi_{diff} = \Psi_{BT} - \Psi_{SV}$



## Hypothesis



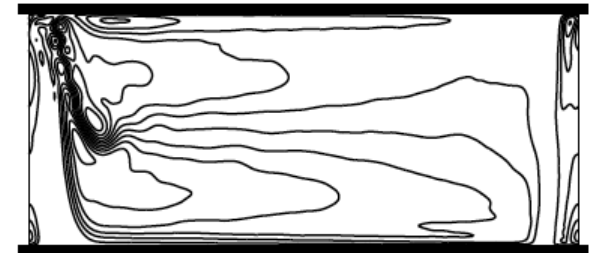
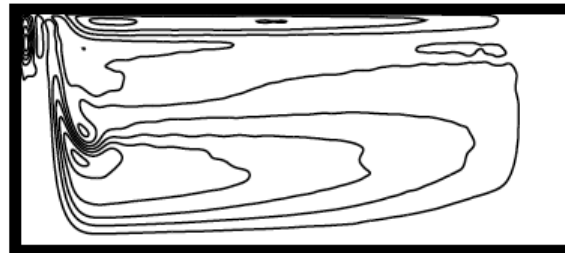
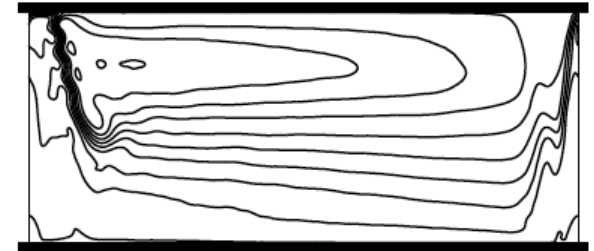
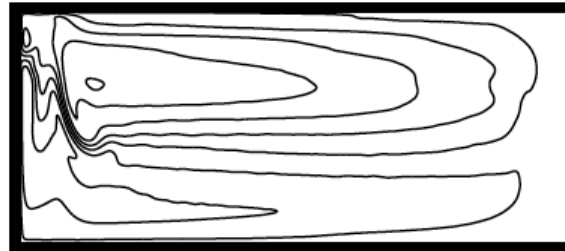
**Total barotropic streamfunction** can be decomposed in a "gyre mode" (Sverdrup gyres), contributing no barotropic transport, and a "circumpolar mode", contributing all barotropic transport.

# ROLE OF EDDIES

## Vertical Structure of the Sverdrup Flow

Box + Ridge

Channel + Ridge



**Bottom circulation:**  
driven by vortex  
stretching associated  
with transient eddies

$\psi_1$

$\psi_2$

**Physical Mechanism:**

**Eddy driven** bottom circulation  
sustains **topographic form drag**  
that balances wind stress  
and causes transport saturation



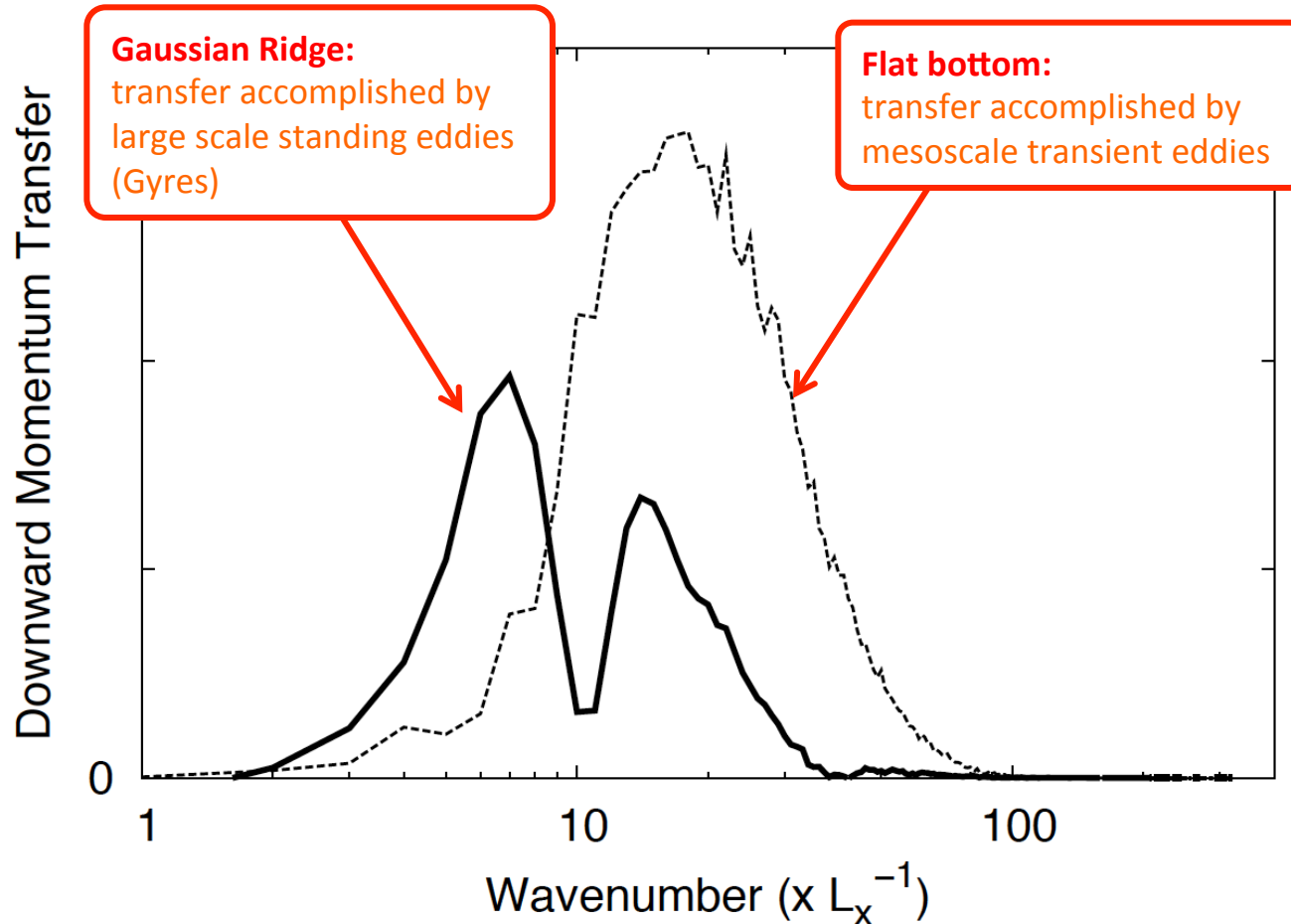
# EDDY MOMENTUM TRANSFER

Dominant zonal momentum balance at statistical equilibrium

$$\begin{array}{lcl}
 & \text{Reynolds} & \text{Interfacial} & \text{Wind} \\
 & & & \downarrow \\
 \text{Upper:} & \langle \psi_{1x} \psi_{1yy} \rangle - \frac{f_0^2}{g' H_1} \langle \psi_2 \psi_{1x} \rangle & \leftarrow \frac{1}{\rho_0 H_1} \langle \tau \rangle & = 0 \\
 & & \downarrow & \\
 \text{Lower:} & \langle \psi_{2x} \psi_{2yy} \rangle - \frac{f_0^2}{g' H_2} \langle \psi_1 \psi_{2x} \rangle & \rightarrow \frac{f_0}{H_2} \langle \psi_{2x} h_b \rangle & = 0 \\
 & & & \downarrow \\
 & & & \text{Topographic}
 \end{array}$$

# EDDY MOMENTUM TRANSFER

Spectral decomposition according to zonal wavenumber  
of the interfacial form stress term



# EFFECT OF THE RIDGE HEIGHT

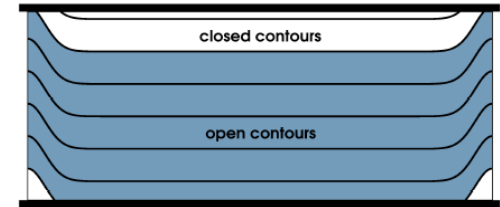
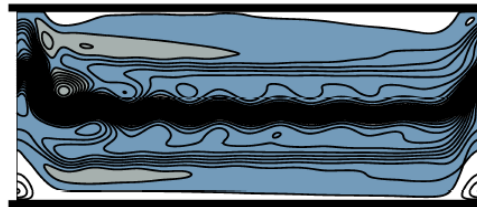
Why do we observe gyres in an open channel?

Increasing height

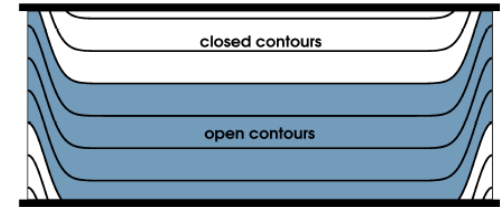
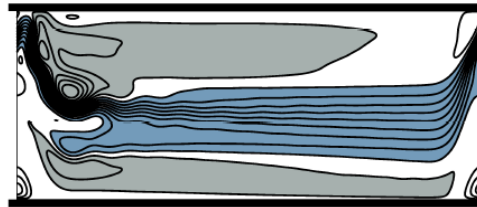
$\psi_{Bt}$

$\beta y + (f_0/H)h_b$

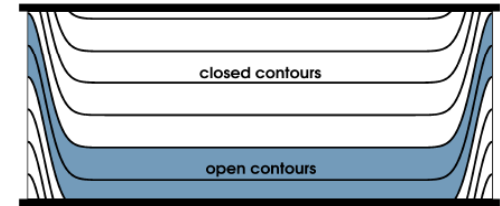
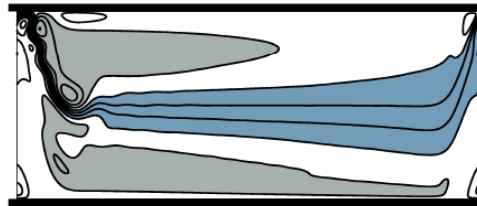
$h_0 = 425\text{m}$



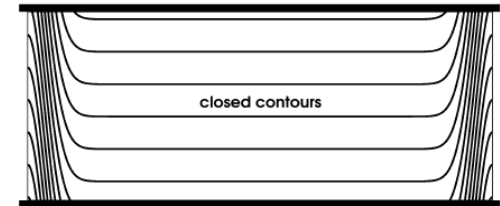
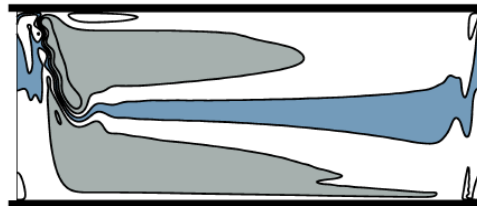
$h_0 = 850\text{m}$



$h_0 = 1275\text{m}$



$h_0 = 2600\text{m}$

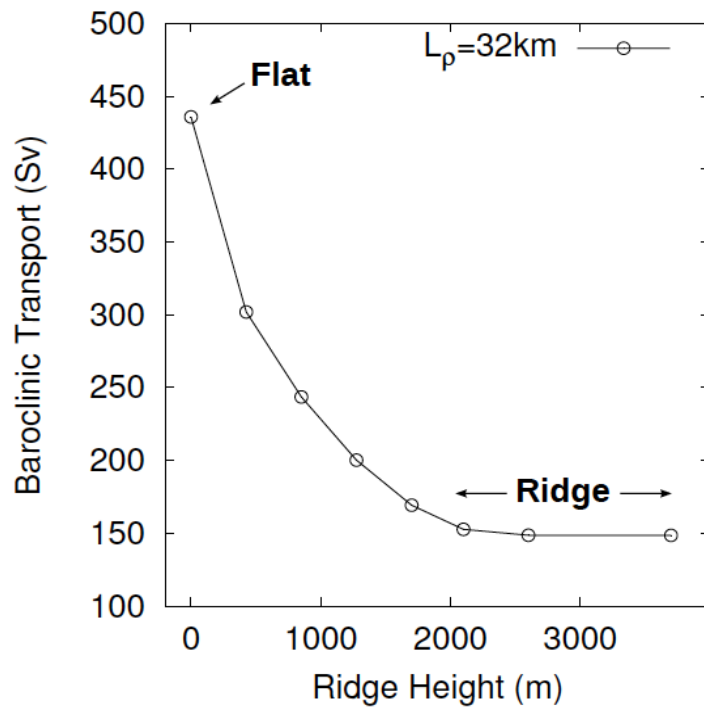


If topography is sufficiently steep to bring geostrophic contours close together, a frictional boundary layer develops that acts as an “effective wall”.

# EFFECT OF THE RIDGE HEIGHT

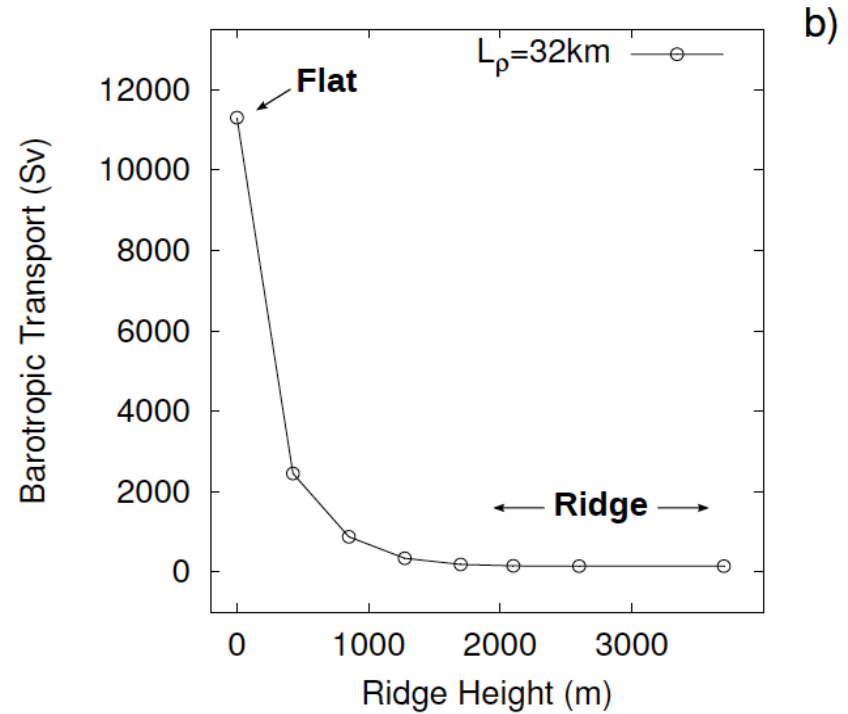
## Effect of the Ridge height on the transport

### Baroclinic



a)

### Barotropic



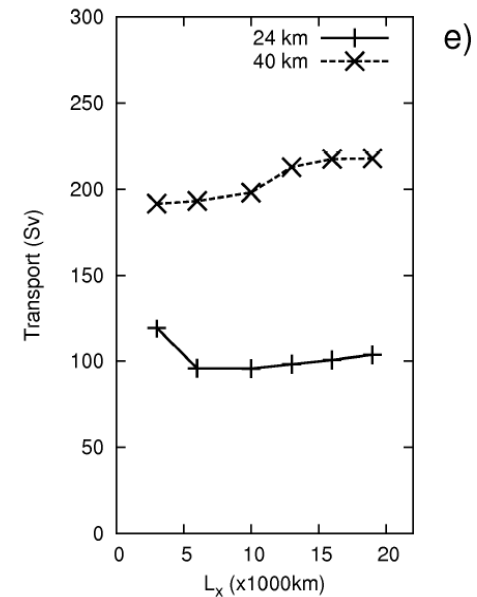
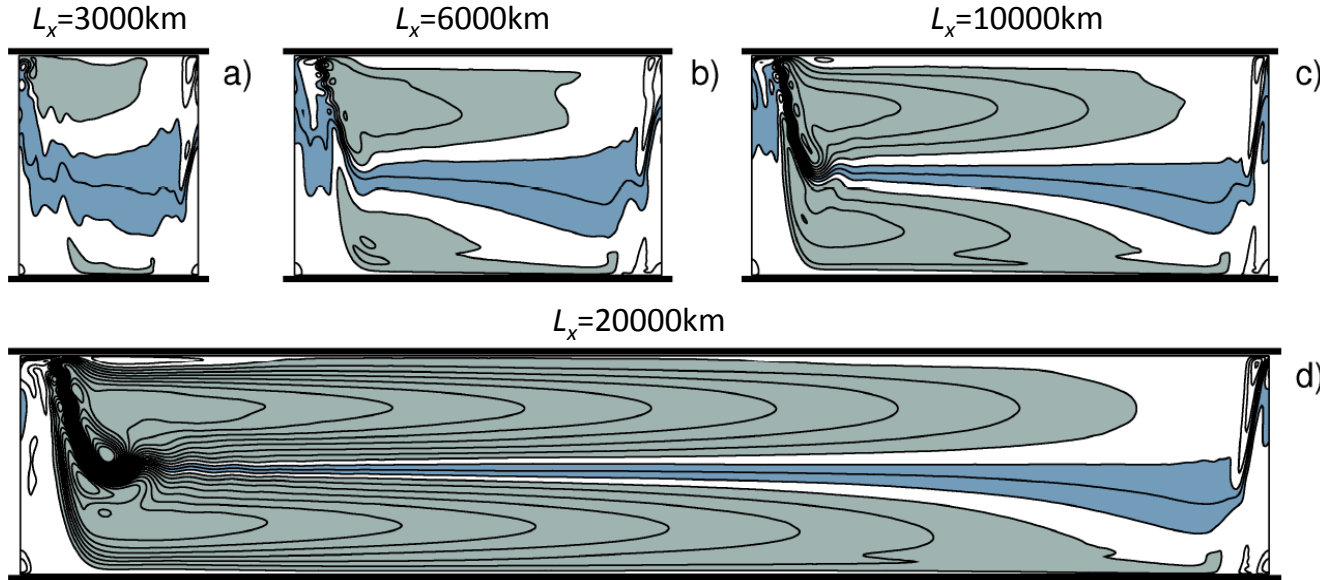
b)

# EFFECT OF THE CHANNEL LENGTH

Varying Channel Length: modes can be decoupled

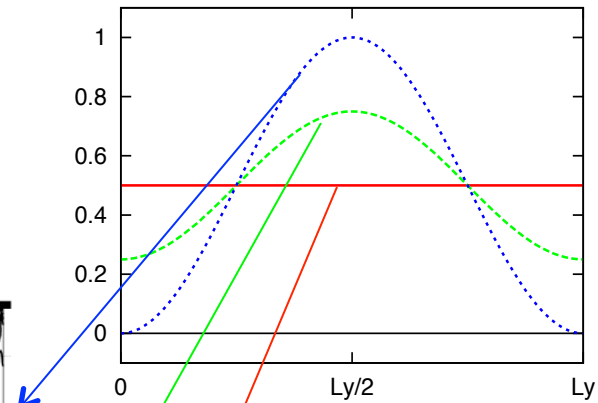
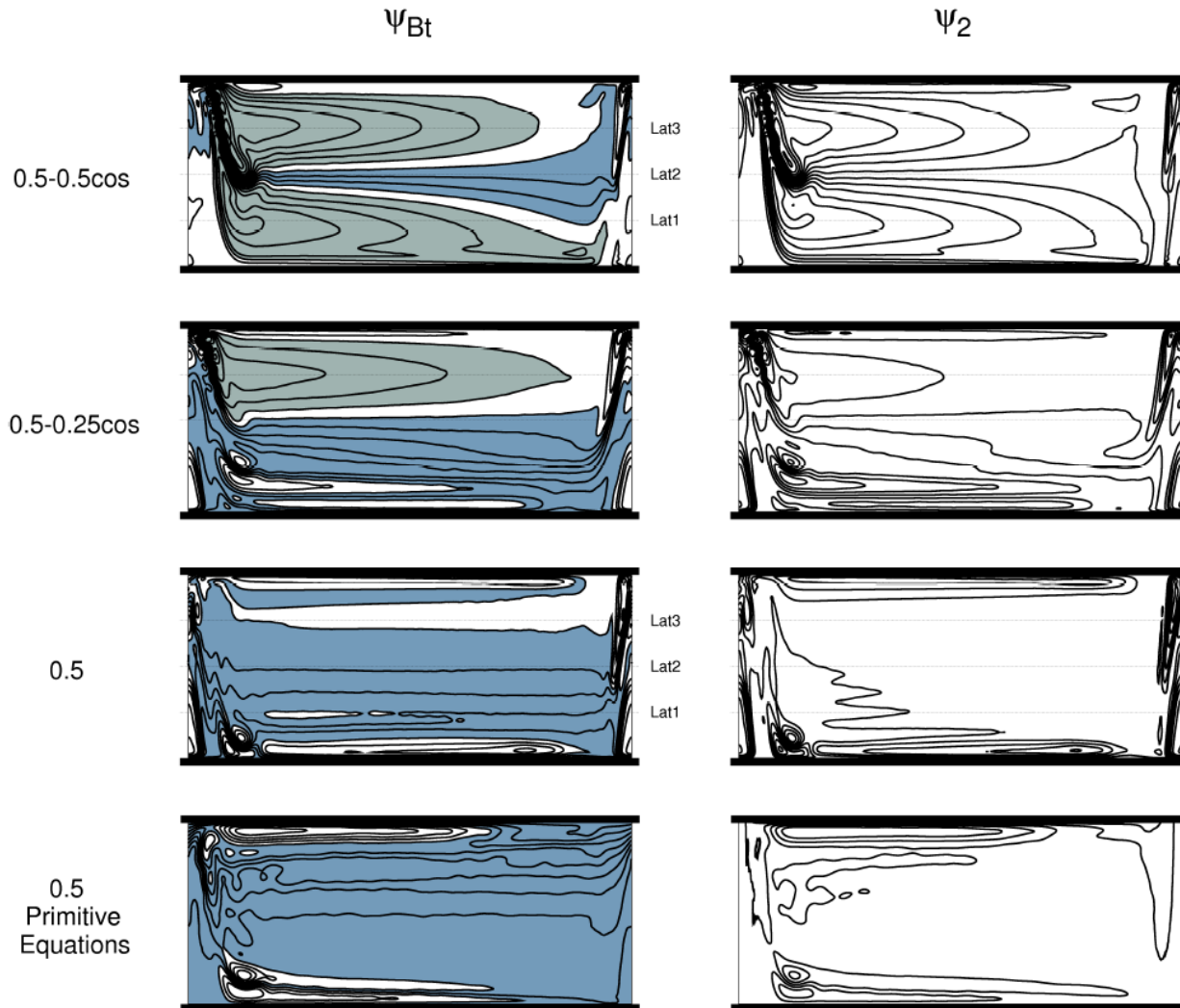
**Gyre mode** increases increases with  $L_x$

**Circumpolar mode** independent of  $L_x$



# EFFECT OF THE WIND STRESS CURL

Varying wind stress curl  
for a fixed averaged wind stress

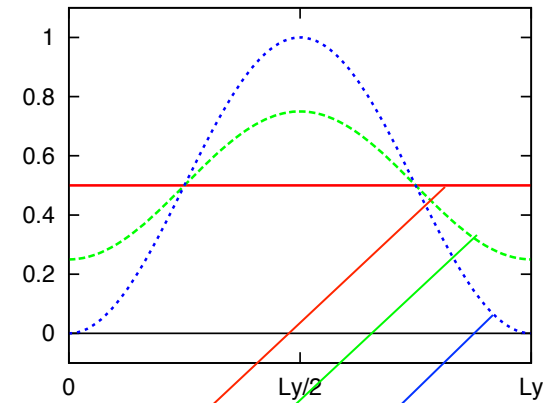
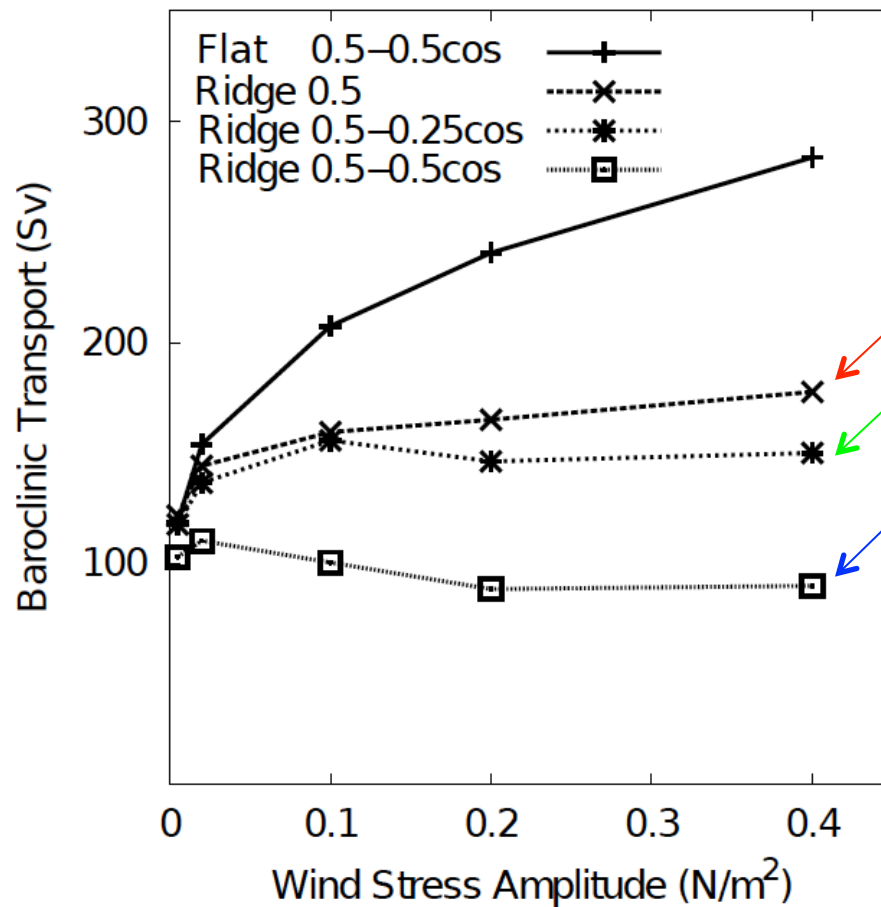


Less bottom  
circulation

Less  
Topographic  
form drag

# EFFECT OF THE WIND STRESS CURL

Varying wind stress curl  
for a fixed averaged wind stress



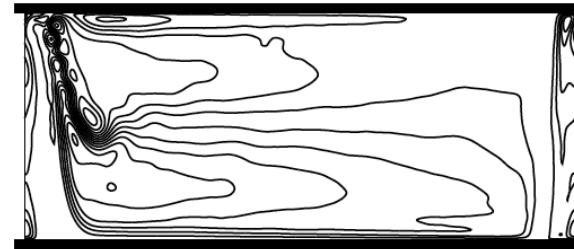
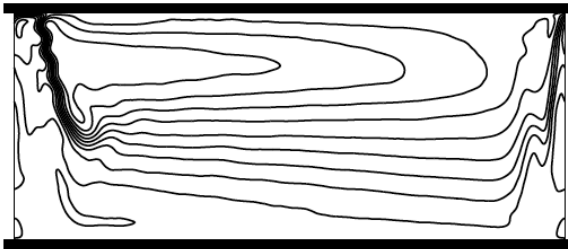
# EFFECT OF THE BOTTOM FRICTION

## Bottom Friction

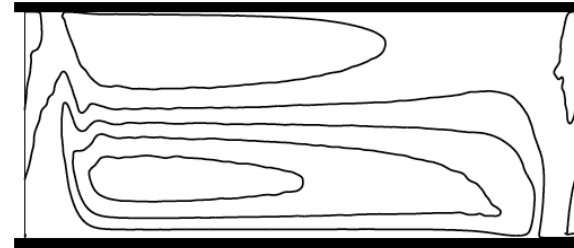
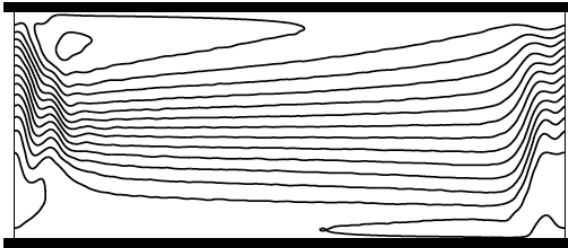
$\Psi_1$

$\Psi_2$

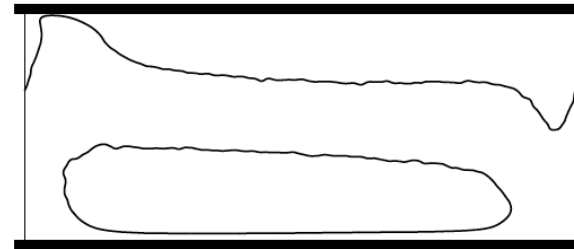
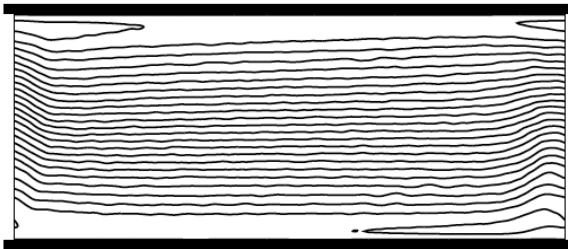
$r = 10^{-7}$



$r = 10^{-6}$



$r = 10^{-5}$

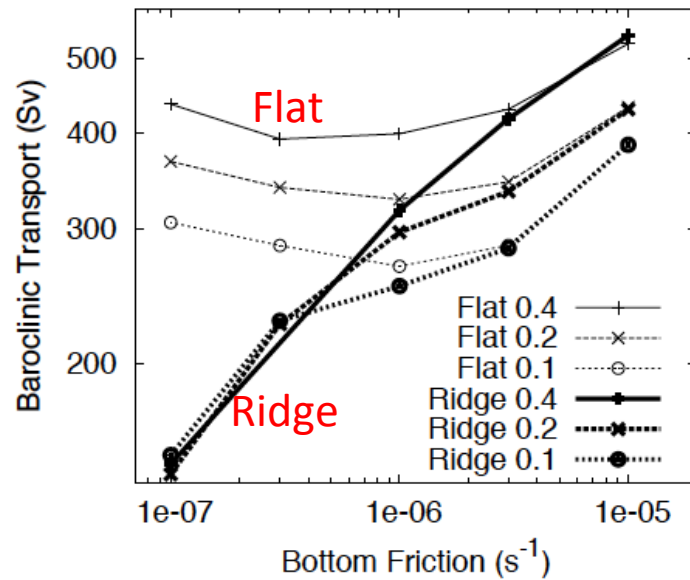




# EFFECT OF THE BOTTOM FRICTION

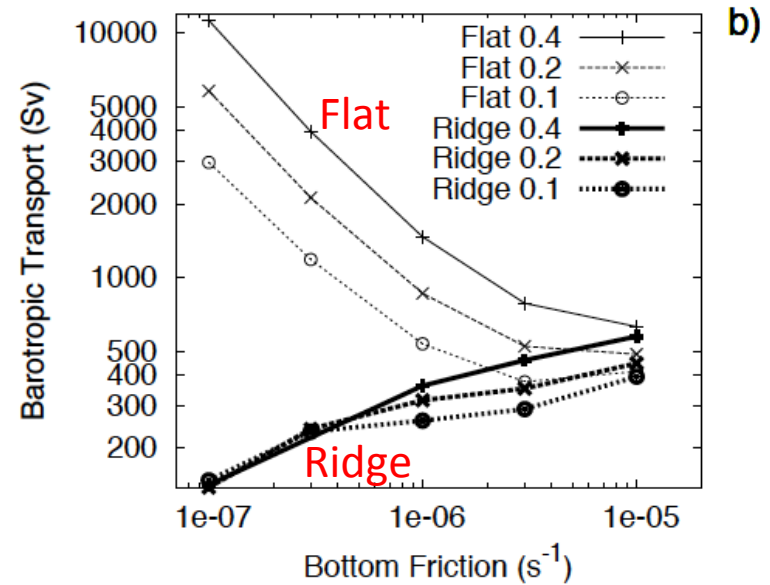
## Bottom Friction

### Baroclinic



a)

### Barotropic



b)

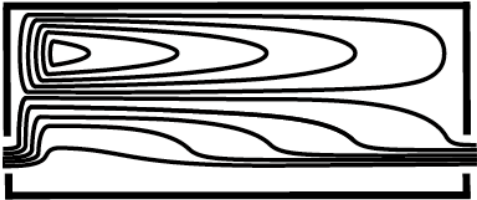
# SUMMARY

- Closed **recirculating gyres** develop in the lee of major topographic ridges.
- Circulation can be decomposed in a **gyre mode** and a **circumpolar mode**.
- Each **mode can be decoupled** (can favor one mode without affecting the other).
- However, specific effect of the form drag generated by the gyre mode onto the circumpolar mode is still not clearly defined.

# SUMMARY

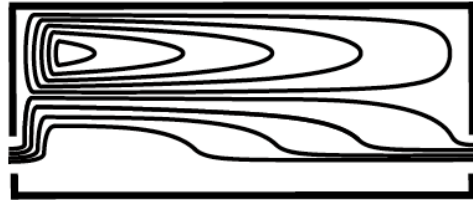
## Initial hypothesis

Total Transport (Stommel Regime)



=

Basin Contribution

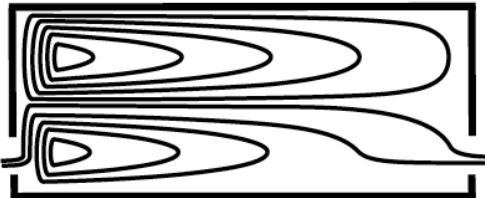


+

Channel Contribution

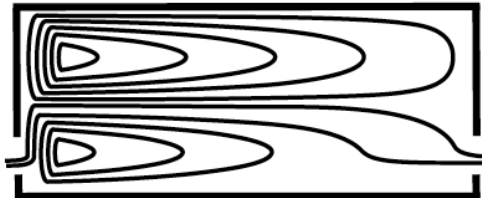


Total Transport (Saturation Regime)



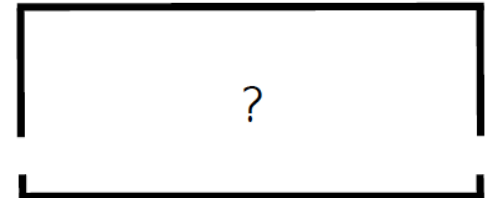
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Basin Contribution



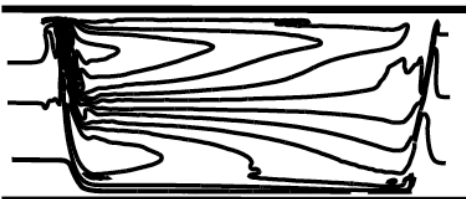
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Channel Contribution



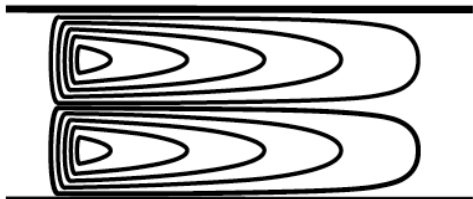
## Revised hypothesis

Total Transport (Saturation Regime)



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Basin Mode



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Circumpolar Mode

